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BULLETIN 136

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GEOLOGY AND GROUND-WATER RESOURCES  
OF CLAY COUNTY, KANSAS

By KENNETH L. WALTERS and CHARLES K. BAYNE

*Prepared by the United States Geological Survey and the State Geological Survey of Kansas, with the cooperation of the Division of Sanitation of the Kansas State Board of Health, and the Division of Water Resources of the Kansas State Board of Agriculture*



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# GEOLOGY AND GROUND-WATER RESOURCES OF CLAY COUNTY, KANSAS

By KENNETH L. WALTERS and CHARLES K. BAYNE

## ABSTRACT

This report describing the geography, geology, and ground-water resources of Clay County, in north-central Kansas, is based on hydrologic and geologic information obtained in the field during the fall of 1953 and the summer of 1954. The field data are given in tables; they include records of 143 wells, chemical analyses of water from 35 representative wells, logs of 28 test holes, and results of pumping (aquifer) tests. In addition, 110 holes were augered to determine the depth to water.

Clay County has an area of about 660 square miles and lies in the Great Plains and Central Lowlands physiographic provinces. Most of it is drained by Republican River; small areas are drained by Fancy Creek and Chapman Creek, which are not tributaries of Republican River. The normal annual precipitation at Clay Center is 27.92 inches, and the mean annual temperature is 55.1°F. Agriculture is the principal occupation in the county.

The rocks that crop out at the surface in Clay County are sedimentary and range in age from Permian to Recent. The oldest formation exposed in the county is the Barneston Limestone. The Dakota Formation, the youngest Cretaceous rock in the county, crops out over a large part of western and northern Clay County. The Permian and Cretaceous rocks are mantled in many places by unconsolidated continental deposits of fluvial and eolian origin representing four stages of the Pleistocene Epoch.

The unconsolidated sand and gravel deposits of Pleistocene age form the principal aquifer in the county. These deposits are thickest and most extensive in the valley of Republican River. The Dakota Formation yields moderate quantities of water to wells in western and northern Clay County. The Barneston Limestone yields small to moderate quantities of water to wells in eastern Clay County. Ground water in the area is recharged principally from local precipitation; underflow from adjacent areas contributes significantly, however. Ground water is discharged mainly by seepage into streams and by transpiration by plants. All municipal and industrial water supplies and most domestic and stock supplies are obtained from wells. Irrigation from wells is practiced extensively in the valley of Republican River.

## INTRODUCTION

### PURPOSE AND SCOPE OF INVESTIGATION

A program of investigation of the ground-water resources of Kansas was begun in 1937 by the United States Geological Survey and the State Geological Survey of Kansas with the cooperation of the Division of Sanitation of the State Board of Health and the Division of Water Resources of the State Board of Agriculture. The investigation of that part of Clay County that lies within the Republican

Valley was integrated with the program of the Interior Department for development of the Missouri River basin. The investigation upon which this report is based was begun in the fall of 1953 and field work was completed in the fall of 1955. It is similar to other investigations that have been completed or are being made in other counties in Kansas. The present status of investigations resulting from this program is shown in Figure 1.

Ground water is one of the principal natural resources of Clay County. Nearly all public, domestic, and industrial water supplies and many stock supplies are obtained from wells. Ground water is being used to some extent for irrigation, and recent interest in irrigation indicates that the use of ground water for this purpose probably will increase greatly in the future. Withdrawal at the present rate has not seriously depleted the ground-water supply or resulted in impairment of the quality of the water by encroachment of water of poor quality, but there is need for an adequate understanding of the quality and quantity of ground water available.

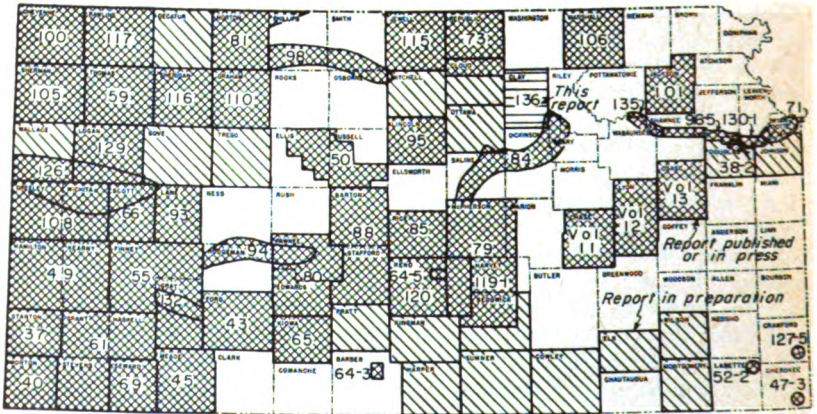


FIG. 1.—Index map of Kansas showing area discussed in this report and other areas for which ground-water reports have been published or are in preparation.

LOCATION AND EXTENT OF AREA

Clay County is in north-central Kansas, in the second tier of counties south of the Nebraska line and about 50 miles east of the center of the state. The county is bounded on the west by Ottawa and Cloud Counties, on the north by Washington County, on the east by Riley and Geary Counties, and on the south by Dickinson County; it has an area of 660 square miles.

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### PREVIOUS INVESTIGATIONS

The principal studies of the geology and ground-water resources of north-central Kansas that have a direct bearing on Clay County are cited below. Specific references are cited by author and date at appropriate places in the text; all are listed in the references at the end of the report.

The geology of the Upper Cretaceous rocks in Kansas was described by Logan (1897). Darton (1905) made reference to wells in north-central Kansas in a preliminary report on the geology and ground-water resources of the central Great Plains. Haworth (1913) prepared a report on well waters in Kansas in which he discussed the availability of ground water in the Republican River valley. Jewett (1941) prepared a report on the geology of Riley and Geary Counties; in the western part of these counties the geology is very similar to that in eastern Clay County. Plummer and Romary (1947) prepared a report on clay in Kansas in which they described the ceramic characteristics of the clays of the Dakota Formation in the area. Schoewe (1952) in a report on the coal resources of the Dakota Formation described coal reserves in the county. Frye and Leonard (1952) in a report on the Pleistocene geology of Kansas described unconsolidated deposits in the Republican River valley in Clay County.

### METHODS OF INVESTIGATION

One month in the fall of 1953, four months in the summer of 1954, and one month in the fall of 1955 were spent in the field collecting the data upon which this report is based. The geology was mapped on aerial photographs from field observations and from stereoscopic study of the photographs. This information was transferred from the photographs to a base map modified from a map prepared by the Soil Conservation Service of the Department of Agriculture.

Data on the character of the water-bearing material, and on the depth, depth to water, and yield of 143 wells were collected (Table 10). Holes were augered with a power auger at 110 locations in order to gather both geologic and hydrologic information. Additional information on the material below the land surface was obtained by drilling 28 test holes with a hydraulic rotary drilling machine owned by the State Geological Survey of Kansas and operated by E. L. Reavis and William Gellinger. Logs of the test holes were prepared in the field, and the drill cuttings were studied

microscopically in the laboratory. Altitudes at the surface of the wells and test holes were determined by a level party headed by Mr. Reavis using an alidade and plane table.

Several pumping (aquifer) tests were made to determine the hydraulic characteristics of the water-bearing deposits.

Thirty-three samples (two composite) of water from 35 wells and test holes were collected, and chemical analyses of the samples were made by Howard Stoltenberg, chemist in the Water and Sewage Laboratory of the Kansas State Board of Health (Table 3).

#### WELL-NUMBERING SYSTEM

In this report the wells, auger holes, and test holes are numbered according to their location as given by the General Land Office system of land classification. The component parts of a well number are the township number, range number, section number, and two or three lowercase letters that indicate respectively the quarter section, quarter-quarter section, and, if there are three, the quarter-quarter-quarter section in which the well is located. The lowercase letters are assigned to the quarter divisions in a counterclockwise direction, beginning in the northeast quarter of each section or subdivision. For example, well 6-1-2baa (Fig. 2) is in the NE $\frac{1}{4}$  NE $\frac{1}{4}$  NW $\frac{1}{4}$  sec. 2, T. 6 S., R. 1 E. If there are two or more wells in the same quarter-quarter-quarter section they are numbered serially according to the order in which they were inventoried.

#### ACKNOWLEDGMENTS

Appreciation is expressed to the many residents of Clay County who supplied information and aided in the collection of field data. Special acknowledgment is due the officials of the cities and industries who provided information about their water supplies. Acknowledgment is made also of information supplied by George Cox, driller, and E. H. Erickerman, irrigation-equipment supplier. F. Turner and H. Rhodes were very cooperative in permitting aquifer tests using their wells. Judge Walter O. Curtis of Clay Center was helpful in supplying information about some of the older wells in the area.

The manuscript of this report has been reviewed by several members of the U. S. Geological Survey and the State Geological Survey of Kansas; Robert Smrha, Chief Engineer, and George S. Knapp, Engineer, Division of Water Resources, Kansas State Board of Agriculture; and Dwight Metzler, Chief Engineer, and Willard Hilton, Geologist, Division of Sanitation, Kansas State Board of Health.

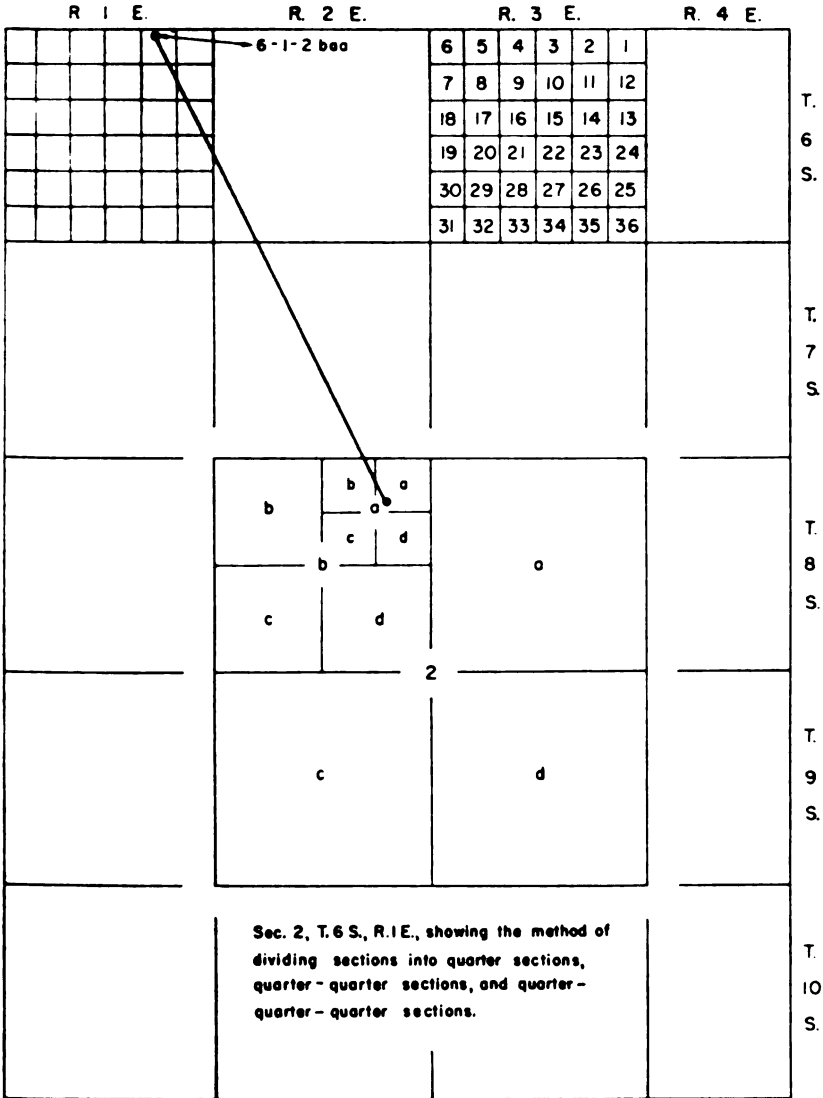


FIG. 2.—Map of Clay County illustrating well-numbering system used in this report.

## GEOGRAPHY

## TOPOGRAPHY AND DRAINAGE

Schoewe (1949) placed the western part of Clay County in the Smoky Hills division of the Dissected High Plains section of the Great Plains physiographic province and the eastern part of the county in the Flint Hills Upland division of the Osage Plains section of the Central Lowlands physiographic province. The border between the two divisions roughly coincides with the contact between the Cretaceous (Dakota Formation) and Permian (Wellington Formation) rocks through the county. The topography is varied. In the western part of the county, areas underlain with Dakota Formation typically are gently rolling surfaces dotted with numerous mounds or buttes capped by sandstone. East of the outcrop of the Dakota Formation the topography is influenced by the thick shales of the Wellington Formation, which produce a gently rolling surface dissected by streams. Local relief is relatively small. The topography in the easternmost part of the county is dominated by the bedrock of the Chase Group. Here gently rolling uplands contrast with sharp breaks into the valley areas, and the relief is considerably greater than that to the west in the outcrop area of the Wellington Formation. The broad, flat terraces in the valley of Republican River are important for agriculture and ground water.

The highest points in Clay County (about 1,500 feet above sea level) are the hills formed by sandstone of the Dakota Formation a short distance northeast of Oak Hill, and the lowest point (altitude about 1,100 feet) is in the Republican River valley in the southeastern part of the county.

Approximately two-thirds of Clay County lies in the drainage basin of Republican River. The river enters the county near the northwest corner, flows southeastward across the county, and leaves the county a short distance north of the southeast corner. An area of about 50 square miles in northeastern Clay County drains into Fancy Creek and thence into Big Blue River. In southwestern Clay County an area of about 120 square miles drains into Chapman Creek and thence into Smoky Hill River.

## CLIMATE

The U. S. Weather Bureau has maintained precipitation and temperature gages at Clay Center since 1902. The normal annual precipitation at Clay Center is 27.92 inches. The annual precipitation and the cumulative departure from normal for the period of record are shown graphically in Figure 3. The precipitation has ranged

from a low of 15.57 inches in 1934 to a high of 53.86 inches in 1951. About three-quarters of the precipitation falls during the 6-month period between April 1 and September 30, which is the period most favorable for growing crops. The normal monthly precipitation is shown graphically in Figure 4.

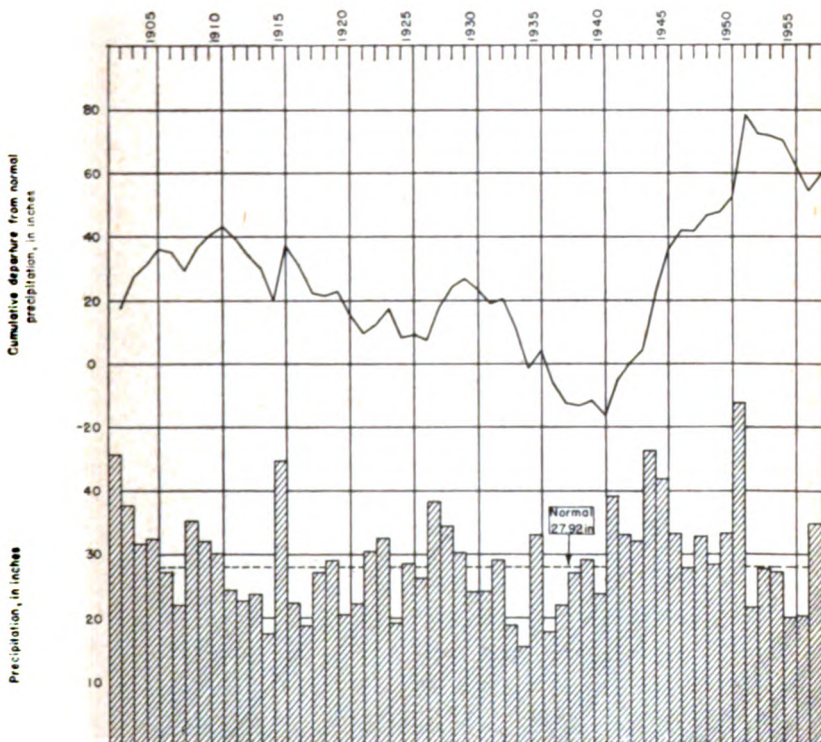


FIG. 3.—Annual precipitation and cumulative departure from normal precipitation at Clay Center.

The mean annual temperature at Clay Center is 55.1°F. The growing season in Clay County has ranged from 139 days to 213 days and averages 173 days. The earliest killing frost of record occurred on September 20, 1918; the latest, May 27, 1907.

POPULATION

According to the 1950 census, the population of Clay County was 11,697 and average density was 17.8 persons per square mile as compared with 23.1 per square mile for the entire state. The census records show that there has been a gradual decline in popu-

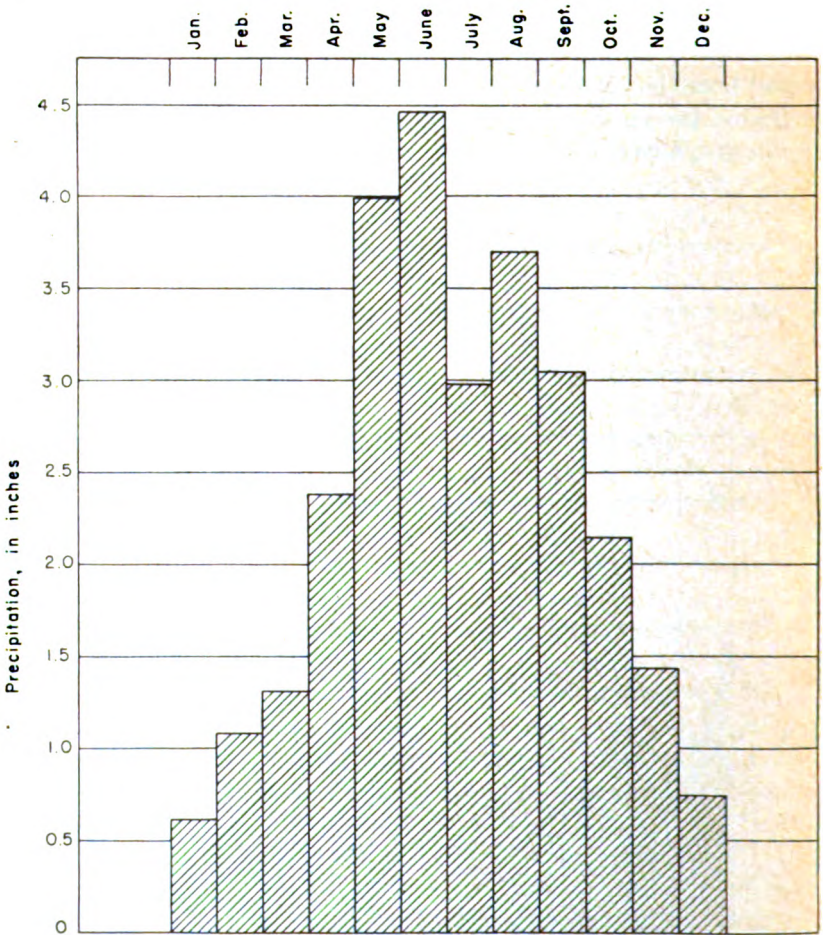


FIG. 4.—Normal monthly precipitation at Clay Center.

lation in the last 20 years, averaging about 140 persons a year. The urban population has remained about the same since the 1940 census, but the rural population has decreased. The principal cities and their population as shown by the 1950 census are as follows: Clay Center, 4,528; Clifton, 365; Green, 219; Longford, 178; Morganville, 278; Oak Hill, 92; Vining, 85; and Wakefield, 591.

#### TRANSPORTATION

Clay County has good transportation facilities. All but one of the cities are served by one or more railroads. Broughton, Clay Center, Morganville, Clifton, and Vining are served by a branch



line of the Chicago, Rock Island and Pacific Railroad Company. Longford and Oak Hill are served by a branch line of the Atchison, Topeka and Santa Fe Railroad Company. Wakefield, Broughton, Clay Center, Idana, and Browndale are served by a branch line of the Union Pacific Railroad. The city of Green was formerly served by another branch line of the Union Pacific, but this line has been abandoned.

Clay County is served by a good state and national highway system. U. S. Highway 24 crosses the county from east to west through Clay Center. Kansas Highway 15 crosses the county from north to south, and Kansas Highway 9 crosses the northwestern part of the county from Kansas Highway 15 west through Clifton. Kansas Highway 82 crosses the southeastern part of the county from Kansas Highway 15 east through Wakefield. All state and federal highways are hard surfaced, as are many of the county highways. Most of the township roads are graded and many of them are all-weather roads.

#### AGRICULTURE

Agriculture is the principal occupation in Clay County. About 85 percent of the land area in the county is used for agriculture, about one-third pasture and two-thirds cultivated. Wheat is the principal crop and is planted on about half the cultivated land.

The principal crops and the acres harvested in 1955 are as follows: Wheat, 94,000 acres; corn, 22,000 acres; alfalfa, 38,500 acres; sorghums, 21,700 acres; oats, 17,200 acres; and wild hay, 8,500 acres.

Ranking equally in value with grain and hay in the county is livestock products. In order of value these animal products are cattle, poultry and eggs, dairy products, and hogs. The value of livestock, poultry, milk, and eggs in 1955 was about \$5,267,000. that of all grains and hay harvested was about \$5,225,000 in 1955.

#### MINERAL RESOURCES

##### Construction Materials

Geologic materials that would be useful in construction projects include the alluvial deposits in the major stream valleys, sandstones and clays of the Dakota Formation, and Permian limestones in the eastern part of the county. For discussion these materials are classified according to use.

*Concrete aggregate.*—Aggregate for concrete consists of fragments of hard, durable minerals or rocks of sand and gravel size.

The constituent particles should be free from adherent coatings or particles that would interfere with the bonding of the cement and the aggregate. An almost unlimited quantity of sand and gravel is available from the alluvial deposits in the Republican River valley. Crushed limestone aggregate can be produced in large quantities from the Permian limestones in the eastern part of the county. Material for the manufacture of lightweight aggregate is available from the Dakota Formation and from thin deposits of the Kiowa Shale.

*Road metal.*—Road-surfacing material is available in large quantities from the sand and gravel in the alluvial deposits of the Republican River valley. Sandstone and iron-cemented sandstone of the Dakota Formation can be crushed for road metal and are used extensively in the western part of the county. Crushed limestone from the Permian beds is used extensively in the eastern part of the county.

*Mineral filler.*—Silts from loess deposits of the Sanborn Group and from terrace deposits in the Republican River valley are available and are used in large quantities for mineral filler in road-surfacing material in the county.

*Structural stone.*—Structural stone is any hard dense rock that can be quarried and cut to the desired size and shape. Materials that meet these requirements are found in the Cresswell Limestone member of the Winfield Limestone and the Fort Riley Limestone member of the Barneston Limestone. These limestones are found in the eastern part of the county, and many buildings in the area have been constructed from materials taken from small local quarries.

#### Agricultural Limestone

Limestone having a calcium carbonate equivalent of 80 percent or more and occurring in ledges sufficiently thick to allow economical quarrying is regarded as a potential source for agricultural limestone. In Clay County most of the limestones, especially the Cresswell Limestone member of the Winfield Limestone and the Fort Riley Limestone member of the Barneston Limestone, meet this requirement. Some zones in many of the shales contain enough calcium carbonate to be used for agricultural limestone.

#### Oil

Oil was discovered in the Wakefield pool in Clay County in 1928. The producing zone was the "Chat" at the top of the Mississippian rocks at a depth of 1,774 feet. After about 2 years the well was abandoned, but in 1951, production was resumed in this pool, and

the Wakefield Northeast pool was discovered. In 1952, one well was added to each pool. No production figures are available for Clay County. Of 25 oil tests drilled in the county, 5 have produced some oil.

### Gypsum

Gypsum is not produced commercially in Clay County, although there is a deposit in the NW¼ SW¼ sec. 19, T. 6 S., R. 4 E. Its areal extent is not known, but the geology and topography are such that the gypsum may extend over a considerable area. Unusually high sulfate content of water samples in the area probably indicates the presence of gypsum, though possibly not in commercial quantity. Test hole 6-4-19cb penetrated 13 feet of gypsum, including one seam of clean gypsum 8 feet thick. This deposit rests on the Herington Limestone member of the Nolans Limestone.

### Ceramic Raw Materials

By NORMAN PLUMMER

Although there is no ceramic industry in Clay County at present, there is available an abundance of clay and shale suitable for use in the manufacture of structural brick and tile, fire brick, and pottery. The most important ceramic clay deposits are in the Dakota Formation, which is exposed in the western and northern parts of the county. Many of the clay deposits in the Dakota Formation fire to light colors ranging from ivory to buff. Most such clays are refractory and could be used for making Low Duty and Intermediate Duty fire brick as well as face brick and pottery. Other clays of the Dakota Formation fire to darker colors ranging from dark buff to deep red (Plummer and Romary, 1947).

Most of the Pleistocene loess deposits of the county are suitable for use in the manufacture of common brick, or as a flux for mixing with the more refractory clays. Such blends are especially suitable for the manufacture of sewer pipe (Frye, Plummer, Runnels, and Hladik, 1949).

Permian shale was at one time used in the manufacture of brick at a plant a short distance southeast of Clay Center. Although acceptable brick can be made from this material, it would be considered inferior in the light of modern standards of brickmaking, and because high-quality raw materials are available it probably would not be used.

The Kiowa Shale of Cretaceous Age, which underlies the Dakota Formation, is an excellent "bloating" shale from which a good grade of lightweight concrete aggregate could be produced.

TABLE 1.—Generalized section of outcropping geologic rock units in Clay County

System	Series	Stage	Group	Formation	Member	Thick-ness (feet)	Physical character	Water supply	
Quaternary	Pleistocene	Recent		Dune sand		0-35	Fine to medium quartz sand, and some silt.	Above the water table, and does not yield water to wells in Clay County.	
				Alluvium		0-70	Stream-deposited clay, silt, sand and gravel. Generally graded, the coarser materials at the base.	Yields large quantities of water to wells in Clay County.	
			Wisconsinan	Sanborn	Peoria		0-20	Wisconsinan terrace deposits and the Crete formation consist of stream-deposited beds of silt, clay, sand, and gravel. Love- and Peoria formations consist of massive eolian deposits	Wisconsinan terrace deposits yield large quantities of water in Republican River valley and smaller amounts along tributary streams. Silt units yield little or no water to wells; Crete formation yields small to moderate quantities of water to wells in Clay County.
					Wisconsinan terrace deposits		0-70		
					Loveland		0-25		
					Crete		0-25		
		Kansas	Meade	Sappa		0-5	Gray clay and silt.	Does not yield water to wells in Clay County.	
				Grand Island		0-10	Locally derived gravel.	Yields moderate amounts of water to wells in Clay County.	
Cretaceous	Gulfian			Dakota Formation	Jansen Clay Terra Cotta Clay	0-165	Varicolored silty and sandy shale and massive sandstone beds.	Yields moderate amounts of water to wells in Clay County.	
				Kiowa Shale		0-50	Dark fissile shale containing a few thin streaks of sand and silt.	Does not yield water to wells in Clay County.	



## GEOLOGY

## SUMMARY OF STRATIGRAPHY \*

The areal geology of Clay County is shown on Plate 1. The rocks cropping out in the county are of sedimentary origin. The oldest rock exposed at the surface is the Barneston Limestone of the Chase Group, Wolfcampian Series, Permian System. It is exposed in the valley wall of Republican River in the vicinity of Wakefield. Successively younger Permian rocks are exposed west of the outcrop of the Barneston Limestone. The Wellington Formation, the youngest Permian formation in Clay County, is exposed in a wide belt extending across the county from north to south.

The Kiowa Shale is known to overlie the Wellington Formation in southwestern Clay County, but it is so poorly exposed that it is mapped with the Dakota. The Dakota Formation, the youngest Cretaceous rock unit in the county, crops out over a large part of western and northern Clay County.

Much of the upland area is blanketed by eolian silt or loess of Pleistocene age. Extensive Pleistocene alluvial deposits occur in the valleys.

A generalized section of the outcropping rock units of Clay County is given in Table 1. More detailed descriptions of these units are given in the section on rock units and their water-bearing characteristics.

## PRE-CENOZOIC GEOLOGIC HISTORY

Although exposed rocks in Clay County total only a few hundred feet in thickness, a great deal is known about the much thicker sequence of deeply buried rocks through the interpretation of deep tests for oil and gas. The geologic history during the Paleozoic Era, as it is discussed here, is based chiefly on reports by Jewett (1949) and by Lee, Leatherrock, and Botinelly (1948).

## Paleozoic Era

Clay County, like the rest of Kansas, is underlain by a basement complex of Precambrian crystalline rocks. The area that is now Clay County was invaded by the sea during Cambrian time, and remained an area of marine deposition during most of the Paleozoic Era. Rocks representing each of the systems of the Paleozoic Era are present, but leveling by erosion of certain beds indicates that the area was not always stable and was elevated above sea level

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\* The stratigraphic classification used in this report is that of the State Geological Survey of Kansas and does not necessarily follow the nomenclature of the U. S. Geological Survey.

from time to time. Clay County is in the central part of the early Paleozoic North Kansas Basin and on the eastern flank of the late Paleozoic Salina Basin. Sedimentary rocks present under Clay County total about 3,000 feet in thickness.

### Mesozoic Era

The Paleozoic Era was brought to a close by uplifting of the entire area by the end of Permian time. The area was above sea level and subjected to erosion throughout all of Triassic and Jurassic time. Erosion continued through Early Cretaceous time until the deposition of the Kiowa Shale during a brief invasion by the sea. As the sea retreated from the area, the Dakota Formation was deposited in shallow fresh water. The sea again invaded the area and marine deposition was resumed.

At the end of the Cretaceous Period an area far to the west was uplifted, forming the Rocky Mountains. The Mississippi Valley area also was uplifted at about the same time, resulting in a north-west regional dip of Cretaceous and older rocks in eastern and central Kansas.

## CENOZOIC GEOLOGIC HISTORY

### Tertiary Period

After the withdrawal of the Cretaceous sea and the tilting of Cretaceous and older rocks, there was a long period of erosion. The Cretaceous rocks overlying the Dakota Formation were eroded from the area that is now Clay County. Late in the Tertiary Period (Pliocene Epoch) the area to the west of Clay County was receiving deposits of sand and gravel from the eroding Rocky Mountains. These sand and gravel deposits, the Ogallala Formation, may not have reached as far east as Clay County, but they are found in several scattered areas a few miles west of Clay County.

### Quaternary Period—Pleistocene Epoch

The major events of the Pleistocene Epoch in this area were the establishment of new drainage lines and downcutting by existing streams, aggradation of the major streams, and the deposition of loess. The relation of the Pleistocene deposits is shown in Plate 3.

At the beginning of the Pleistocene Epoch, Clay County probably was drained by a system of small streams, none of whose headwaters were very far from the county. Climatic changes associated with the advance of the glacial ice fronts during the Pleistocene Epoch resulted in increased streamflow and downcutting by streams. There was little or no deposition along streams in Clay County during the

first glacial stage, the Nebraskan. The second glacial stage, the Kansan, continued much as the Nebraskan stage had been. In early Illinoian time, the ancestral Republican River became blocked by ice or choked by sediments. The ancestral Republican River followed approximately the course of the present Republican River from its headwaters to the city of Republic in Republic County. From this point the ancestral river flowed northeastward into Nebraska (Lohman, in Fishel, 1948). After the damming of the ancestral Republican River in early Illinoian time, the impounded waters spilled over the lowest part of a divide between Republic and Scandia in Republic County, and the present course of Republican River was established from Republic to Junction City. Republican River and several smaller streams in Clay County have alternated since late Kansan time between periods of deposition and periods of downcutting, so that the valleys now contain terrace and alluvial deposits of Kansan, Illinoian, Wisconsinan, and Recent age.

During Recent time, sand has been carried from the flood plain of Republican River by the prevailing westerly winds and has been deposited in several places as sand dunes on the alluvium and Wisconsinan terraces.

## GROUND WATER

### PRINCIPLES OF OCCURRENCE

The rocks and surficial deposits that form the crust of the earth are, in general, not solid throughout, but contain many open spaces, called voids or interstices, and it is in these spaces that water is found below the surface of the earth and from which it is recovered, in part, through springs and wells. There are many types of rocks and they differ greatly in the number, size, and arrangement of their interstices and therefore in their water-bearing properties. The occurrence of ground water in any region, therefore, is determined by the geology of that region.

The interstices of rocks range in size from pores of microscopic dimensions to openings several inches or feet in width. These interstices can be divided into two classes, primary interstices and secondary interstices. The primary or original interstices were formed during the formation of the rocks. Secondary interstices were developed by the different processes that affected the rocks after deposition. In Clay County all the water-bearing rocks are sedimentary, and the openings that hold the water are either open pore spaces between the grains of rock (primary interstices) or joints and open bedding planes, which have resulted from deforma-



tion of the rocks, and openings caused by solution of the rocks. These are secondary interstices.

The amount of water that can be stored in any rock depends upon the porosity of that rock. Porosity is expressed quantitatively as the percentage of the total volume of the rock that is occupied by interstices. When all the interstices in a rock are filled with water the rock is said to be saturated. The amount of water that a saturated rock will yield to the force of gravity is known as the specific yield. The amount of water that a rock can hold is determined by its porosity, but the rate at which it will yield water to wells is determined by its permeability. The permeability of a rock is its ability to transmit water under hydraulic gradient and is measured by the rate at which the rock will transmit water through a given cross section under a given loss of head per unit of distance. Some beds of clay or shale may have a high porosity, but because the interstices are small and poorly connected, they transmit little or no water, and the rock may be regarded as virtually impervious. Rocks differ greatly in their degree of permeability, according to the number, size, and interconnection of their interstices.

#### SOURCE

Ground water is the part of the water below the surface of the earth that supplies wells and springs. In Clay County, ground water is derived entirely from precipitation, in the form of rain or snow, that falls directly on the county or on nearby areas. Part of the precipitation that falls as rain or snow is carried away by surface runoff and is discharged by streams, a part of it may evaporate, and a part may be absorbed by vegetation and transpired into the atmosphere. The part that escapes runoff, evaporation, and transpiration percolates slowly downward through the soil and underlying strata until it reaches the water table, where it joins the body of ground water in the zone of saturation. After reaching the ground-water body the water percolates slowly through the rocks in directions determined by the geology, topography, and geologic structure until it is discharged through wells and springs or by evaporation and transpiration in areas where the water table is shallow.

#### ARTESIAN CONDITIONS

Artesian conditions may exist where a water-bearing bed is overlain by a relatively impermeable bed that dips from its outcrop area toward the discharge area. Water entering the water-bearing bed in the outcrop area percolates downward to the water table and moves downdip beneath the confining bed. The weight of the

water at higher levels in the confined bed creates a hydrostatic pressure in the water reservoir. When the confining bed is penetrated the water will rise in the drill hole to a level equal to the hydrostatic head of the aquifer at the point of discharge. If the pressure in the aquifer is sufficient to lift the column of water above the top of the aquifer, artesian conditions exist. If the pressure in the aquifer is sufficient to lift the water above the land surface, the well will flow naturally. Such a well is called a flowing artesian well. Well 8-4-21ac flowed when drilled, but as the flow continued only a few days, a pump was installed. A flowing well was observed a few hundred yards west of the Clay County line, in the NE $\frac{1}{4}$  sec. 13, T. 8 S., R. 1 W. Geologic and topographic conditions indicate that similar flowing wells could be drilled in an adjacent area in Clay County.

Although only one flowing well is known to have existed in the county, it is probable that artesian conditions exist locally over much of the county. In western Clay County the Dakota Formation contains lenticular sandstone bodies semiconfined beneath clay and shale, and much of the water is probably under some hydrostatic pressure. In eastern Clay County the rocks are a sequence of limestones and relatively impermeable shales. The limestones are in general the best aquifers, and local structures create hydrostatic head within the aquifers; although wells may not flow at the surface, in many of them the water rises above the aquifer.

#### THE WATER TABLE AND MOVEMENT OF GROUND WATER

The configuration and gradient of the water table and the direction of ground-water movement generally can be shown by constructing a water-table contour map. In Clay County, however, the water table is not everywhere continuous, so the water-table contours shown on Plate 2 cover only part of the areas discussed below. Clay County can be divided into three general areas on the basis of the occurrence and movement of ground water. The largest of these areas is that underlain by rocks of Permian age; it includes roughly the area east of a north-south line passing through the center of Range 2 East (Pl. 1). Although the largest, this area is probably the least important as a source of ground water, because it is underlain chiefly by shale. Water is derived principally from limestones but occurs also in appreciable quantity in solution channels in gypsum deposits in the lower shales of the Wellington Formation. The regional dip in this area is westward, and the water moves downdip from the recharge area that lies to the east.

The second largest ground-water area is that underlain by the Dakota Formation (Pl. 1). In this area water is obtained from sandstones in the lower part of the Dakota Formation. Water-table contours were drawn over a part of this area (Pl. 2), but in the areas where the Dakota Formation is thin it was not feasible to construct contours. Also, in many areas the water levels in the shallow wells did not fit the contour pattern of the water levels in deeper wells. The shallow water seemed to occupy a perched or semiperched position in relation to the rest of the Dakota Formation, and for this reason the shallow water levels were not used in constructing the water-table contours. The movement of ground water in this area is generally eastward toward the valley of Republican River.

The third and most important ground-water area in the county is the Republican River valley. Water is obtained from unconsolidated alluvial deposits, and large supplies are available for irrigation. The Republican River valley can be divided into two areas. One area is that part of the valley above Clay Center, where the Nolans Limestone crops out in the valley walls, and the second area is that part of the valley below the outcrop of the Nolans Limestone. In the upper area the valley is broad and flat and the contours cross the valley nearly at right angles to the river and valley walls, indicating that the river is neither gaining water from the ground-water reservoir nor losing water to it. The portion of the valley below the Nolans Limestone outcrop is narrower, and the water-table contours intersect the river at sharp angles, indicating that the river is draining the ground-water reservoir in this area.

#### RECHARGE

The addition of water to the zone of saturation is known as ground-water recharge. Ground-water recharge in Clay County occurs by infiltration from precipitation within the county, by percolation from influent streams, and by subsurface inflow from adjacent areas.

#### Recharge from Precipitation

The areas in Clay County that are most favorable to ground-water recharge are the small dune tracts near Clifton and Wakefield, where poor drainage and permeable soils tend to reduce runoff and induce infiltration. Next in importance is the valley area, which is relatively flat and is underlain by very permeable materials. Much water enters the ground-water reservoir in the outcrop area

of the Dakota Formation where permeable sandstones are at or near the surface, although the steep slopes of this area tend to encourage runoff.

#### Percolation from Outside the Area

Most of the recharge in the area underlain by the Permian rocks necessarily penetrates through the more porous limestone beds in their outcrop areas and moves downdip to places where it is discharged. East of Clay County the lower part of the Barneston Limestone is cherty and fractured, and this area is probably the most important recharge area contributing water to the Permian rocks in Clay County.

Much water moves across the west county line of Clay County, through the Dakota Formation (Pl. 2). This water moves in a generally eastward direction to be discharged by seeps and springs and directly into the Republican River valley.

#### Seepage from Streams and Ponds

Much water is stored in ponds in Clay County. A large part of this water is lost through evaporation and transpiration and a part is used by livestock. The remainder seeps into the ground and becomes ground water. Under normal conditions the streams in Clay County contribute but little water to the ground-water reservoir, but when the stream stage is high a part of the water enters the stream bank as "bank storage". This process may become an important source of recharge in the future if large-scale development of irrigation in the valley area lowers the water table in these deposits and permits infiltration from the streams.

#### DISCHARGE

Ground water is discharged in Clay County by evaporation and transpiration, by seeps and springs, by wells, and by percolation to areas outside the county.

#### Discharge by Evaporation and Transpiration

In the areas in Clay County where the water table is only a few feet below the surface, water may be evaporated from storage. This is especially true in the valley of Republican River and to a lesser extent along the steeper slopes of the upland area where the water table is relatively near the surface and where seeps and springs occur. Transpiration accounts for much discharge in areas where the water table can be reached by the roots of plants.

### Discharge by Seeps and Springs

The water-table contours (Pl. 2) indicate that ground water in the outcrop area of the Dakota Formation moves toward Republican River or its tributaries. Many springs and seeps occur along the valley walls and steep slopes adjacent to tributary streams, and a considerable quantity of water is discharged from storage by this means. Springs occur along the outcrop of the Permian limestones in eastern Clay County where structural conditions are favorable. This water has moved down dip from the recharge area farther east and is discharged in considerable quantity.

### Discharge by Wells

Although large quantities of water are discharged through wells in Clay County, it is probable that the discharge of water by wells is small in comparison with discharge by other means. With increased development of irrigation, discharge by wells will become a more important factor.

## RECOVERY

### Principles of Recovery

When water is standing in a well the head of the water in the aquifer outside the well is in equilibrium with that in the well. When water is withdrawn from a well a difference in head is created between the water inside the well and the water outside the well for some distance from the well. The water table in the vicinity of the well develops a cone of influence (Fig. 5), which is deepest at the wall of the well and extends some distance around the well. An increase in the pumping rate of the well produces a greater drawdown.

The specific capacity of a well is the rate of yield per unit of drawdown and is generally given in gallons a minute per foot of drawdown. Specific capacity is generally determined after the well has been pumped for a period long enough to stabilize the drawdown.

The character of the water-bearing material and the type of construction control the yield, drawdown, and specific capacity of a well. If the water-bearing material is coarse and well sorted it will readily yield large quantities of water to wells and have a minimum drawdown. If on the other hand the material is fine and poorly sorted, it will offer much resistance to the movement of water and

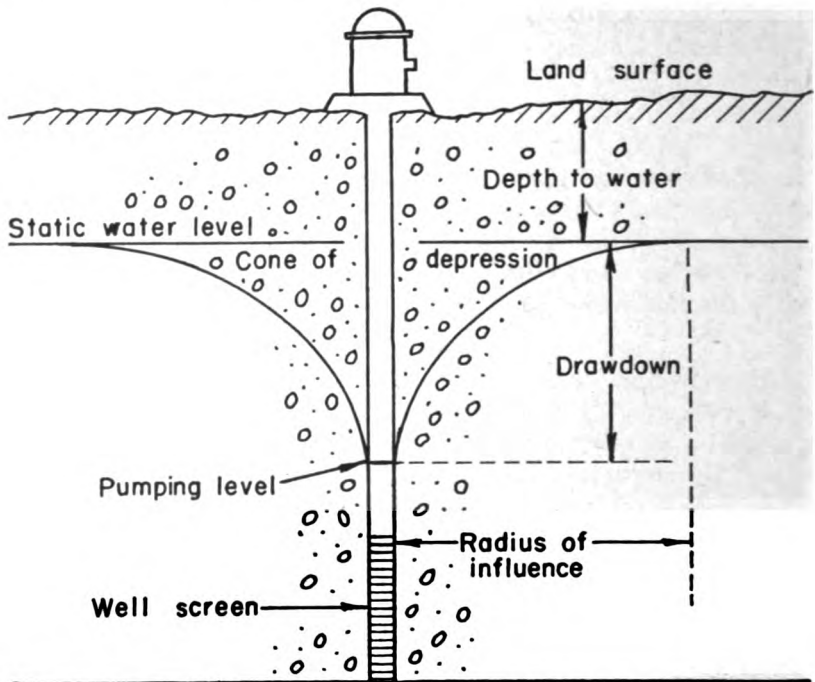


FIG. 5.—Diagrammatic section of well that is being pumped, showing drawdown, cone of depression, and radius of influence.

thereby yield less water and have a greater drawdown. All other things being equal, the drawdown of a well varies inversely with the permeability of the water-bearing material.

#### Types of Wells

Several different types of wells are used in Clay County for the recovery of ground water. Wells generally are classed according to the method of construction. The particular type of well used at any location depends on the geology of the area, the depth to water and thickness of the saturated material at the location, and the use for which the well is constructed.

*Dug wells.*—Dug wells are of large diameter, generally constructed with pick and shovel or other hand tools, and walled with stone, tile, or concrete. Such wells generally are constructed in areas where large supplies of ground water are not available, and the large hole is used as a storage reservoir, which provides water during periods of pumping and slowly fills when the pump is not running.

*Driven wells.*—Driven wells are constructed by driving a pipe, generally 1½ to 2 inches in diameter and equipped with a screen, to a point below the water table. Such wells are pumped with some type of shallow-well pump. Most driven wells are in valley areas, for they are limited to areas where the water table is not much below 20 feet and where the deposits are unconsolidated. This type of well also is generally used when only small supplies of water are required.

*Drilled wells.*—Drilled wells are constructed by use of power-driven equipment to operate a percussion or rotary drill rig. This is the most common type of well in Clay County and is found in all types of material and in all localities. The depth of drilled wells is determined by the depth to the water table and the use for which the well is constructed. The use in turn may be limited by the quantity of water available. The diameter of drilled wells is generally dependent upon the use and desired quantity of water. Industrial, irrigation, and municipal wells require larger quantities of water and larger pumps and therefore require large-diameter holes to accommodate these pumps.

#### UTILIZATION OF WATER

Ground water in Clay County is used chiefly for domestic, stock, and public supplies. Recent interest in irrigation in the valley areas in the county has caused an increase in the use of ground water for this purpose, and it is probable that there will be a considerable future increase. The quantity of water used for industrial purposes is small compared to that for other uses.

#### Domestic and Stock Supplies

Nearly all the domestic and stock supplies of water in rural areas and in small communities that do not have a public supply are obtained from privately owned wells. In the valley areas these supplies are obtained chiefly from drilled and driven wells. In the upland areas the water supplies are obtained from either dug or drilled wells. In several small areas underlain by shales of the Wellington Formation, there is a scarcity of wells and there are many abandoned farmsteads. A lack of dependable water supplies in these areas was a contributing factor in the abandoning of the farmsteads. In other areas the Wellington Formation yields water of good quality to wells.

#### Industrial Supplies

The quantity of water used for industrial purposes in Clay County is small in comparison with that for other uses. The Northern Natural Gas Company has four wells that yield water from terrace

deposits in the Republican River valley in the NW¼ sec. 1, T. 6 S., R. 1 E. The wells are 50 feet deep, are equipped with turbine pumps, and pump an average of 325,000 gallons of ground water a day for cooling. In Clay Center, Swift and Company uses water from a well (8-3-8bb1) for cooling. This well is about 65 feet deep, is equipped with a turbine pump having a capacity of 120 gpm, and is pumped at an average rate of about 170,000 gpd. Well 8-3-8bc is owned by the Clay Center Coca Cola Bottling Company. This well is equipped with a jet pump having a capacity of about 50 gpm. The water is softened and the iron is removed. The average daily use is about 1,500 gallons.

### Public Supplies

Seven cities in Clay County have public water supplies. These supplies are discussed in the following paragraphs.

*Clay Center.*—The Clay Center water supply is obtained from five wells drilled in the low terrace deposits in the Republican River valley. Three of these wells in Utility Park are 55 feet deep and the depth to water is about 22 feet. Well 8-3-8bb4 was tested for 8 hours at a rate of 850 gpm, and had a drawdown of 14 feet. Well 8-3-8bd in Dexter Park, which was drilled in the summer of 1954, yields water from terrace gravels. The well is 60 feet deep and the depth to water is about 29 feet. After the well was pumped 8 hours at a rate of 500 gpm, the drawdown was 4 feet. Well 8-3-8db, in the southeastern part of Clay Center, is a drilled well 58 feet deep and yields water from terrace gravels. In August 1954 the depth to water was 28 feet. All the wells are equipped with electric turbine pumps. The maximum capacity of the five wells is about 4 million gallons per day, and the average daily use is about 800,000 gallons. The water is chlorinated but receives no other treatment. The water is stored in an elevated steel tank having a capacity of 500,000 gallons. In addition to the five wells used for the public water supply, the city has four wells at the municipal power plant that provide water for cooling. These are drilled wells about 50 feet deep equipped with turbine pumps. The average daily pumpage from these wells is 1.7 million gallons.

*Clifton.*—The Clifton water supply is obtained from two wells (6-1-2bac1 and 6-1-2bac2) in the southwestern part of the city. The wells, which are equipped with turbine pumps, yield water from terrace deposits in the Republican River valley and are about 65 feet deep. The water level in August 1954 was about 30 feet. The drawdown in each of these wells is 7 feet at a pumping rate of 160 gpm. The maximum capacity of the two wells is about 360,000



gallons per day and the average daily use is 100,000 gallons. Hardness is reduced by pressure zeolite softeners and the water is chlorinated. A steel standpipe has a capacity of 90,000 gallons.

*Green.*—The Green municipal water supply is obtained from two wells (7-4-20ad and 7-4-21bc), which yield water from the Fort Riley Limestone member of the Barneston Limestone at depths of 170 and 190 feet respectively. The water is hard but receives no treatment. The capacity of the wells is about 30,000 gpd, and the average daily use is about 10,000 gallons. Water is pumped directly to an elevated concrete storage tank having a capacity of 60,000 gallons.

*Idana.*—The Idana water supply is obtained from one well (8-1-13ca) south of town in a creek valley. It yields water from terrace deposits along the edge of the valley. Average use at Idana is about 5,300 gpd. Storage capacity is 10,000 gallons in an underground steel tank.

*Longford.*—The Longford municipal water supply is obtained from two wells about half a mile west of the city. These wells (10-1-17dc1 and 10-1-17dc2) yield water from the Dakota Formation. The wells are 100 and 110 feet deep, respectively. The average daily pumpage is about 15,000 gallons. The water is chlorinated and delivered to the mains from an elevated concrete tank having a capacity of 60,000 gallons.

*Morganville.*—The Morganville water supply is obtained from two wells within the city limits. These wells (7-2-3cc and 7-2-4ddc) obtain water from terrace deposits in the Republican River valley. The wells are 53 and 57 feet deep, respectively, and are equipped with turbine pumps. Maximum daily yield is 160,000 gallons, and average daily pumpage is 30,000 gallons. The water is chlorinated and delivered to the mains from an elevated steel tank having a capacity of 50,000 gallons.

*Wakefield.*—The Wakefield water supply is obtained from two wells at the south edge of the city. These wells (10-4-5db1 and 10-4-5db2) obtain water from terrace deposits in the Republican River valley. The wells are 57 feet deep and are equipped with turbine pumps. The water levels in the wells were 34 feet below the surface in August 1954. After the wells are pumped for 24 hours at the rate of 100 gpm each, the drawdown is 2 feet. The yield of the wells as equipped is about 300,000 gpd, and average daily use is 80,000 gallons. The water is chlorinated and delivered to the mains from an elevated steel tank having a capacity of 30,000 gallons.

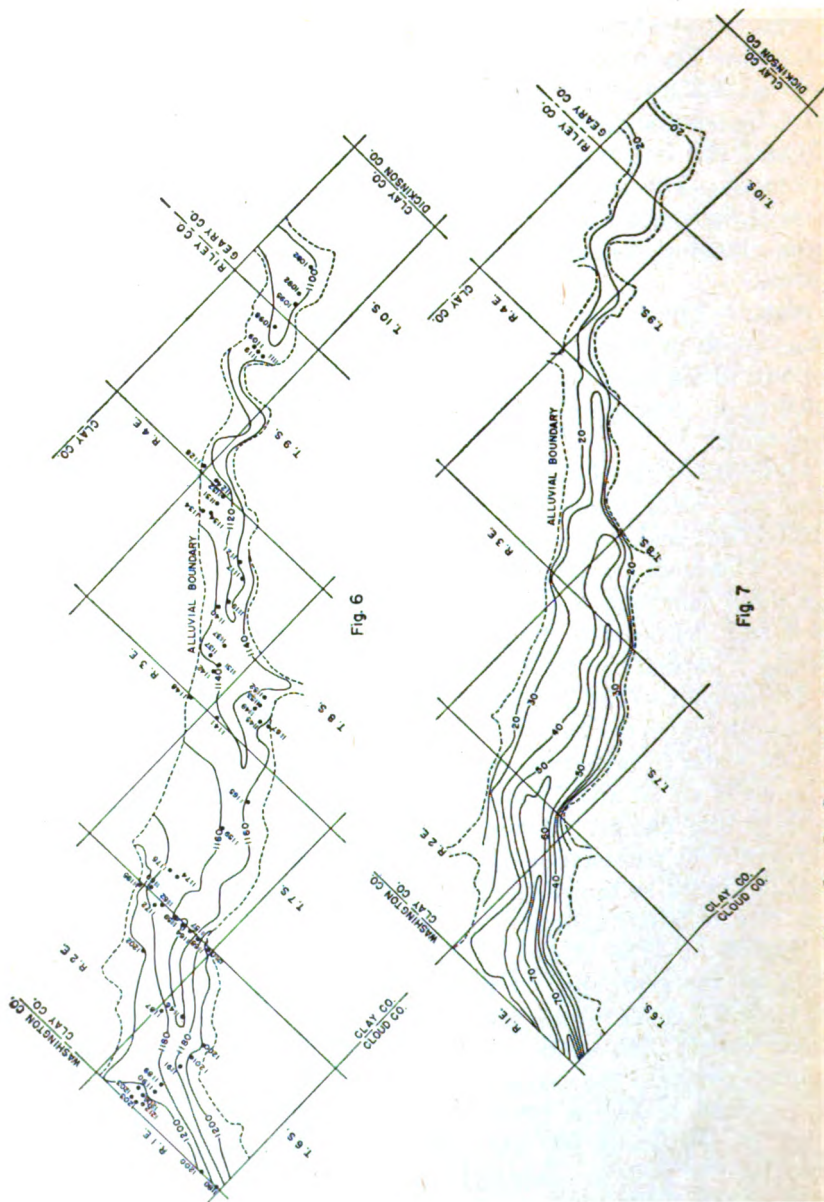


FIG. 6.—Configuration of bedrock surface in Republican River valley.  
FIG. 7.—Thickness of saturated material in Republican River valley.

## Irrigation Supplies

In the fall of 1954 six irrigation wells having a combined capacity of about 5,000 gpm were in operation in Clay County. Continued drought conditions during the growing season in 1955 intensified interest in irrigation, and by autumn 29 additional wells had been drilled, which had a combined capacity of about 21,000 gpm. In 1955 about 2,730 acre-feet of water was pumped for irrigation.

Four of the irrigation wells drilled in 1955 obtain water from the Barneston Limestone. The yields of these wells range from about 50 to 250 gpm. The other 25 wells obtain water from alluvial materials in the valley area. The yields of these wells range from 350 to 2,000 gpm.

Maps were prepared showing the configuration of the bedrock surface in the valley area (Fig. 6) and the water-table contours (Pl. 2). The water-table map was superposed over the bedrock map and a saturated-thickness map was prepared by connecting the points of equal thickness (Fig. 7). The bedrock contour map (Fig. 6) indicates that the pre-Pleistocene surface of the valley above Clay Center was a broad V-shaped valley in which the deeper part generally is near the center of the valley. Below Clay Center, where the limestones of Permian age cross the valley, the valley is also V shaped but is much narrower. The saturated-thickness map (Fig. 7) indicates that the thickest deposits of saturated material follow the deepest channels of the pre-Pleistocene surface shown in Figure 6, and the greatest thickness is in the northwest corner of the county, where the saturated materials are 80 feet thick. The thickness of these materials gradually decreases downstream to a point just west of Clay Center where the saturated material is only 40 feet thick, and below a point a short distance downstream from Clay Center the thickness of saturated materials is about 20 feet except in a very narrow channel cut in the pre-Pleistocene surface. The volume of saturated material in the valley area was determined from Figure 7, and by applying a specific yield of 20 percent the volume of water in storage was calculated. The volume of saturated material and volume of water in storage are given by townships in Table 2. Table 2 indicates that about 275,000 acre-feet of water is in storage west of the west line of Range 3 East and only about 45,000 acre-feet east of this line. The large amount of storage upstream from Clay Center in comparison with the small amount downstream is due to the great amount of alluvium upstream; the valley becomes narrower and shallower downstream because it is

TABLE 2.—Volume of saturated water-bearing material in Republican River valley and volume of water in storage based on specific yield of 20 percent.

Township	Acre-feet of saturated material	Acre-feet of water in storage
T. 6 S., R. 1 E.....	560,000	110,000
T. 6 S., R. 2 E.....	240,000	49,000
T. 7 S., R. 2 E.....	510,000	100,000
T. 7 S., R. 3 E.....	7,000	1,400
T. 8 S., R. 2 E.....	100,000	20,000
T. 8 S., R. 3 E.....	140,100	29,000
T. 9 S., R. 3 E.....	13,000	2,600
T. 9 S., R. 4 E.....	39,000	7,800
T. 10 S., R. 4 E.....	32,000	6,400
Total.....	1,600,000	320,000

cut in the more resistant Permian limestones that cross the valley downstream, whereas upstream the less resistant Wellington and Dakota Formations form the bedrock underlying the valley area. The amount of water in storage (320,000 acre-feet) is equal to the quantity of water that would be pumped in 60 to 65 years at the 1955 rate of withdrawal for industrial and municipal use, which is about 5,000 acre-feet per year. This does not take into consideration any recharge to the valley during that period. The water-table contours indicate that water moves into the valley area, and although no quantitative data are available a considerable quantity of water must be contributed to the valley area in this way. Surface water is contributed to the valley through local drainage and through Republican River. Recent completion of reservoirs on Republican River in Colorado and Nebraska will aid in maintaining a minimum flow in the river. There seems to be no immediate danger of a serious depletion of ground water in storage in the Republican Valley, but a continued increase in the irrigation development will cause a lowering of the water table in the area and may diminish the streamflow during periods of heavy pumping.

#### CHEMICAL CHARACTER OF WATER

The chemical character of the ground water in Clay County is indicated by the analyses of 33 samples (2 composite) of water from 35 wells and test holes distributed as uniformly as practicable within the area and among the principal water-bearing formations (Table 3). A graphic representation of the chemical analyses of water from representative wells in the county is shown in Figures 8 and 9.

The depth given in Table 3 for the samples of water from test holes is the depth from which the sample was pumped. For all other samples of water the depth indicated is the total depth of the well. The samples of water were analyzed by Howard A. Stoltenberg, chemist, in the Water and Sewage Laboratory of the Kansas State Board of Health.

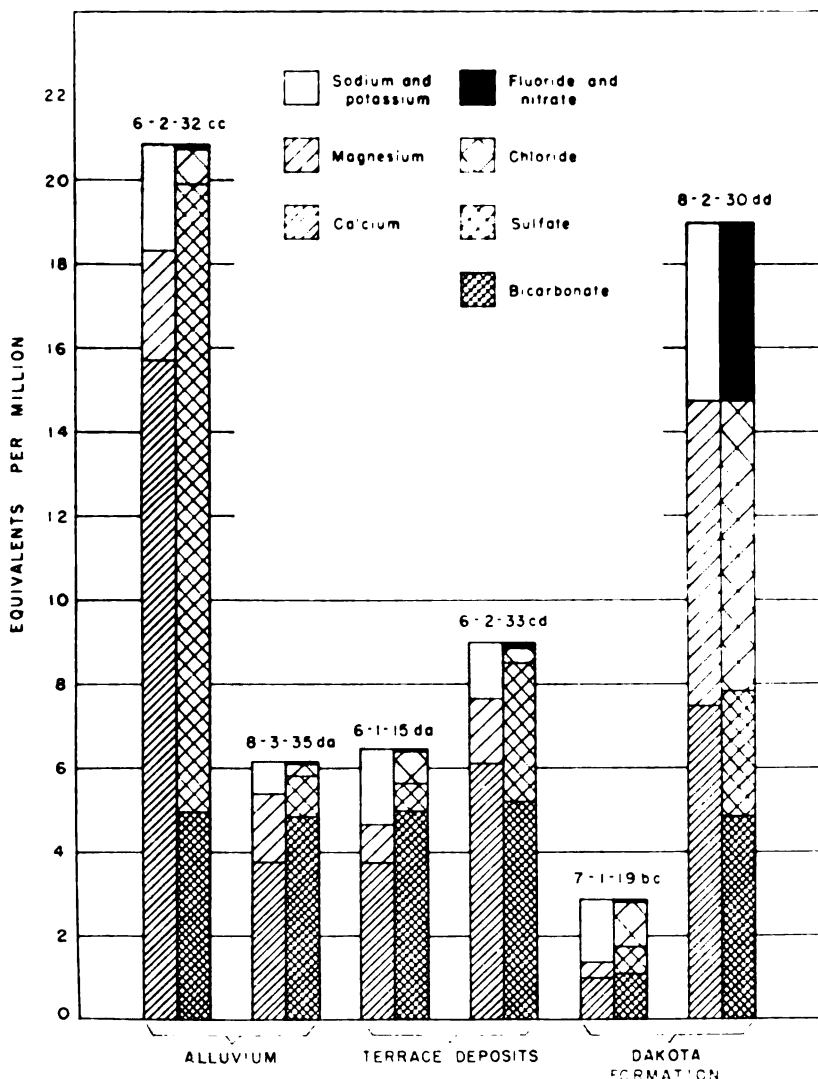


FIG. 8.—Graphic representation of chemical analyses of water from principal water-bearing formations in representative wells in Clay County.

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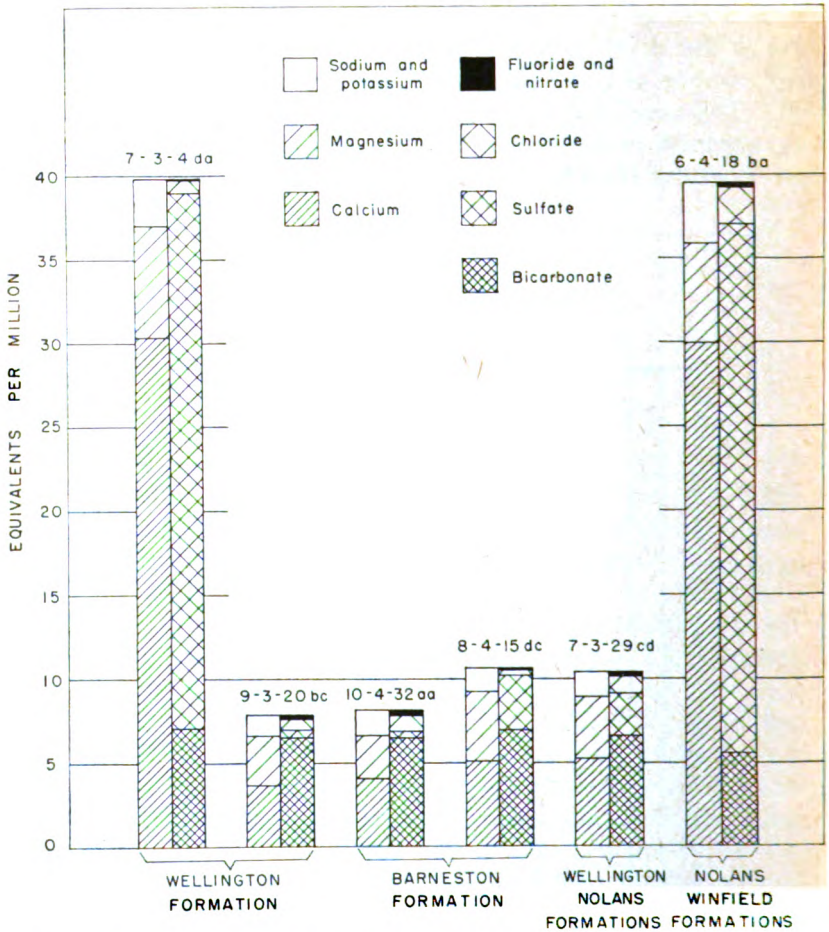


FIG. 9.—Graphic representation of chemical analyses of water from minor water-bearing formations in representative wells in Clay County.

### Chemical Constituents in Relation to Use

The following discussion of the chemical constituents of ground water has been adapted from publications of the State Geological Survey of Kansas. The analyses of samples of water collected in Clay County are given in Table 3.

*Dissolved solids.*—The residue left after a natural water has evaporated consists of rock materials, but may include some organic material and water of crystallization. Waters containing less than 500 parts per million of dissolved solids generally are satisfactory for domestic use, except for difficulties resulting from hardness and,

in some cases, excessive iron content or corrosiveness. Waters having more than 1,000 ppm are as a rule not satisfactory, for they are likely to contain enough of certain constituents to produce a noticeable taste or to make the water unsuitable in some other respect.

The concentration of dissolved solids was less than 500 ppm in 14 of the 33 samples of water collected in Clay County, ranged from 500 to 1,000 ppm in 14 samples, and was more than 1,000 ppm in 5 samples. The water sample having the lowest concentration of dissolved solids (58 ppm) was from a spring (9-2-30bb) in the Dakota Formation. The water sample having the highest concentration of dissolved solids (3,600 ppm) was from well 10-2-17cd1 in the Wisconsinan terrace deposits. This sample had a very high concentration of calcium and sulfate ions.

*Hardness.*—The hardness of water, which is the property of water that generally receives the most attention, is most commonly recognized by its effects when soap is used with water. Hard water is objectionable because it forms with soap a sticky, insoluble curd difficult to remove from containers and fabrics, and because it requires much soap to form a lather. Hard water forms scale in boilers and pipes, which reduces efficiency of heat transfer and may even result in boiler failure. Calcium and magnesium cause most of the hardness of ordinary water, and hence also the greater part of the scale in steam boilers and in other vessels in which water is heated or evaporated.

In addition to total hardness the table of analyses shows carbonate hardness and noncarbonate hardness. The carbonate hardness is that due to the presence of calcium and magnesium bicarbonates. It may be almost completely removed by boiling. This type of hardness has been called temporary hardness. The noncarbonate hardness is due to the presence of the sulfates or chlorides of calcium and magnesium and cannot be removed by boiling, and for this reason has sometimes been called permanent hardness. With reference to use with soap there is no difference between the carbonate and noncarbonate hardness.

Water having a hardness of less than 50 ppm is generally considered soft, and under ordinary circumstances treatment for removal of hardness is not necessary. Hardness between 50 and 150 ppm does not seriously interfere with the use of water for most purposes, but it does increase somewhat the consumption of soap, and its removal by a softening process is profitable for laundries and other industries using large quantities of soap. Hardness of more than 150 ppm can be noticed by most users, and if the hardness is 200

TABLE 3.—Analyses of water from typical wells, springs, and test holes in Clay County  
 Analyzed by H. A. Stoltenberg. Dissolved constituents given in parts per million.\*

Location	Depth, feet	Geologic source	Date of collection	Temper- ature (°F)	Dis- solved solids	Silica (SiO <sub>2</sub> )	Iron (Fe)	Cal- cium (Ca)	Mag- nesium (Mg)	Sodium and potas- sium (Na+K)	Bicar- bonate (HCO <sub>3</sub> )	Sulfate (SO <sub>4</sub> )	Chlo- ride (Cl)	Fluo- ride (F)	Ni- trate (NO <sub>3</sub> )	Hardness as CaCO <sub>3</sub>		
																Total	Car- bonate	Noncar- bonate
6-1-2bac1	65	Terrace deposits	7-20-45	58	345	24	0.61	70	12	37	290	40	16	0.3	1.5	224	224	0
6-1-2cc	56-61	Alluvium	6-17-54	58	373	28	3.8	88	13	29	345	26	18	0.3	0.9	273	273	0
6-1-15da	56-61	Terrace deposits	6-18-54	57	367	28	0.88	75	11	42	303	33	28	0.3	0.8	232	232	0
6-2-3ad	59	Dakota Formation	10-26-54	56	571	14	18	100	17	78	261	54	121	0.3	68	320	214	106
6-2-32ccc	60-65	Alluvium	6-23-54	56	1,320	18	4.1	315	32	57	300	721	30	0.6	2.3	918	246	672
6-2-33cd	56-61	Terrace deposits	6-24-54	.....	525	23	1.2	122	19	30	317	160	13	0.4	1.7	332	260	132
6-4-18ba	77	Nolans and Winfield Limestones	10-26-54	56	2,580	10	11	601	71	83	349	1,510	77	0.8	1.5	1,790	286	1,500
6-4-33ad	103	Barneston Limestone	10-26-54	55	431	13	0.50	76	47	20	459	13	10	0.3	26	382	376	6
7-1-65b	74-76	Dakota Formation	7- 8-54	.....	502	9.5	.....	89	27	53	337	134	20	0.4	3.3	333	276	57
7-1-18beb	65	do	6- 9-54	57	863	8.2	4.9	145	23	106	262	54	171	0.2	217	456	215	241
7-1-19be	50	do	3- 4-54	54	174	7.4	5.4	20	4.6	35	66	31	37	0.2	6.2	69	56	13
7-2-46de	57	Terrace deposits	11-24-53	.....	514	.....	0.18	122	24	28	364	140	16	0.3	6.6	403	290	113
7-3-4da	Spring	Wellington Formation	3- 1-54	57	2,540	17	0.42	608	83	60	432	1,580	25	0.8	1.5	1,860	384	1,500
7-3-20ed	77	Wellington Formation and Nolans Ls.	10-26-54	55	574	18	0.10	107	44	34	407	121	34	0.4	16	448	334	114
7-4-20ed	190	Barneston Limestone	3- 2-54	.....	835	15	0.23	143	67	39	368	311	49	0.6	26	632	302	330
7-4-21bc	.....	.....	.....	.....	.....	.....	.....	.....	.....	.....	.....	.....	.....	.....	.....	.....	.....	.....
7-4-32ee	105	Barneston Limestone	10-26-54	56	668	10	15	107	50	26	429	143	18	0.8	1.7	472	352	120



8-1-13ca.....	40	Terrace deposits.....	10-13-55	889	32	0.77	160	35	44	261	246	79	0.3	71	543	214	329
8-2-30ld.....		Dakota Formation.....	6-9-54	1,140	9.4	3.2	151	88	97	295	146	244	0.6	261	738	242	496
8-3-8db.....	58	Terrace deposits.....	1-31-48	655	30	3.5	144	24	29	445	125	15	0.4	5.3	458	365	93
8-3-35da2.....	16-21	Alluvium.....	6-29-54	330	12	.....	75	20	18	294	47	11	0.4	1.1	269	241	28
9-4-15dc.....	118	Barneston Limestone.....	10-26-54	580	13	9.0	102	51	32	427	157	13	0.6	1.3	464	360	114
9-2-30bb.....	Spring	Dakota Formation.....	10-27-54	58	11	0.23	7.4	1.8	7.6	29	4.9	7.0	0.2	3.6	26	24	2
9-3-1ec.....	93	Barneston Limestone.....	10-26-54	546	18	2.0	100	37	46	454	93	16	0.4	12	402	372	30
9-3-20bc.....	78	Wellington Formation.....	6-9-54	401	11	1.9	75	36	26	404	18	20	0.2	16	335	331	4
9-3-23cc.....	99	Barneston Limestone.....	10-26-54	407	20	1.5	84	38	16	455	9.1	10	0.2	4.4	366	366	0
9-4-21aa.....	144	do.....	10-26-54	371	16	0.26	76	23	32	383	14	14	0.2	6.6	284	284	0
9-4-32adc.....	31-36	Terrace deposits.....	7-1-54	278	22	.....	67	9.5	19	254	28	6.0	0.4	1.3	206	206	0
10-1-17dc.....	100	Dakota Formation.....	5-16-44	119	11	1.2	16	3.8	12	70	13	5.0	0.1	2.2	56	56	0
10-2-17ed.....	55	Terrace deposits.....	6-9-54	3,600	22	29	549	198	292	371	2,170	178	0.9	1.9	2,180	304	1,880
10-2-23bb.....	65	Nolans Limestone.....	6-9-54	476	9.4	1.5	73	31	64	387	46	54	0.2	8.4	310	310	0
10-3-13ba.....	107	Barneston Limestone.....	10-26-54	668	14	1.5	98	40	77	366	90	76	0.3	93	409	300	109
*10-4-5db1.....	57	Terrace deposits.....	8-11-54	605	25	.04	109	28	40	371	56	43	0.2	62	387	304	83
*10-4-5db2.....																	
10-4-32aa.....	134	Barneston Limestone.....	10-26-54	427	16	1.7	84	30	34	398	16	36	0.2	15	333	326	7

a. One part per million is equivalent to one pound of substance per million pounds of water or 8.33 pounds per million gallons of water.

\* Composite sample from two wells.

ppm or more the water is commonly softened. Where municipal water supplies are softened, an attempt is generally made to reduce the hardness to about 80 to 100 ppm. Further softening of a public supply is deemed not worth the additional cost.

Of the 33 samples of water collected in Clay County only 3 samples had a hardness of less than 150 ppm, and these 3 samples were from the Dakota Formation. Six samples had a hardness of 150 to 300 ppm, and 24 samples had a hardness of more than 300 ppm. All but one of the samples collected from wells in Permian rocks had a hardness of more than 300 ppm.

*Iron.*—If the water contains much more than 0.3 ppm of iron, the excess may separate out and settle as a reddish sediment when exposed to air. Iron may be present in sufficient quantity to give a disagreeable taste or to stain cooking utensils; it may be removed by aeration and filtration, but some water requires additional treatment to remove the iron.

Water samples from the test holes in Clay County were pumped by the air-lift method, and the resulting aeration resulted in precipitation of much of the iron. These samples were not analyzed for iron. Only 6 of the 30 samples of water from Clay County that were analyzed for iron contained less than 0.3 ppm.

*Fluoride.*—The fluoride content of water used by children should be known, because fluoride in water has been shown to be associated with the dental defect known as mottled enamel. Mottled enamel may appear on the teeth of children who, during the formation of the permanent teeth, drink water containing excessive amounts of fluoride. According to standards promulgated by the U. S. Public Health Service (1946), water containing more than 1.5 ppm of fluoride should not be used by children. If water contains as much as 4 ppm of fluoride, 90 percent of the children habitually drinking it are likely to have mottled enamel. Concentrations of fluoride of about 1 to 1.5 ppm have been shown to be beneficial in reducing tooth decay, and fluoride is now being added to some municipal supplies to bring the concentration up to about 1 to 1.5 ppm. The water samples collected in Clay County had a fluoride content of less than 1 ppm.

*Chloride.*—Chloride is widely distributed in nature; it is an abundant constituent of sea water and oil-field brines and is dissolved from most rock materials. Chloride has little effect on the suitability of water for ordinary use unless there are enough chloride salts in solution to impart a salty taste or to cause the water to be corrosive. The removal of chloride from water is difficult and expensive.

Water containing chloride concentrations of less than 250 ppm is regarded as satisfactory for domestic uses. Concentrations of chloride salts giving a chloride content between 250 and 500 ppm may impart a slight salty taste, but the water may be used for drinking and for household uses if water of better quality is not available. Cattle have a fairly high tolerance for mineralized water. Although fresh water is preferable, it is reported that cattle can drink water having chloride content of 5,000 ppm.

All the samples of water from Clay County contained less than 250 ppm of chloride.

*Nitrate.*—The use of water containing an excessive amount of nitrate in the preparation of a baby's formula can cause cyanosis (blue baby), or oxygen starvation. Some authorities advocate that water containing more than 45 ppm of nitrate should not be used in formula preparation for infant feeding. Water containing 90 ppm of nitrate is generally considered dangerous to infants, and water containing 150 ppm may cause severe cyanosis. Cyanosis is not produced in adults and older children by the concentration of nitrate found in drinking water. Boiling of water containing excessive nitrate does not render it safe for use by infants; therefore, only water that is known to be free from excessive nitrate should be used for preparing baby formulas.

The nitrate content of the water from some wells is somewhat seasonal, being highest in the winter and lowest in the summer (Metzler and Stoltenberg, 1950). Of the 33 water samples for which analyses are given in Table 3, 27 contained less than 45 ppm of nitrate, 3 contained 45 to 90 ppm, 1 contained 90 to 150 ppm, and 2 contained more than 150 ppm of nitrate.

*Sulfate.*—Sulfate in ground water is derived principally from gypsum (calcium sulfate), and from the oxidation of pyrite (iron sulfide). Magnesium sulfate (Epsom salt) and sodium sulfate (Glauber's salt), if present in sufficient quantity, will impart a bitter taste to the water and may have a laxative effect upon persons who are not accustomed to drinking it.

Of 33 samples of water from Clay County that were analyzed for sulfate, 14 contained less than 50 ppm, 13 contained 51 to 200 ppm, and 6 contained more than 200 ppm.

#### Chemical Constituents in Relation to Irrigation

This discussion of the suitability of water for irrigation is adapted from Agriculture Handbook 60, U. S. Department of Agriculture (U. S. Salinity Laboratory Staff, 1954).

The development and maintenance of successful irrigation projects involve not only the supplying of irrigation water to the land, but also the control of the salinity and alkalinity of the soil. Irrigation practices, drainage conditions, and quality of irrigation water all are involved in salinity and alkali control. Soil that was originally nonsaline and nonalkali may become unproductive if excessive soluble salts or exchangeable sodium are allowed to accumulate because of improper irrigation and soil-management practices or inadequate drainage.

In areas of sufficient rainfall and ideal soil conditions the soluble salts originally present in the soil or added to the soil with water are carried downward by the water and ultimately reach the water table. This process of dissolving and transporting soluble salts by the downward movement of water through the soil is called leaching. If the amount of water applied to the soil is not in excess of the amount needed by plants, there will be no downward percolation of water below the root zone, and mineral matter will accumulate at that level. Likewise, impermeable soil zones near the surface can retard the downward movement of water, resulting in waterlogging of the soil and in deposition of salts. Unless drainage is adequate, attempts at leaching may not be successful, because leaching requires the free passage of water through and away from the root zone.

The characteristics of an irrigation water that seem to be most important in determining its quality are (1) total concentration of soluble salts; (2) relative proportion of sodium to total principal cations (magnesium, calcium, potassium, and sodium); (3) concentration of boron or other elements that may be toxic; and (4) under some conditions, the bicarbonate concentration as related to the concentration of calcium plus magnesium.

For diagnosis and classification of irrigation water, the total concentration of soluble salts can be adequately expressed in terms of electrical conductivity. Electrical conductivity is the measure of the ability of the inorganic salts in solution to conduct an electric current, and is usually expressed in terms of micromhos per centimeter at 25° C. The electrical conductivity can be determined accurately in the laboratory, or an approximation of the electrical conductivity may be obtained by multiplying the total equivalents per million of calcium, sodium, magnesium, and potassium by 100, or by dividing the total dissolved solids in parts per million by a factor which in this area is about 0.64. Table 4 gives the factors for converting parts per million to equivalents per million. In general.

waters having an electrical conductivity less than 750 micromhos per centimeter are satisfactory for irrigation insofar as salt content is concerned, although salt-sensitive crops such as strawberries, green beans, and red clover may be adversely affected by irrigation water having an electrical conductivity in the range of 250 to 750 micromhos per centimeter. Waters in the range of 750 to 2,250 micromhos per centimeter are widely used, and satisfactory crop

TABLE 4.—Factors for converting parts per million of mineral constituents to equivalents per million.<sup>1</sup>

Cation	Conversion factor	Anion	Conversion factor
Ca <sup>++</sup> .....	0.0499	HCO <sub>3</sub> <sup>-</sup> .....	0.0164
Mg <sup>++</sup> .....	.0822	SO <sub>4</sub> <sup>-</sup> .....	.0208
Na <sup>+</sup> .....	.0435	Cl <sup>-</sup> .....	.0281
		NO <sub>3</sub> <sup>-</sup> .....	.0161
		F <sup>-</sup> .....	.0526

1. Equivalents per million equals parts per million multiplied by conversion factor. For example, 487 ppm of calcium × 0.0499 = 24.3 epm.

growth is obtained under good management and favorable drainage conditions, but saline conditions will develop if leaching and drainage are inadequate. Use of waters having a conductivity greater than 2,250 micromhos per centimeter is the exception, and very few instances can be cited where such waters have been used successfully.

In the past the relative proportion of sodium to other cations in irrigation water has been expressed simply as the percentage of sodium, generally called "percent sodium". According to the U. S. Department of Agriculture, however, the sodium-adsorption ratio (SAR), used to express the relative activity of sodium ions in exchange reactions with soil, is a much better measure of the suitability of water for irrigation. The sodium-adsorption ratio may be

determined by the formula 
$$SAR = \frac{Na^+}{\sqrt{\frac{Ca^{++} + Mg^{++}}{2}}}$$
 where the ionic

concentrations are expressed in equivalents per million. The sodium-adsorption ratio may be determined also by use of the nomogram shown in Figure 10. In using the nomogram to determine the sodium-adsorption ratio of a water, the concentration of sodium expressed in equivalents per million is plotted on the left scale (A),

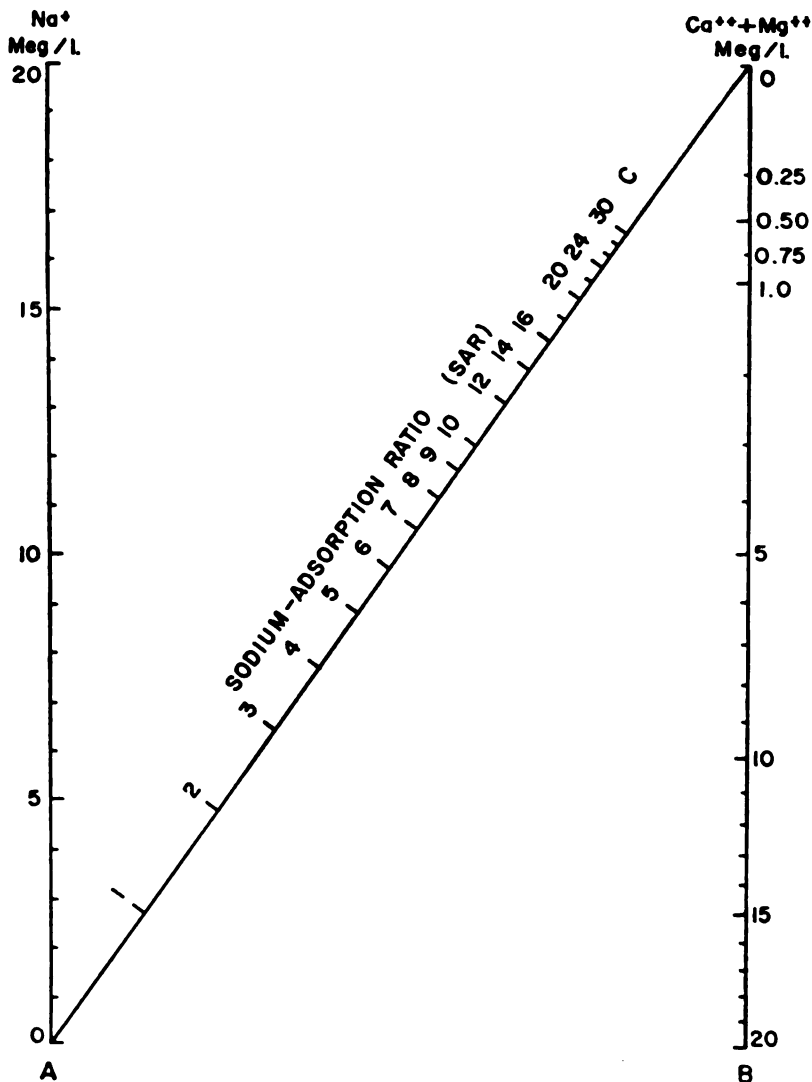


FIG. 10.—Nomogram for determining sodium-adsorption ratio of irrigation water.

and the concentration of calcium plus magnesium expressed in equivalents per million is plotted on the right scale (B). The point at which a line connecting these two points intersects the sodium-adsorption-ratio scale (C) indicates the sodium-adsorption ratio of the water. When the sodium-adsorption ratio and the electrical conductivity of a water are known, the suitability of the water

for irrigation can be determined by plotting these values on the diagram shown in Figure 11. Low-sodium water (S1) can be used for irrigation on almost all soils with little danger of the development of harmful levels of exchangeable sodium. Medium-sodium water

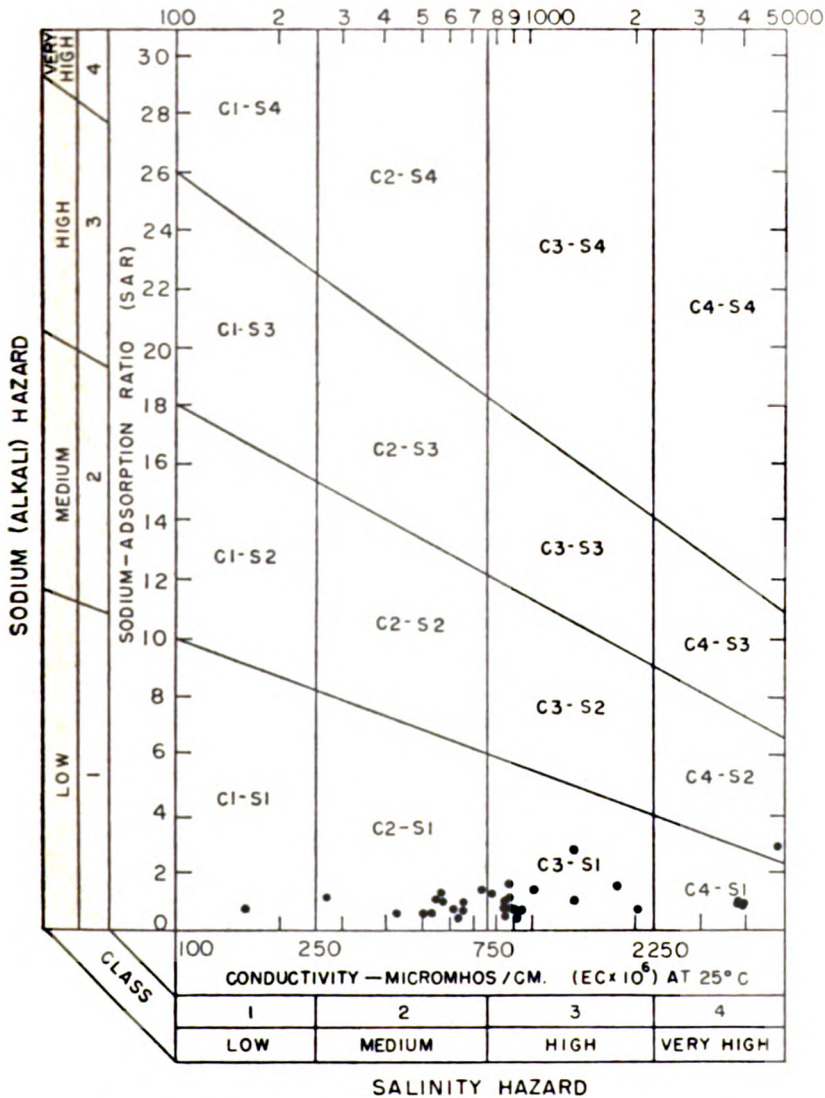


FIG. 11.—Diagram showing classification of water in Clay County for irrigation use.

(S2) will present an appreciable sodium hazard in certain fine-textured soils, especially under low-leaching conditions. This water may be safely used on coarse-textured or organic soils having good permeability. High-sodium water (S3) may produce harmful levels of exchangeable sodium in most soils and will require special soil management such as good drainage, high leaching, and additions of organic matter. Very high sodium water (S4) is generally unsatisfactory for irrigation unless special action is taken, such as addition of gypsum to the soil.

Low-salinity water (C1) can be used for irrigation of most crops on most soils with little likelihood that soil salinity will develop. Medium-salinity water (C2) can be used if a moderate amount of leaching occurs. Crops having moderate salt tolerances, such as potatoes, corn, wheat, oats, and alfalfa, can be irrigated with (C2) water without special practices. High-salinity water (C3) cannot be used on soils with restricted drainage. Very high salinity water (C4) can be used only on certain crops and then only if special practices are followed.

Boron is essential to normal plant growth, but the quantity required is very small. Crops vary greatly in their boron tolerances, but in general it may be said that the ordinary field crops common to Kansas are not adversely affected by boron concentrations of less than 1 ppm.

Prolonged use, under adverse conditions, of water having a high concentration of bicarbonate could have an undesirable effect upon the soil texture and plant growth.

Water samples (1 composite) from 33 wells and test holes are classified on Figure 11 to show the suitability of the water for irrigation use. The wells or test holes corresponding to points shown in Figure 11 are given in Table 5. Of 32 samples of water from Clay County, 2 are of low alkali hazard and low salinity hazard and are very good irrigation waters; 7 analyses indicate a low alkali hazard and medium salinity hazard, and the waters are satisfactory for irrigation of most crops grown in Clay County; 20 samples represent water of low alkali hazard but high salinity hazard, but as most of these samples fall within the lower part of the high-salinity area, only moderate precautionary measures would be necessary. Three samples were of low or medium alkali hazard but very high salinity hazard, and such waters are unfit for irrigation use. In general, the aquifers in Clay County that produce water that is unfit for irrigation do not produce large enough quantities to be considered for irrigation.



TABLE 5.—Sodium-adsorption ratios (SAR), conductivity (C), and class of water for irrigation for wells plotted on Figure 11.

Well number	SAR	C	Class	number	SAR	C	Class
6-1-2bac1	1.07	537	Z1-C2	8-1-13ca	0.82	1,370	Z1-C3
6-1-2ec	.75	563	Z1-C2	8-2-30dd	1.56	1,780	Z1-C3
6-1-15da	1.21	573	Z1-C2	8-3-84b	.59	900	Z1-C3
6-2-3ad	1.80	892	Z1-C3	8-3-35da <sub>2</sub>	.47	515	Z1-C2
6-2-32cc	.78	2,070	Z1-C3	8-4-15le	.65	906	Z1-C3
6-2-33cd	.62	820	Z1-C3	9-2-30bb	.65	90	Z1-C1
6-4-18ba	.85	3,950	Z1-C4	9-3-1cc	1.00	853	Z1-C3
6-4-33ad	.44	673	Z1-C2	9-3-20bc	.61	626	Z1-C2
7-1-6bb	1.26	784	Z1-C3	9-3-23cc	.37	635	Z1-C2
7-1-15beb	2.16	1,330	Z1-C3	9-4-21aa	.83	579	Z1-C2
7-1-19bc	1.8	272	Z1-C2	9-4-32ade	.58	434	Z1-C2
7-2-4dde	.60	803	Z1-C3	10-1-17de	.70	186	Z1-C1
7-3-4da	.60	3,970	Z1-C4	10-2-17ed <sub>1</sub>	2.72	5,620	Z2-C4
7-3-29cd	.69	896	Z1-C3	10-2-23bb	1.58	743	Z1-C2
*7-4-20ad-21be)	.68	926	Z1-C3	10-3-13ba	1.65	1,040	Z1-C3
7-4-32cc	.52	887	Z1-C3	10-4-32aa	.75	667	Z1-C2

\* Composite sample from wells 7-4-20ad and 7-4-21be.

HYDROLOGIC PROPERTIES OF WATER-BEARING MATERIALS

The quantity of water that a water-bearing formation will yield to wells depends upon the hydrologic properties of the material from which the wells produce. The two hydrologic properties of greatest significance are the coefficients of transmissibility (T) and storage (S) and are used in making quantitative estimates of water available in an aquifer and of future water-level decline that will result from continued pumping. Controlled aquifer tests in the field provide the data required to compute these coefficients.

The coefficient of transmissibility (T) may be defined as the number of gallons of water at the prevailing temperature that will move in 1 day through a vertical strip of the aquifer 1 foot wide, having a height equal to the full thickness of the aquifer, under a hydraulic gradient of 1 foot per foot, or it is the number of gallons of water that will move in 1 day through a cross-sectional area equal to the saturated thickness of the aquifer and 1 mile wide under a hydraulic gradient of 1 foot per mile. The coefficient of storage (S) may be defined as the change in volume of water stored per unit surface area of aquifer per unit change in head. Under water-table conditions the coefficient of storage (S) is practically the same as the specific yield of the aquifer.

The coefficient of permeability (P) of an aquifer is the discharge per unit of area per unit of hydraulic gradient. It may be measured in terms of the number of gallons of water a day, at 60° F., con-

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ducted laterally through each mile of aquifer under investigation (measured at right angles to the direction of flow) for each foot of thickness of the aquifer, and for each foot per mile of hydraulic gradient. The field coefficient of permeability ( $P_f$ ) is the same except that it is measured at the prevailing temperatures of the water rather than at 60° F. The coefficient of permeability may be expressed by the formula  $P_f = \frac{T}{m}$  where  $P_f$  is the field coefficient of permeability,  $T$  is the coefficient of transmissibility, and  $m$  is the thickness of the aquifer, in feet.

#### AQUIFER-TEST DETERMINATION OF TRANSMISSIBILITY AND PERMEABILITY

The coefficients of transmissibility and permeability of the alluvium and terrace deposits in the Republican River valley in the vicinity of Clifton were determined by an aquifer test using well 6-1-2ac owned by Mr. F. Turner and operated by Mr. H. Rhodes. Values for the coefficient of transmissibility ( $T$ ), storage ( $S$ ), and permeability ( $P$ ) were computed from the aquifer-test data by formulas developed by C. V. Theis and C. E. Jacob.

##### Theis Recovery Method of Determining Transmissibility

The recovery method of computing transmissibility ( $T$ ) developed by Theis (1935) utilizes a series of measurements of the water level in a well after a period of pumping.

The Theis recovery formula is expressed as:  $T = \frac{264Q \log_{10} t/t'}{s'}$

in which  $T$  is the coefficient of transmissibility, in gallons per day per foot,  $Q$  is the pumping rate, in gallons a minute,  $t$  is the time since pumping started, in minutes,  $t'$  is the time since pumping stopped, in minutes, and  $s'$  is the residual drawdown in the pumped well, in feet, at time  $t'$ .

The residual drawdown ( $s'$ ) is computed by subtracting the static-water-level measurement from the depth-to-water measurements at time  $t'$  after pumping ceases.

The proper ratio of  $\log_{10} t/t'$  to  $s'$  is determined graphically by plotting  $\log_{10} t/t'$  on the logarithmic coordinate and  $s'$  on the arithmetic coordinate of semilogarithmic paper. If  $\log_{10} t/t'$  is taken over one log cycle it will become unity. The formula may then be expressed as  $T = \frac{264Q}{\Delta s}$ , where  $\Delta s$  is the difference in drawdown

over one log cycle. In practice, water levels at time  $t'$  may be plotted and the residual drawdown need not be computed. Because the value of  $T$  is directly related to the slope of the line formed by plotting water level against  $t/t'$ , the selection of the proper points on the plot to determine  $\Delta s$  is very important. Theoretically, the relation  $s'$  to  $t/t'$  should plot as a straight line that passes through the point where residual drawdown is 0 and where  $t/t'$  approaches unity. In unconsolidated deposits this is not always true, as the observed line is nearly always a curve and seldom passes through the point of origin. Generally the early points on the recovery curve are erratic and do not fall in a straight line. This may be due in part to head loss in the well or to the surge of the column of water, which falls back into the well after the pump is shut off. For this reason it would seem that the points on the curve corresponding to the late part of the recovery period would be the most reliable and should be used in determining  $\Delta s$ . For many aquifer tests the recovery data will not plot on a line that passes through the point where  $s' = 0$  and will not plot in a straight line in the latter part of the recovery period, and it is difficult to establish a line that will determine  $\Delta s$ . If the latest data where  $t/t'$  approaches 1 do not fall on a straight line, they should not be used. If the earlier points fall on a straight line they should be used, but if there is an appreciable curve to the line almost any value can be obtained for  $\Delta s$ , and hence the value of  $T$  can be considerably in error.

#### Straight-line Graphical Method of Determining Transmissibility and Storage Coefficient

The Theis (1935) nonequilibrium formula requires the use of a "type curve" on which the test-data curve is superimposed to determine the coefficients of transmissibility and storage. Cooper and Jacob (1946) devised a straight-line graphical method that does not require the type curve to accomplish the same purposes. In this method the data are plotted on semilogarithmic paper, on which the plotted points should fall in a straight line. Three equations were devised for determination of coefficients of transmissibility and storage by a distance-drawdown graph, a time-drawdown graph, and a composite graph. In the distance-drawdown graph, drawdown is plotted on the arithmetic scale and distance of observation wells in feet from the pumped well is plotted on the logarithmic scale. The formulas for coefficients of transmissibility and storage are:

$$T = \frac{264 Q}{\Delta s} \text{ and } S = \frac{0.3 T t}{r_o^2}.$$

In the time-drawdown graph, drawdown is plotted on the arithmetic scale and time in days is plotted on the logarithmic scale. The formulas for coefficients of transmissibility and storage are:

$$T = \frac{264 Q}{\Delta s} \text{ and } S = \frac{0.3Tt_0}{r^2}.$$

In the composite drawdown graph, drawdown is plotted on the arithmetic scale and the value  $t/r^2$  is plotted on the logarithmic scale. In this method the plots of all observation wells should fall in a straight line. The formulas for coefficients of transmissibility and storage are:

$$T = \frac{264 Q}{\Delta s} \text{ and } S = 0.3T \times (t/r^2)_0.$$

In the above formulas, the symbols are:

**T** = coefficient of transmissibility, in gallons per day per foot

**S** = coefficient of storage

**Q** = rate of discharge of pumped well, in gallons per minute

**$\Delta s$**  = drawdown over one log cycle

**t** = time since pumping started, in days

**r** = distance from pumped well, in feet

**$t_0$**  = value of t where drawdown = 0

**$r_0$**  = value of r where drawdown = 0

**$(t/r^2)_0$**  = value of  $t/r^2$  where drawdown = 0

#### Turner Aquifer Test

An aquifer test was made in the fall of 1955 in which well 6-1-2ac was used. This well yields water from late Wisconsinan terrace deposits in the valley of Republican River. The well is of gravel-wall construction and is 18 inches in diameter and 70 feet deep. Three observation wells, a, b, and c, were drilled in a line at distances of 25, 50, and 100 feet respectively from the pumped well. The well was pumped at a rate of 800 gpm from 9:25 a. m. until 9:55 a. m., when pumping stopped for a period of 45 minutes, owing to a power failure. Pumping was resumed at 10:40 a. m. and continued until 1:13 p. m. Water-level measurements made during the pumping period and values for t are given in Table 6.

In Figure 12 the depth-to-water measurements in the pumped well are plotted against time since pumping started. A line drawn

TABLE 6.—Water-level measurements made during pumping and recovery periods in well 6-1-2ac, and values for (t).

Time	Time since pumping started, t, in minutes	Depth to water, in feet
9:20 a.m.	Static water level	22 21
9:25	Pumping started	
9:25:20	0 3	31 35
9:25:45	0 75	32 87
9:26:15	1 25	33 55
9:26:45	1 75	33 99
9:27:15	2 25	34 17
9:28	3	34 44
9:29	4	34 77
9:29:30	4 5	34 86
9:30	5	34 95
9:30:30	5 5	35 04
9:31	6	35 11
9:31:30	6 5	35 19
9:32	7	35 26
9:32:30	7 5	35 32
9:33	8	35 36
9:33:30	8 5	35 42
9:34	9	35 48
9:34:30	9 5	35 52
9:35	10	35 57
9:36	11	35 67
9:37	12	35 78
9:38	13	35 83
9:39	14	35 89
9:40	15	35 95
9:45	20	36 15
9:50	25	36 47
9:55	Pump off (power failure)	
10:39		22 98
10:40	Pumping started	
10:40:30	75 5	33 87
10:41	76	34 77
10:41:30	76 5	35 16
10:42	77	35 42
10:42:30	77 5	35 67
10:43	78	35 81
10:44	79	36 00
10:44:30	79 5	36 10
10:45	80	36 20
10:45:30	80 5	36 29
10:46	81	36 38
10:46:30	81 5	36 44
10:47	82	36 52

TABLE 6.—Water-level measurements made during pumping and recovery periods in well 6-1-2ac, and values for (t).—Continued.

Time	Time since pumping started, t, in minutes	Depth to water, in feet
10:47:30 a.m.	82.5	36.57
10:48	83	36.66
10:48:30	83.5	36.70
10:49	84	36.74
10:49:30	84.5	36.78
10:50	85	36.83
10:50:30	85.5	36.87
10:51	86	36.90
10:51:30	86.5	36.93
10:52	87	36.95
10:52:30	87.5	36.97
10:53	88	37.00
10:53:30	88.5	37.01
10:54	89	37.03
10:54:30	89.5	37.06
10:55	90	37.09
10:57	92	37.16
10:59	94	37.23
11:02	97	37.35
11:05	100	37.48
11:10 a.m.	105	37.61
11:15	110	37.78
11:20	115	37.90
11:25	120	38.00
11:30	125	38.13
11:35	130	38.22
11:40	135	38.30
11:50	140	38.48
12:00 noon	155	38.55
12:10 p.m.	165	38.60
12:20	175	38.67
12:30	185	38.81
12:40	195	38.86
12:50	205	38.90
1:00	215	38.94
1:10	225	38.96
1:13	228 Pump off	.....
1:14	229	26.83
1:15	230	26.38
1:16	231	26.17
1:17	232	26.01
1:18	233	25.90
1:19	234	25.80
1:20	235	25.71
1:21	236	25.63

TABLE 6.—Water-level measurements made during pumping and recovery periods in well 6-1-2ac, and values for (t).—Concluded.

Time	Time since pumping started, t, in minutes	Depth to water, in feet
1:22 p.m.	237	25 57
1:23	238	25 49
1:24	239	25 44
1:25	240	25 39
1:26	241	25 35
1:28	243	25 26
1:30	245	25 16
1:32	247	25 09
1:34	249	25 02
1:36	251	24 95
1:38	253	24 90
1:41	256	24 80
1:44	259	24 74
1:47	262	24 66
1:50	265	24 58
1:54	269	24 52
1:58	273	24 45
2:02	277	24 38
2:06	281	24 32
2:10	285	24 24
2:15	290	24 15
2:21	296	24 09
2:28	303	24 01
2:35	310	23 98
2:45	320	23 91
2:55	330	23 81
3:05	340	23 74
3:15	350	23 66
3:25	360	23 60
3:35	370	23 54
3:50	385	23 46
4:05	400	23 40
4:20	415	23 32

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through these points gives a value for  $\Delta s$  of 2.18. Applying the Cooper-Jacob formula,  $T = \frac{(264)(800)}{2.18} = 97,000$  gpd per foot.

The saturated thickness is 43 feet, and from the relation  $P = \frac{T}{m}$ ,  $P = 2,250$ . The pumped well had a drawdown of 17 feet while pumping 800 gpm; hence the specific capacity of the well was 47 gpm per foot of drawdown.

Depth-to-water measurements were made in observation wells a, b, and c during the test and are shown in Tables 7, 8, and 9.

Early in the test it was evident that observation well a was not responding to pumping as it should. The water-level measurements in Tables 7 and 8 indicate that the water level in observation well a was drawing down less than that in observation well b throughout the test. This was probably due to the lenticular nature of the aquifer rather than to any defect in the observation well, as tests made during the power failure indicated that the observation wells were open and functioning normally.

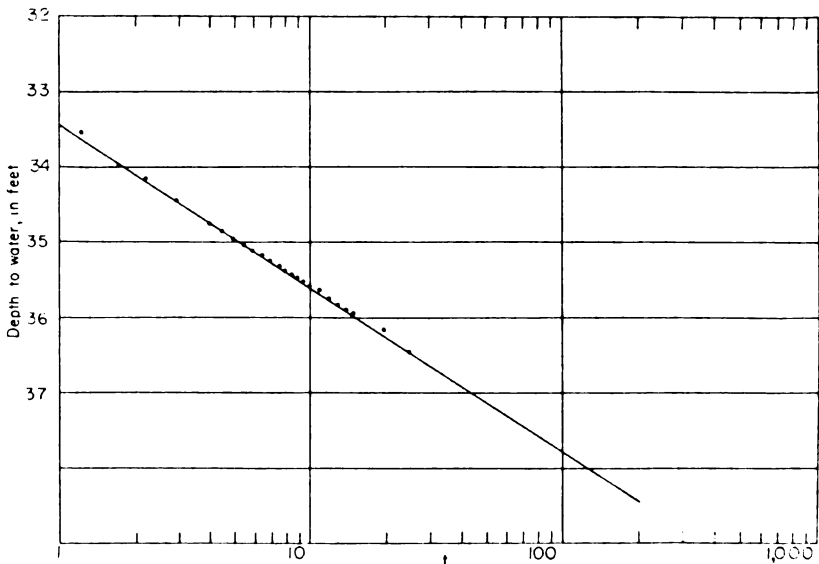


FIG. 12.—Depth to water in well 6-1-2ac plotted against time since pumping started.



TABLE 7.—Water-level measurements in observation well a, 25 feet from pumped well 6-1-2ac.

Time	Time since pumping started, t, in minutes	Depth to water, in feet
9:20 a.m. ....	Static water level	23 40
9:25 .....	Pumping started	
9:26 .....	1	24 31
9:27 .....	2	24 69
9:29 .....	4	25 05
9:30 .....	5	25 23
9:32 .....	7	25 56
9:34 .....	9	25 72
9:36 .....	11	25 93
9:38 .....	13	26 08
9:40 .....	15	26 23
9:42 .....	17	26 33
9:44 .....	19	26 47
9:46 .....	21	26 60
9:48 .....	23	26 69
9:50 .....	25	26 78
9:53 .....	28	26 98
9:55 .....	Pumping stopped	
10:38 .....		24 24
10:40 .....	Pumping started	
10:41 .....	1	25 12
10:42 .....	2	25 37
10:43 .....	3	25 75
10:44 .....	4	25 84
10:47 .....	7	26 05
10:49 .....	9	26 30
10:51 .....	11	26 50
10:53 .....	13	26 67
10:55 .....	15	26 84
10:57 .....	17	26 98
11:00 .....	20	27 10
11:05 .....	25	27 26
11:10 .....	30	27 48
11:15 .....	35	27 68
11:20 .....	40	27 86
11:25 .....	45	27 99
11:30 .....	50	28 13
11:35 .....	55	28 24
11:40 .....	60	28 32
11:45 .....	65	28 42
11:50 .....	70	28 50

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TABLE 7.—*Water-level measurements in observation well a, 25 feet from pumped well 6-1-3ac.—Concluded.*

Time	Time since pumping started, t, in minutes	Depth to water, in feet
11:55 a.m.....	75	28.58
12:00.....	80	28.64
12:10 p.m.....	90	28.69
12:20.....	100	28.80
12:30.....	110	28.89
12:40.....	120	28.98
12:50.....	130	29.03
1:00.....	140	29.09
1:10.....	150	29.15

TABLE 8.—Water-level measurements in observation well b, 50 feet from pumped well 6-1-2ac.

Time	Time since pumping started, t, in minutes	Depth to water, in feet
9:20 a.m.	Static water level	23.34
9:25	Pumping started	.....
9:26:30	1.5	26.63
9:27:30	2.5	27.12
9:28:15	3.25	27.32
9:29		27.48
9:30	5	27.61
9:31	6	27.71
9:32	7	27.82
9:33	8	27.90
9:34	9	27.98
9:35	10	28.04
9:36	11	28.11
9:38	13	28.19
9:40	15	28.34
9:42	17	28.43
9:44	19	28.51
9:46	21	28.59
9:48	23	28.67
9:50	25	28.75
9:53	28	28.84
9:54	Pumping stopped	.....
10:38		24.00
10:40	Pumping started	.....
10:41:30	1.5 76.5	27.48
10:43	3 78	27.98
10:44	4 79	28.22
10:45	5 80	28.36
10:46	6 81	28.47
10:48	8 83	28.84
10:50	10 85	28.89
10:52	12 87	28.90
10:55	15 90	29.04
10:58	18 93	29.15
11:00	20 95	29.23
11:05	25 100	29.38
11:10	30 105	29.51
11:15	35 110	29.66
11:20	40 115	29.77
11:25	45 120	29.85
11:30	50 125	29.95

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## Geological Survey of Kansas

TABLE 8.—Water-level measurements in observation well b, 50 feet from pumped well 8-1-2ac.—Concluded

Time	Time since pumping started, t, in minutes		Depth to water, in feet
11:40 a.m.....	60	135	30.11
11:50.....	70	145	30.23
12:00 noon.....	80	155	30.33
12:10 p.m.....	90	165	30.45
12:20.....	100	175	30.54
12:30.....	110	185	30.63
12:40.....	120	195	30.69
12:50.....	130	205	30.73
1:00.....	140	215	30.79

TABLE 9.—Water-level measurements in observation well c, 100 feet from pumped well 6-1-2ac.

Time	Time since pumping started, t, in minutes	Depth to water, in feet
9:20 a.m.	Static water level	23 53
9:25	Pumping started	
9:26	1	23 57
9:27	2	23 64
9:28	3	23 73
9:29	4	23 81
9:30	5	23 88
9:31	6	23 97
9:32	7	24 04
9:33	8	24 12
9:34	9	24 18
9:35	10	24 24
9:36	11	24 31
9:37	12	24 38
9:38	13	24 44
9:39	14	24 50
9:40	15	24 56
9:42	17	24 66
9:44	19	24 78
9:46	21	24 90
9:48	23	25 00
9:50	25	25 09
9:53	28	25 21
9:54	Pumping stopped	
10:38		24 37
10:40	Pumping started	
10:41	1	24 39
10:42	2	24 45
10:43	3	24 52
10:44	4	24 58
10:45	5	24 66
10:46	6	24 73
10:47	7	24 79
10:48	8	24 86
10:49	9	24 93
10:50	10	24 99
10:52	12	25 10
10:54	14	25 21
10:56	16	25 23
10:58	18	25 42
11:00	20	25 51
11:03	23	25 63
11:06	26	25 76
11:09	29	25 86

TABLE 9.—Water-level measurements in observation well c, 100 feet from pumped well 6-1-2ac—Continued.

Time	Time since pumping started, t, in minutes	Depth to water, in feet
11:12 a.m.	32	26.00
11:15	35	26.09
11:20	40	26.25
11:25	45	26.38
11:30	50	26.52
11:35	55	26.62
11:40	60	26.71
11:45	65	26.94
11:50	70	26.95
11:55	75	27.02
12:00 noon	80	27.07
12:10 p.m.	90	27.20
12:20	100	27.30
12:30	110	27.43
12:40	120	27.49
12:50	130	27.57
1:00	140	27.59
1:10	150	27.67
1:13	<b>Pumping stopped</b>	
1:14	1	27.62
1:15	2	27.58
1:16	3	27.53
1:17	4	27.45
1:18	5	27.42
1:19	6	27.39
1:20	7	27.29
1:21	8	27.28
1:22	9	27.22
1:23	10	27.17
1:24	11	27.13
1:25	12	27.09
1:26	13	27.05
1:27	14	27.01
1:28	15	26.97
1:29	16	26.93
1:30	17	26.88
1:31	18	26.86
1:32	19	26.83
1:33	20	26.78
1:34	21	26.76
1:35	22	26.73
1:36	23	26.69
1:37	24	26.67
1:38	25	26.64
1:40	27	26.58

**TABLE 9.—Water-level measurements in observation well c, 100 feet from pumped well 6-1-2ac. —Continued.**

Time	Time since pumping stopped, t, in minutes	Depth to water, in feet
1:45 p.m. ....	32	26.45
1:50 .....	37	26.34
1:55 .....	42	26.23
2:00 .....	47	26.14
2:05 .....	52	26.05
2:10 .....	57	25.97
2:15 .....	62	25.89
2:20 .....	67	25.82
2:25 .....	72	25.75
2:30 .....	77	25.69
2:40 .....	87	25.58
2:50 .....	97	25.49
3:00 .....	107	25.39
3:10 .....	117	25.31
3:20 .....	127	25.23
3:31 .....	138	25.14
3:45 .....	152	25.05
4:00 .....	167	25.00
4:15 .....	182	24.88

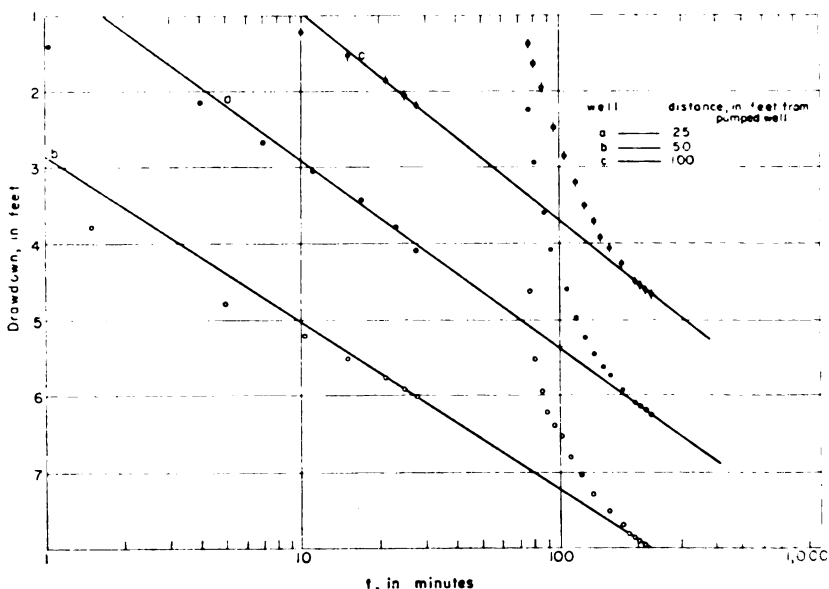


FIG. 13.—Drawdown in observation wells a, b, and c plotted against time since pumping started, during test using well 6-1-2ac.

The drawdowns in observation wells a, b, and c are plotted against the time since pumping started (Fig. 13). Lines through the latest points prior to the power failure pass through the latest points in the test in all the plots. Applying the Cooper-Jacob formula, values for the coefficient of transmissibility from the fluctuations in observation wells a, b, and c are 87,000, 97,000, and 79,000 gpd per foot respectively. The storage coefficient, from measurements in wells a and c, was 0.03 and 0.02, respectively, indicating water-table conditions.

A line through the latest points in the recovery curve for well c (Fig. 14) gives a value for  $s = 2.17$ . Applying the formula

$$T = \frac{(264)(800)}{2.17}$$

gives a result of 97,000 gpd/per foot.

**Summary.**—The values of  $T$  obtained from the aquifer test using well 6-1-2ac ranged from 79,000 to 97,000 gpd per foot. The lowest value was 79,000 gpd per foot from observation well a and the highest value, 97,000 gpd per foot, from the pumped well, observation well b, and the recovery curve for observation well c. The test is not conclusive, but 90,000 gpd per foot is probably about the cor-



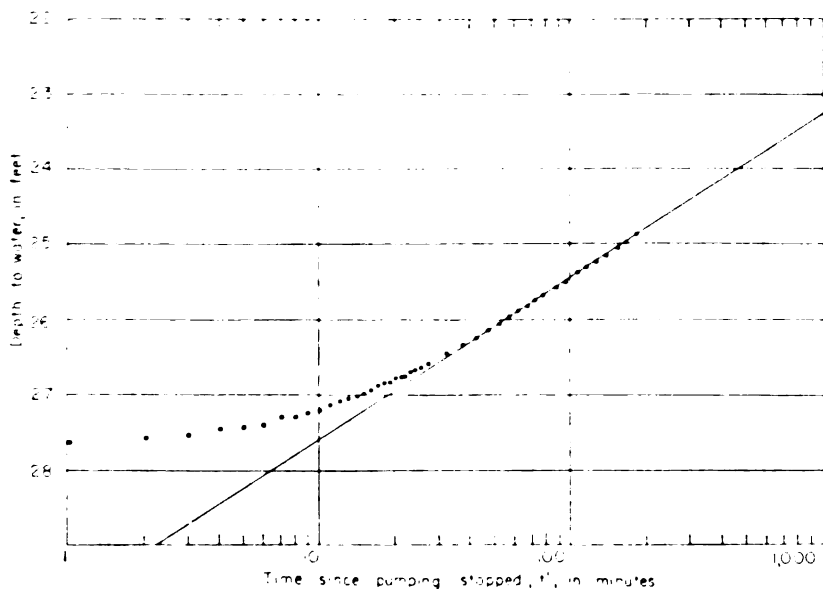


FIG. 14.—Depth to water in observation well c plotted against time since pumping stopped, during test using well 6-1-2ac.

rect value of T in the area of the well. The test shows that the aquifer is not homogeneous and that conditions are variable within short distances. The values for S are too small to have any significance, because of the shortness of the period of pumping.

### ROCK UNITS AND THEIR WATER-BEARING PROPERTIES

#### PERMIAN SYSTEM

#### Wolfcampian Series—Chase Group

#### *Barneston Limestone*

In Clay County the Barneston Limestone crops out in the vicinity of Wakefield in the valley walls of Republican River and its tributaries. The Barneston Limestone contains the thickest and most resistant limestone beds in the county, and in the area to the east it is one of the principal scarp-forming rocks of the Flint Hills. In Clay County the Barneston Limestone occupies a fairly low topographic position and does not form a prominent scarp. The members of the Barneston Limestone, in ascending order, are the Florence Limestone, the Oketo Shale, and the Fort Riley Limestone members.

3—6124

The Florence Limestone member is very easily recognized in Clay County because it is the only rock unit more than a few feet thick that contains a preponderance of chert or flint. The member is primarily a series of beds of limestone and beds of chert. The chert beds are generally of about the same thickness as the limestone beds. At most places a very dark gray or black impure limestone bed about 2 feet thick forms the base of the Florence Limestone member, and a persistent shale bed about 2 feet thick occurs in the upper part of the member. The Florence contains abundant fossils of many varieties. The thickness of the member ranges from 35 to 40 feet.

In Clay County the Oketo Shale member averages about 5 feet in thickness. It is composed chiefly of fossiliferous hard gray calcareous shale.

The lower few feet of the Fort Riley Limestone member is composed chiefly of thin-bedded yellow limestone. The limestone beds are overlain by a few feet of gray shale, which is easily confused with the underlying Oketo Shale member unless the Florence Limestone member is exposed also. Next above this shale bed is a massive bed of yellow limestone about 5 feet thick. This bed, which weathers to a pitted surface, is the rimrock, characteristic of nearly all exposures of the Fort Riley Limestone member (Pl. 4A). The rimrock bed is overlain by 15 to 20 feet of thin-bedded sandy yellow limestone and thin calcareous shale beds. In some exposures, a massive bed that has much the same appearance as the rimrock occurs in the upper part of the member. The average thickness of the member is about 35 feet.

Both the Florence and Fort Riley Limestone members are important aquifers in Clay County, but the Oketo Shale member does not yield appreciable quantities of water. A properly constructed well that penetrates the entire thickness of the Barneston Limestone may produce as much as 350 gpm.

In areas where the Barneston Limestone is deeply buried it may yield water that is too mineralized for domestic use. Several wells are known to obtain water of satisfactory quality from the Barneston Limestone at a depth of almost 200 feet, however.

#### *Doyle Shale*

The Doyle Shale is composed of two shale members separated by a limestone member—in ascending order, the Holmesville Shale, the Towanda Limestone, and the Gage Shale members.

A



PLATE 4. A, Rimrock formed by Fort Riley Limestone member, SE $\frac{1}{4}$  NE $\frac{1}{4}$  sec. 27, T. 9 S., R. 4 E.; B, Exposure of Winfield Limestone, SE $\frac{1}{4}$  NE $\frac{1}{4}$  sec 19, T. 10 S., R. 4 E.

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The Holmesville Shale member consists of 20 to 25 feet of vari-colored calcareous shale. Calcite veinlets and thin impure limestone beds are present in the lower few feet of the member in most places.

The Towanda Limestone member is composed chiefly of platy and nodular beds of gray to yellow limestone. In most weathered exposures the member has a very broken and deformed appearance, and the upper limestone beds may be limonite stained and may contain many calcite veinlets. The average thickness of the Towanda Limestone member is about 10 feet.

The Gage Shale member, which in Clay County has an average thickness of about 35 feet, is composed of a lower unfossiliferous varicolored shale and an upper very fossiliferous gray and yellow shale.

The Doyle Shale is not an important aquifer in Clay County, but locally the Towanda Limestone member may yield as much as 5 gpm to properly constructed wells.

#### *Winfield Limestone*

The members of the Winfield Limestone, in ascending order, are the Stovall Limestone, the Grant Shale, and the Cresswell Limestone members.

The Stovall Limestone member consists of a single bed of gray to tan hard cherty limestone about 1 foot thick. It is nonresistant to weathering and is not well exposed in many places. Echinoid spines are the most common fossils in the Stovall. It is easily recognized by its uniform thickness and chert content.

The Grant Shale member has a uniform thickness of about 10 feet in Clay County. It is composed chiefly of tan and gray shale but includes some darker shale in the upper part. It contains many fossil brachiopods, and many quartz geodes are present in the upper part.

The Cresswell Limestone member now includes all the sequence of rocks that in some earlier reports was divided into the Cresswell Limestone member and the Luta Limestone member. The Cresswell Limestone member as now defined consists of a lower massive bed of yellow limestone, containing numerous echinoid spines and other fossils and characterized by pinkish-brown splotches on weathered surfaces, and an upper thin-bedded to concretionary sequence of limestone and shale (Pl. 4B). The thickness of the Cresswell Limestone member in Clay County is about 13 feet.

The Winfield Limestone yields water in quantities sufficient for most domestic and stock supplies, and where the formation is not deeply buried, the water is of good quality.

#### *Odell Shale*

The Odell Shale in Clay County consists of about 23 feet of shale. The lower few feet of the shale is principally yellow, and the middle and upper parts of it are mostly red and more calcareous than the lower part.

The Odell Shale does not yield water to wells in Clay County.

#### *Nolans Limestone*

The Nolans Limestone includes, in ascending order, the Krider Limestone, Paddock Shale, and Herington Limestone members. The thickness of the formation in Clay County ranges from 18 to 23 feet.

The Krider Limestone member consists of one bed of limestone or of two beds separated by a thin shale bed. The total thickness of the Krider Limestone member is not known to exceed 2 feet. In most places the limestone portion of the member is yellow sandy-appearing nodular limestone containing many fossil brachiopods and pelecypods.

The Paddock Shale member consists of 11 to 13 feet of gray fossiliferous shale.

The Herington Limestone member is the most conspicuous part of the Nolans Limestone. Although it is not as resistant to weathering as the Cresswell Limestone member of the Winfield Limestone, the Herington does form an escarpment in favorable topographic situations. In most places the member consists of 6 to 8 feet of yellow to brown, pelecypod-bearing limestone. The upper part is almost everywhere more massive than the lower part of the member (Pl. 5A).

The Nolans Limestone is not a good aquifer in Clay County. The Krider Limestone and Paddock Shale members are not known to yield water. Under favorable conditions the Herington Limestone member might yield some water to domestic and stock wells.

### Leonardian Series—Sumner Group

#### *Wellington Formation*

The Wellington Formation is not well exposed in Clay County, hence can be described only from observations in other areas and from the interpretation of well logs.

The lower part of the Wellington Formation in Clay County consists chiefly of greenish-gray and red clay shale. Locally, a massive

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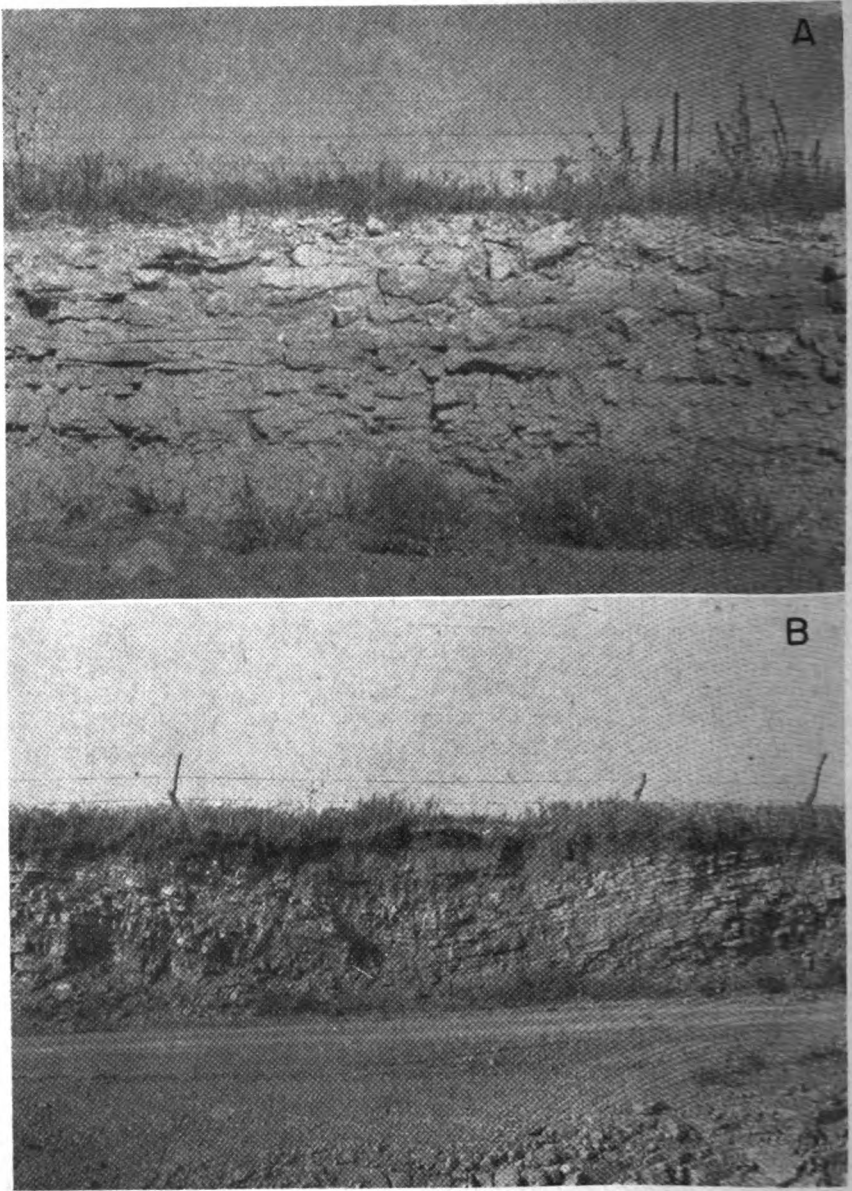


PLATE 5. A, Herington Limestone member exposed in road cut, NW $\frac{1}{4}$  sec. 9, T. 9 S., R. 3 E.; B, Hollenberg Limestone member in NW $\frac{1}{4}$  sec. 32, T. 6 S., R. 4 E.

bed of gypsum forms the base of the formation in contact with the Herington Limestone member of the Nolans Limestone. This massive bed of gypsum is well exposed in the southwest bank of a creek about 100 feet east and 300 feet south of the center of the west line of sec. 19, T. 6 S., R. 4 E.

A bed of impure dolomitic limestone, the Hollenberg Limestone member, lies about 40 feet above the base of the Wellington Formation in the northern part of the county. In the area southwest of Broughton the sequence between the Herington Limestone member and the Hollenberg Limestone member may be as thin as 20 feet. The Hollenberg is typically a succession of light-yellow earthy limestone and shale beds, each of which is about 3 to 5 inches thick. In almost every observed exposure in Clay County, the Hollenberg is deformed by many small anticlines and synclines (Pl. 5B) having an amplitude of only a few feet and a distance between crests of about 20 feet. Small-scale faulting also is very common in exposures of the Hollenberg (Pl. 6A). The thickness of the Hollenberg ranges from slightly less than 1 foot to 5 feet.

That part of the Wellington Formation that overlies the Hollenberg Limestone member in Clay County is in general darker than that below. At several widely scattered points a thin discontinuous bed of dark-gray to black chert was found about 12 to 15 feet above the Hollenberg Limestone member. The Hutchinson Salt member, which occurs in the middle part of the formation farther southwest, does not underlie any of Clay County. The greatest thickness of Wellington Formation underlying Clay County is not known, but probably is about 250 feet.

The Wellington Formation yields only very small quantities of water of poor quality to wells in Clay County. One spring (7-3-4da) discharges about 200 gpm of mineralized water from the Wellington Formation.

## CRETACEOUS SYSTEM

### Comanchean Series

#### *Kiowa Shale*

The base of the Cretaceous rocks is marked by a major unconformity indicated by a cobble zone. The cobbles consist of quartzite, chert, igneous rock, and many types of metamorphic rocks. At many places in Clay County the Dakota Formation rests on the Wellington Formation, the cobble zone being incorporated in the base of the Dakota Formation, and the Kiowa Shale is absent. At a few localities in southern Clay County a dark clay shale resem-

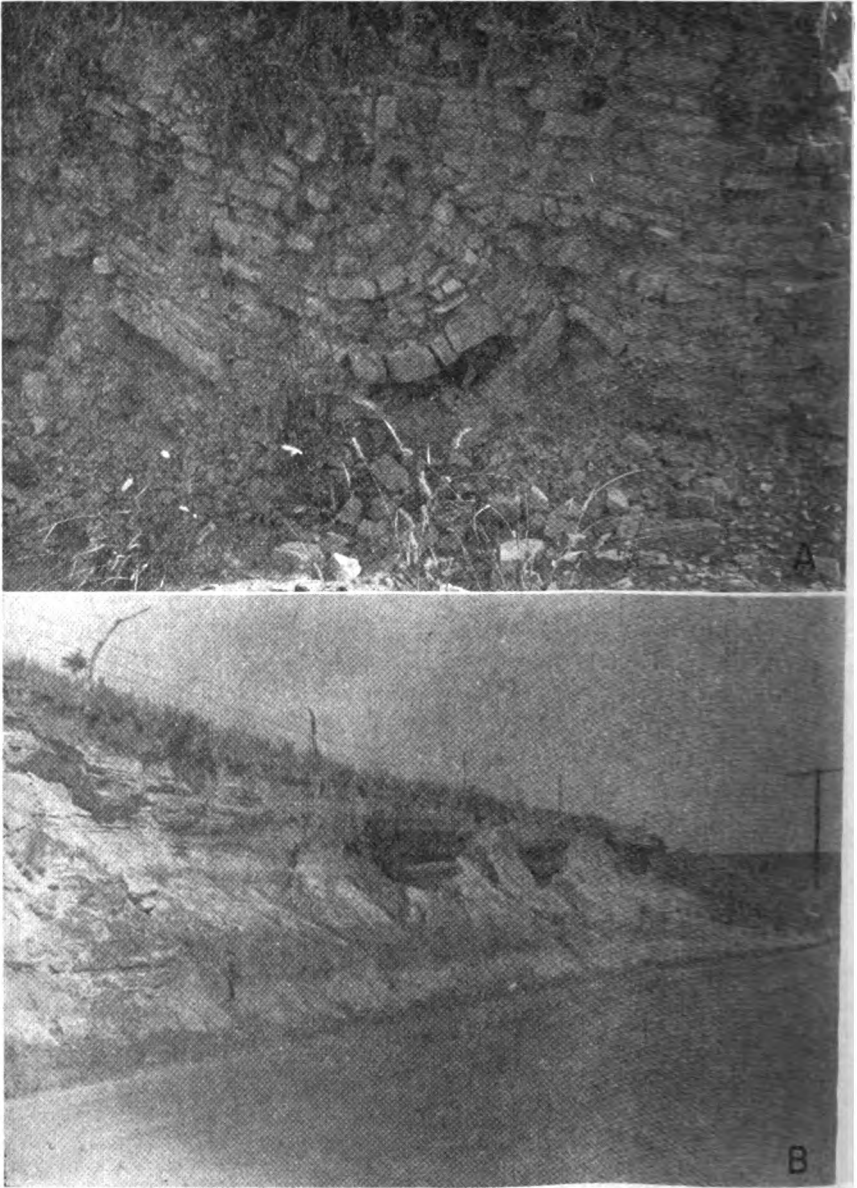


PLATE 6. A, Faulting in Hollenberg Limestone member in NW $\frac{1}{4}$  sec. 32, T. 6 S., R. 4 E.; B, Concretions in Dakota Formation, NE $\frac{1}{4}$  sec. 17, T. 10 S., R. 1 E.



bling the Kiowa crops out at about the contact between the Permian and Cretaceous, but at these localities the cobble zone is not noticeably present at the base of the shale, and the age of the shale is not known. Because the Kiowa Shale was not definitely recognized in Clay County and because of the similarity of their water-bearing characteristics to those of the Wellington Formation, these deposits are included with the Wellington Formation, and the Kiowa Shale is not shown on the geologic map.

Gulfian Series—Colorado Group  
*Dakota Formation*

The Dakota Formation in Kansas consists of two members, the lower Terra Cotta Clay member and the upper Janssen Clay member. The Dakota Formation popularly is believed to consist chiefly of sandstone, but as the member names indicate, clay is the dominant material in most of Kansas. The formation consists entirely of continental or near-shore deposits. Study of surface exposures and test-hole cuttings shows that the formation consists principally of clay of various colors but contains discontinuous beds and lenses of sandstone and siltstone. Minor amounts of pyrite, limonite, hematite, and siderite as concretions (Pl. 6B), thin beds of lignite, and small lenses of quartzitic sandstone are scattered throughout the formation.

In the subsurface most of the sandstone is light gray, but upon being exposed to weathering it becomes yellow, dark brown, or almost red. Some of the sandstone beds are very resistant to weathering and form bold outcrops.

The Dakota Formation is the youngest Cretaceous formation exposed in Clay County, and where it occurs its top is an erosional surface. The thickness of the Dakota Formation in Clay County ranges from a featheredge to about 165 feet.

The Dakota Formation is one of the principal aquifers in Clay County. The water is generally of good quality. Few wells in Clay County that obtain water from the Dakota Formation yield more than 30 gpm, but it is likely that test drilling in some areas might locate a sufficient thickness of sandstone to yield as much as 100 gpm. Water in the Dakota Formation is under artesian pressure in most parts of Clay County. No flowing wells were observed in the county, but a well flowing from the Dakota Formation is lo-

cated a few feet west of the county line in Cloud County. Chemical analyses of water from the Dakota Formation are shown in Table 3 and in Figure 8.

#### QUATERNARY SYSTEM

##### Pleistocene Series

Pleistocene deposits of Nebraskan age possibly are present in Clay County but are unimportant quantitatively. During the first stage of the Pleistocene Epoch there were probably no well-established lines of through drainage, and erosion was the dominant process. Remnants of conglomerate (Pl. 7A), which are exposed in several places in northeastern and south-central Clay County, probably were deposited during Nebraskan or early Kansan time. The conglomerate contains pebbles and fragments of sandstone and limestone derived from Cretaceous rocks and is overlain by loess of Illinoian and Wisconsinan age. These deposits are not known to exceed 4 feet in thickness, and because they cover a very small area they are not shown on the geologic map.

##### *Kansan Stage*

A few feet of locally derived gravel of probable Kansan age underlies the alluvium of Republican River upstream from Clay Center. Although Republican River did not flow through Clay County during Kansan time, a stream whose headwaters were in the drainage area of what is now White Rock Creek in Jewell County probably flowed where Republican River now has its course through Clay County. In most places this deposit of limestone and sandstone gravel (Grand Island Formation) is 6 to 10 feet thick and in some places is overlain by a few feet of clay of the Sappa Formation.

The water in the basal gravels has about the same quality as that in the overlying Recent alluvium. The Kansan gravel deposits are made up of more angular particles than other valley-fill deposits in the county and probably do not yield as much water as an equal thickness of younger gravel. They do add considerably to the total yield of wells in northwestern Clay County, however.

##### *Illinoian Stage*

*Terrace deposits.*—Deposits of Illinoian age in Clay County are represented by the Crete Formation along Republican River and its major tributaries. The Crete Formation in Clay County is in the form of terrace deposits and consists chiefly of silt and clay contain-

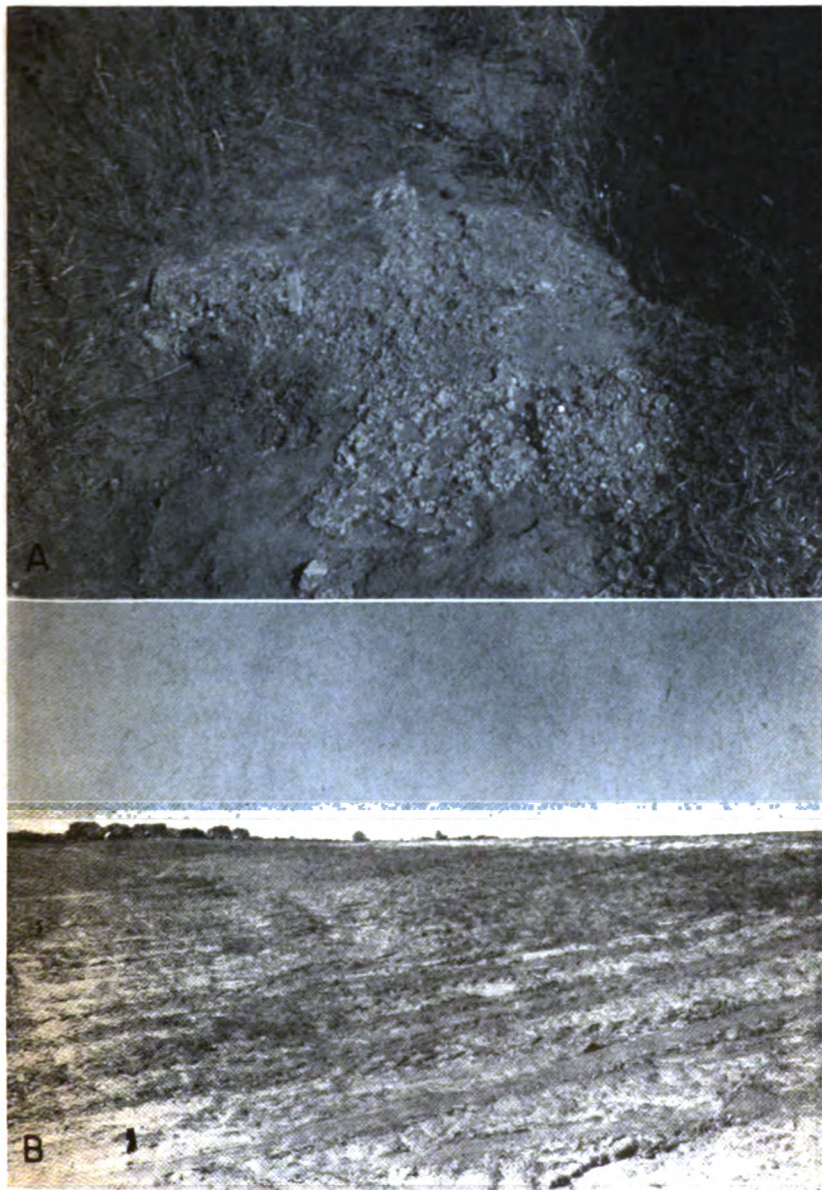


PLATE 7. A, Exposure of Pleistocene conglomerate, center south side sec. 10, T. 10 S., R. 2 E.; B, Prominent scarp formed by Wisconsin terrace, NE 1/4 sec. 2, T. 6 S., R. 1 E.

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ing minor amounts of sand and gravel. Nearly all the deposits of the Crete Formation in Clay County are blanketed by a thick covering of Peoria loess, and are not shown on the geologic map. The upper surface of the Crete Formation at most places in Clay County is about 30 to 40 feet above flood plain level.

The Crete Formation in Clay County is generally only partly saturated; however, a bed of locally derived gravel, which is present in many places at the base of the formation, is generally saturated and yields small to moderate supplies of water.

*Eolian deposits.*—Much of the upland area of Clay County as well as the Crete Formation is blanketed with wind-deposited silt commonly known as loess. All the loess in Clay County is included in the Sanborn Group, but only the oldest silt formation, the Loveland Formation, is of Illinoian age. In fresh exposures the Loveland consists of reddish-yellow silt and clay. Most deposits in Clay County are only a few feet thick, and at some localities they are entirely contained within the Sangamon soil, which was developed on the Loveland during the interglacial period between the Illinoian and Wisconsinan glaciers. The Loveland Formation is above the water table and does not yield water to wells.

#### *Wisconsinan Stage*

*Terrace deposits.*—Deposits of sand, gravel, and clay locally almost 70 feet thick were laid down along Republican River and other major streams during Wisconsinan time. The streams have subsequently cut to a lower level and these alluvial deposits now form low terraces above the flood plain. These terraces were mapped solely on the basis of topographic expression and probably include both early and late Wisconsinan deposits (Pl. 1).

The Wisconsinan terrace deposits along Chapman Creek and the streams tributary to Republican River are composed chiefly of silt and clay but contain minor amounts of locally derived sand and gravel in the lower part. These deposits yield small to moderate quantities of water to wells, adequate only for domestic or stock use.

The Wisconsinan terrace deposits of Republican River in Clay County are composed almost entirely of arkosic sand and gravel, with the exception of the upper 15 to 20 feet, which in most places is silt and clay. No evidence of a continuous clay bed at the base of the Wisconsinan terrace deposits such as occurs near Concordia, in Cloud County, was found in Clay County. The Wisconsinan

terrace deposits lack the hummocky topography typical of the Recent alluvium, and the upper surface of the terrace deposits lies about 12 feet above the upper surface of the alluvium (Pl. 7B). In the area northwest of Clay Center, the Wisconsinan terrace deposits along Republican River yield large quantities of water to wells. Downstream from Clay Center, where the valley is cut through more resistant rocks, the terrace deposits are thinner and more restricted laterally, and the available supply of ground water is somewhat less than in the valley upstream from Clay Center. Water from the Wisconsinan terrace deposits in Clay County is hard but otherwise satisfactory for most uses.

*Eolian deposits.*—Large areas of Clay County are mantled by deposits of loess, or wind-deposited silt, of Wisconsinan age. This silt of the Peoria Formation is contemporaneous with or older than the Wisconsinan terrace deposits. The silt of the Peoria Formation is gray buff and in most places is separated from the underlying reddish-yellow silt of the Loveland Formation by the Sangamon soil. The silt of the Peoria Formation, like all silts, has low permeability and retards the downward movement of rain water, and thus hinders recharge in the upland areas of Clay County. The Peoria Formation lies above the water table and does not yield water to wells.

#### *Recent Stage*

*Alluvium.*—Most of the streams of Clay County traverse flood plains underlain by alluvium of Recent age. The alluvium consists predominantly of sand and gravel but includes some silt and clay. The upper surface of the alluvium is characteristically very hummocky and in most places lies about 12 feet lower than the upper surface of the Wisconsinan terraces. Much of the alluvium is covered by water during major floods; it is subjected to erosion in some places, but in other places a small amount of additional material is added. The composition of the alluvium is nearly the same as that of the Wisconsinan terrace deposits, but as the alluvium is somewhat thicker and generally has a greater percentage of saturation, somewhat larger yields of water are obtained from wells in the alluvium.

The quality of the water in the alluvium of Republican River in Clay County is about the same as that of the water in the Wisconsinan terrace deposits. The water is hard but not strongly mineralized.

*Dune sand.*—In an area west of Vining, an area north of Wakefield, and in several smaller areas along Republican River, the wind has piled up sand from the channel and flood plain to form dunes. These dunes are still active in spots, and very little soil has developed on them. The dune areas are composed of many sand hills or dunes interspersed with undrained depressions. The dunes are above the water table and do not yield water to wells, but because of their permeability and lack of drainage they form important recharge facilities for the underlying alluvium and terrace deposits.

**RECORDS OF WELLS**

Records of 143 wells, 28 test holes, and 110 auger holes in Clay County are given in Table 10. All information classed as reported was obtained from the owner or tenant. Reported depths of wells are given in feet; measured depths are in feet and tenths. Reported depths to water level are given in feet; measured depths to water level are given in feet, tenths, and hundredths. The well-numbering system used in this table is explained on page 12.

TABLE 10.—Records of wells, springs, and test holes in Clay County

Well number	Location	Owner or tenant	Type of well (1)	Depth of well, feet, (2)	Diameter of casing, inches (3)	Principal water-bearing bed		Method of lift (4)	Use of water (5)	Measuring point			Depth to water below land surface, feet, (6)	Date of measurement	REMARKS (Yield given in gallons a minute; draw-down in feet)	
						Character of material	Geologic source			Description	Distance above land surface, feet	Height above mean sea level, feet				
6-1-1ba	T. 6 S., R. 1 E. NE NW sec. 1.	U. S. Geol. Survey	A	29.0	4	N	Sand	Terrace deposits	N	N	Land surface	0.0	1266.6	22.70	4-16-54	Did not reach bed-rock.
6-1-1bc	SW NW sec. 1.	Northern Natural Gas Co.	D	50	16	S	Sand and gravel	do	T, E	Ind.	do	0.0	1264.0	22	10-12-55	Four wells; one symbol on Pl. 2
6-1-2ac	SW NE sec. 2.	F. Turner	D	70.0	18	GI	do	do	T, T	I	Hole in pump base	0.2	1265.4	22.01	4-16-54	Aquifer test run 10-12-55.
6-1-2baa	NE NW sec. 2.	U. S. Geol. Survey	A	29.0	4	N	do	do	N	N	Land surface	0.0	1269.1	25.00	4-16-54	Did not reach bed-rock.
*6-1-2bae1	NE NW sec. 2.	City of Clifton	D	65.0	12	S	do	do	T, E	P	do	0.0	1268.4	21	4-5-49	Drillers log of near-by test hole is given.
6-1-2bae2	NE NW sec. 2	do	D	66	12	S	do	do	T, E	P	do	0.0	1267.9	21		Test hole drilled by Layne Western.
6-1-2bae3	NE NW sec. 2.	do	D	62	4	N	do	do	N	N	do	0.0	1272.0		6-17-54	Test hole.
6-1-2bb	NW NW sec. 2	U. S. Geol. Survey	D	70	4	N	do	do	T, T	I	Hole in pump base	2.0	1261.6	10.17	10-25-55	Drilled in 1955.
6-1-2bc	SW NW sec. 2.	C. Nelson	D	56.5	18	GI	do	do								Not used that season.
*6-1-2cc	SW SW sec. 2.	U. S. Geol. Survey	D	70	4	N	do	Alluvium	N, T	N	Land surface	0.0	1256.8	7.10	6-19-54	Test hole.
6-1-2dbs	NW SE sec. 2.	C. Anderson	D	60.4	18	GI	do	Terrace deposits	T, T	I	Hole in pump base	0.2	1264.5	23.07	10-25-55	Test hole.
6-1-3aa	NE NE sec. 3.	U. S. Geol. Survey	A	29.0	4	N	Sand	do	N	N	Land surface	0.0	1269.9	24.50	4-16-54	Did not reach bedrock.
6-1-3abb	NW NE sec. 3	do	A	29.0	4	N	do	do	N	N	do	0.0	1270.8	18.00	4-16-54	do
6-1-3ad	SE NE sec. 3	do	A	14.0	4	N	do	do	N	N	do	0.0	1269.9	10.70	4-16-54	do
6-1-3add	SE SE sec. 3	do	A	9.0	4	N	do	Alluvium	N	N	do	0.0	1256.8	7.20	4-16-54	do
6-1-3aaa	NE NE sec. 4.	do	A	34.0	4	N	Sand and gravel	Terrace deposits	N	N	do	0.0	1272.1	19.20	4-15-54	do
6-1-4abb	NW NE sec. 4.	do	A	24.0	4	N	Sand	do	N	N	do	0.0	1271.8	20.70	4-15-54	do
6-1-4bbb	NW NW sec. 4.	do	A	24.0	4	N	do	do	N	N	do	0.0	1272.7	19.10	4-15-54	do





TABLE 10.—Records of wells, springs, and test holes in Clay County—Continued

Well number	Location	Owner or tenant	Type of well (1)	Depth of well, feet (2)	Diameter of well, inches (3)	Principal water-bearing bed		Method of lift (4)	Use of water (5)	Measuring point			Depth to water level below land surface, feet (6)	Date of measurement	REMARKS (Yield given in gallons a minute; drawdown in feet)
						Character of material	Geologic source			Description	Distance above land surface, feet	Height above mean sea level, feet			
6-2-28cd	SE SW sec. 28. . . .	U. S. Geol. Survey	A	19 0	4	N	Sand . . . . .	Terrace deposits. . . . .	N	N	Land surface. . . . .	0 0	1246.1	4-20-54	Did not reach bedrock.
6-2-29ac	SW NE sec. 29. . . .	Millies . . . . .	D	40 5	18	GI	do. . . . .	Alluvium. . . . .	Ce, T	I	Top of casing	0 3	1242.8	10-25-55	Yields about 1000 gpm.
6-2-29dad	NE SE sec. 29. . . .	do. . . . .	D	53 5	18	GI	do. . . . .	do. . . . .	Ce, T	I	do. . . . .	0 3	1242.4	10-25-55	do
6-2-29dad	SE SE sec. 29. . . .	U. S. Geol. Survey	A	9 0	4	N	do. . . . .	do. . . . .	N	N	Land surface. . . . .	0 0	1239.9	4-20-54	Did not reach bedrock.
6-2-30ccc	SW SW sec. 30. . . .	do. . . . .	A	19 0	4	N	do. . . . .	Terrace deposits. . . . .	N	N	do. . . . .	0 0	1244.8	4-19-54	do
6-2-30dc	SW SE sec. 30. . . .	do. . . . .	A	24 0	4	N	do. . . . .	do. . . . .	N	N	do. . . . .	0 0	1243.5	4-19-54	do
6-2-31cd	SE SW sec. 31. . . .	do. . . . .	D	48	4	N	do. . . . .	do. . . . .	N	N	do. . . . .	0 0	1253.8	6-25-54	Test hole.
6-2-31dc	SW SE sec. 31. . . .	do. . . . .	D	70	4	N	do. . . . .	do. . . . .	N	N	do. . . . .	0 0	1232.9	6-22-54	do
6-2-32ab	NW NE sec. 32. . . .	do. . . . .	A	14 0	4	N	do. . . . .	Alluvium. . . . .	N	N	do. . . . .	0 0	1239.3	4-20-54	Did not reach bedrock.
6-2-32bbb	NW NW sec. 32. . . .	do. . . . .	A	24 0	4	N	do. . . . .	Terrace deposits. . . . .	N	N	do. . . . .	0 0	1241.5	4-19-54	do
6-2-32ccc	SW SW sec. 32. . . .	do. . . . .	D	80	4	N	do. . . . .	Alluvium. . . . .	N	N	do. . . . .	0 0	1225.6	6-25-54	Test hole.
6-2-32cd	SE SW sec. 32. . . .	do. . . . .	D	70	4	N	do. . . . .	do. . . . .	N	N	do. . . . .	0 0	1233.7	6-23-54	do
6-2-32dd	SE SE sec. 32. . . .	do. . . . .	D	70	4	N	do. . . . .	do. . . . .	N	N	do. . . . .	0 0	1225.2	6-28-54	do
6-2-33cd	SE SW sec. 33. . . .	do. . . . .	D	80	4	N	do. . . . .	Terrace deposits. . . . .	N	N	do. . . . .	0 0	1236.9	6-24-54	do
6-2-33dcb	SW SE sec. 33. . . .	Bill Silver . . . .	D	69 6	18	GI	Sand and gravel	do. . . . .	T, NG	I	Hole in pump base	1 5	1241.2	10-25-55	Yields about 1000 gpm.
6-2-34abb	NW NE sec. 34. . . .	U. S. Geol. Survey	A	44 0	4	N	Clay . . . . .	Colluvium (?) . . . . .	N	N	Land surface. . . . .	0 0	1306.9	4-20-54	May have reached Dakota Formation.
6-2-34bbb	NW NW sec. 34. . . .	do. . . . .	A	34 0	4	N	Sand . . . . .	Terrace deposits. . . . .	N	N	do. . . . .	0 0	1258.1	4-20-54	Did not reach bedrock.
6-2-34cd	SE SW sec. 34. . . .	do. . . . .	D	46	4	N	do. . . . .	do. . . . .	N	N	do. . . . .	0 0	1239.2	6-28-54	Test hole.
6-3-2dd	T. 6 S., R. 3 E.	do. . . . .	D	46 7	6	GI	Limestone	Nolans Limestone	Cy, W	S	Top of casing	1 1	1311.9	8-19-54	do
6-3-3cd	SE SE sec. 2. . . . .	M. Sanneman. . . . .	D	41 4	6	GI	Sandstone	Dakota Formation	Cy, W	S	Base of pump	1 5	1342.0	10-26-54	do
6-3-15-b	NW SW sec. 15. . . .	F. J. Doberer. . . . .	Du	10 4	40	R	do. . . . .	do. . . . .	Cy, W	S	Top of platform	0 0	1321.2	8-19-54	do

6-3-23cd	SW SE sec. 23	Alvonia Francon.	D	67 7	6	GI	Limestone	Nolans Limestone	Cy. W	N	Base of pump	0 2	1302 6	85 86	8 10 51	Abandoned stuck well.
6-4-4bb	T. S. R. & E. NW NW sec. 4	T. R. Cynamina	Du	35 5	48	R	Limestone	Nolans Limestone	Cy. W	N	Base of pump	1 3		28 77	8 16 54	Abandoned stuck well.
6-4-6bb	NW NW sec. 9	A. Boughren	Du	45 2	40	R	do	do	Cy. W	N	do	0 4		21 81	6 16 54	Abandoned stuck well.
6-4-18ba	NE NW sec. 18	R. Johnson	D	76 8	6	GI	do	Nolans and Winsfield Limestones	Cy. W	S	do	0 5	1314 0	54 10	8-17 54	
6-4-19cb	NW SW sec. 19	U. S. Geol. Survey	D	52	4	N										Drilled to measure 8 1/2 ft. lead.
6-4-31da	NE SE sec. 31	F. M. Carrithers	D	63	6	GI	Limestone	Nolans Limestone	Cy. W	S	do	0 2		34 15	8 17 54	
6-4-33ad	SE NE sec. 33	J. Yeager	D	103 2	6	GI	do	Barrenstone Limestone	Cy. W, H	S	do	0 6		63 60	8 16 54	
7-1-2bh	T. S. R. & E. NW NW sec. 2	School District	D	60 7	8	S	Sand	Terrace deposits	Cy. H	D	Base of pump	0 4	1283 3	21 15	9-3 54	
7-1-6bb	NW NW sec. 6	U. S. Geol. Survey	D	133	4	N	Sandstone	Dakota Formation	Cy. H	N	Land surface	0 0	1300 0			
7-1-8bb	NW NW sec. 8	School District	D	55 2	8	GI	do	do	Cy. H	D	Base of pump	0 4	1354 4	30 00	9 2 51	
7-1-13bb	NW NW sec. 13	A. Hansen	Du, D	43 0	5-36	GI, R	do	do	Cy. W	S	do	0 5	1350 4	45 0	9 30 51	Originally a dug well has been deepened by drilling.
7-1-15bb	SW NW sec. 15	M. M. Couray	Du, D	65 0	6-36	GI, R	do	do	Cy. W, H	D, S	do	0 8	1330 4	18 72	6 9 54	do
7-1-15lda	SE SE sec. 15	School District	D	31 3	6	GI	do	do	Cy. H	D	do	0 1	1315 9	21 90	9-3 54	
7-1-19be	SW NW sec. 19	D. B. Molen	Du	50	48	R	do	do	Cy. W	D, S	Land surface	0 0	1298 5	47	9-3 54	
7-1-20be	SW NW sec. 20	H. Lupton	D	41 9	6	GI	do	do	Cy. H	S	do	0 6	1372 9	30 98	9 2 51	
7-1-31bb	SE NW sec. 31	C. W. Crawford	Du	34 8	48	R	do	do	Cy. W	S	do	0 8	1341 0	26 52	9-3 51	
7-1-36bb	NW NW sec. 36	A. Daensthaer	D	51 0	6	GI	do	do	Cy. W	D, S	do	0 4	1313 8	30 88	9-9 51	
7-2-3bb	T. S. R. & E. SW NW sec. 3	U. S. Geol. Survey	D	80	4	N	Sand and gravel	Terrace deposits	N	N	Land surface	0 0	1247 9	23 10	6 28 51	Test hole.
7-2-4de	SW SE sec. 3	City of Moranville	D	53	12	S	Gravel	do	T, R	P	do	0 0	1228 3	22		Near water tower
7-2-4dde	SE SE sec. 3	do	D	57	12	S	do	do	T, E	P	do	0 0	1231 4	20		Four blocks west of water tower.
7-2-6bb	SW NW sec. 6	U. S. Geol. Survey	D	110	4	N	Sandstone	Dakota Formation	N	N	do	0 0	1315 0	82 20	10 22 51	Test hole.
7-2-7ab	SW NE sec. 7	do	A	11 0	4	N	Sand	Alluvium	N	N	do	0 0	1224 4	12 10	4 20 51	Did not reach bedrock.
7-2-7ad	SE NE sec. 7	do	A	14 0	4	N	do	do	N	N	do	0 0	1221 5	11 50	5 26 51	do
7-2-7c	SW SE sec. 7	do	A	45 0	4	N	do	Terrace deposits	N	N	do	0 0	1245 1	42 50	5 26 51	do
7-2-8a-1	SW NE sec. 8	do	A	11 0	4	N	do	Alluvium	N	N	do	0 0	1218 8	12 10	4 20 51	do
7-2-8a-2	SW NE sec. 8	do	A	19 0	4	N	do	do	N	N	do	0 0	1217 5	9 80	4 20 51	do
7-2-8a	SE SW sec. 8	do	A	11 0	4	N	do	do	N	N	do	0 0	1219 2	10 10	4 20 51	do
7-2-9ab	SE SW sec. 9	do	A	3 0	4	N	do	do	N	N	do	0 0	1232 1	18 90	4 20 51	do
7-2-9bd	SE NW sec. 9	do	A	21 0	4	N	do	do	N	N	do	0 0	1224 5	17 20	4 20 51	do
7-2-10ab	SE NW sec. 10	do	A	31 0	4	N	do	do	N	N	do	0 0	1230 1	22 40	4 20 51	do
7-2-10bb	NW NW sec. 10	do	A	24 0	4	N	do	do	N	N	do	0 0	1225 3	17 00	4 20 51	do

TABLE 10.—Records of wells, springs, and test holes in Clay County—Continued

Well number	Location	Owner or tenant	Type of well (1)	Depth of well, feet (2)	Diameter of well, inches	Principal water-bearing bed		Method of lift (4)	Use of water (5)	Measuring point			Depth to water level below land surface, feet, (6)	Date of measurement	REMARKS (Yield given in gallons a minute; drawdown in feet)	
						Character of material	Geologic source			Description	Distance above land surface, feet	Height above mean sea level, feet				
7-2-11abb.	NW NE sec. 11.	U. S. Geol. Survey	A	29.0	4	N	Sand	Terrace deposits	N	N	Land surface	0.0	1236.6	12.30	4-20-54	Did not reach bedrock.
7-2-14cd	SE SW sec. 14	do	A	29.0	4	N	do	do	N	N	do	0.0	1219.3	19.50	5-25-54	do
7-2-15dd	SE SE sec. 15	do	A	19.0	4	N	do	do	N	N	do	0.0	1220.2	18.50	5-25-54	do
7-2-15ed	SE SW sec. 15	do	A	23.0	4	N	do	do	N	N	do	0.0	1221.2	22.70	4-21-54	do
7-2-16fd	SE SW sec. 16	do	A	14.0	4	N	do	do	N	N	do	0.0	1212.0	12.00	4-21-54	do
7-2-16fd	SE SE sec. 16	do	A	39.0	4	N	do	do	N	N	do	0.0	1227.9	23.50	4-21-54	do
7-2-18aa	NE NE sec. 18	do	A	30.0	4	N	do	do	N	N	do	0.0	1230.7	22.00	5-26-54	do
7-2-18dd	SE SE sec. 18	do	A	34.0	4	N	do	do	N	N	do	0.0	1233.3	25.00	5-26-54	do
7-2-19bb	NW NW sec. 19	E. Affolter	D	44.1	5	GI	Sandstone	Dakota Formation	N	N	Top of casing	1.1	1292.4	31.90	9-30-54	Abandoned stock well.
7-2-21ac	SW NE sec. 21	E. H. Eickerman	D	53.0	18	GI	Sand and gravel	Terrace deposits	N	N	do	0.3	1212.0	10.95	10-25-55	Unused irrigation well.
7-2-21bb	NW NW sec. 21	U. S. Geol. Survey	A	24.0	4	N	Sand	do	N	N	Land surface	0.0	1218.3	14.70	4-21-54	Did not reach bedrock.
7-2-23aa	NE NE sec. 23	do	A	34.0	4	N	do	do	N	N	do	0.0	1227.0	30.30	5-26-54	do
7-2-29dac	NE SE sec. 29	Oecil Cole	D	41.0	6	GI	Sand and gravel	do	Ce, T	I	Top of casing	1.2	1206.4	10.50	10-29-55	A four-wall battery.
7-2-34aa	NE NE sec. 34	U. S. Geol. Survey	A	34.0	4	N	Sand	do	N	N	Land surface	0.0	1214.9	22.20	5-26-54	Did not reach bedrock.
7-2-34ba1	NE NW sec. 34	do	A	28.0	4	N	do	do	N	N	do	0.0	1212.4	19.20	5-26-54	do
7-2-34ba2	NE NW sec. 34	do	A	29.0	4	N	do	do	N	N	do	0.0	1213.7	22.00	5-26-54	do
7-2-35aa	NE NE sec. 35	do	A	29.0	4	N	do	do	N	N	do	0.0	1212.7	23.50	5-26-54	do
7-2-35ab	NW NE sec. 35	do	A	29.0	4	N	do	do	N	N	do	0.0	1213.2	23.50	5-26-54	do
7-2-36aa	NE NE sec. 36	do	A	34.0	4	N	do	do	N	N	do	0.0	1217.5	31.00	5-26-54	do
7-2-36ba	NE NW sec. 36	do	A	24.0	4	N	do	do	N	N	do	0.0	1209.5	22.00	5-26-54	do
7-3-1cc	T. 7 S. R. 9 E. 9W 5W sec. 1	School District	Du	9.5	48	R	Sandstone	Dakota Formation	N	N	Top of platform	1.8	1396.5	6.02	8-19-54	Abandoned.

Well ID	Section	Surveyor	Sp	Shale	Formation	Notes	Depth	Flow	Remarks
7-3-4da	NE SE sec. 4	G. W. Root	Sp	Shale	Wellington Formation		1250 0	19 20	Estimated flow of 200 gpm.
7-3-5bb	NW NW sec. 5	U. S. Geol. Survey	D	N Sandstone	Dakota Formation	N	1333 7	10 22 54	Test hole
7-3-6dd	SE SE sec. 6	G. Hongsmerover	Du	R Shale	Wellington Formation	Cy, H	1321 5	10 26 54	Formerly domestic
7-3-6aa	NE NE sec. 9	W. H. Rundle	Du	R Shale	Wellington Formation	N	1291 7	8 19 54	do
7-3-19ad	NE NE sec. 19	C. Engquist	D	GI Limestone	Nolans Limestone	Cy, W	1293 3	10 26 54	
7-3-22cc	SW SW sec. 22	M. Smith	D	GI do	Wellington Formation	Cy, W	35 72	8-14 54	
7-3-26cc	SW SW sec. 26	School District	D	GI do	Nolans Limestone	Cy, H	53 63	8-19-54	School is abandoned.
7-3-26cd	SE SW sec. 29	Nicholas Schlitz	D	GI do	Wellington Fm. and Nolans Lk. Terrace deposits	Cy, W	1296 2	10 26 54	
7-3-31dde	SE SE sec. 31	T. Gerrits	D	S Sand and gravel	Terrace deposits	T, T	1205 2	10 28 55	
7-4-10ab	T. 7 S., R. 4 E.	F. F. Harner	D	GI Limestone	Winfield Limestone	Cy, H	77 40	8 16 54	
7-4-18aa	NW NE sec. 10	Geo. Masterson	D	GI do	Nolans Limestone	Cy, W	85 98	8 17 54	
7-4-20ad	SE NE sec. 20	City of Green	D	S do	Barnston Limestone	T, E			Estimated yield 15 gpm.
7-4-21bc	SW NW sec. 21	do	D	S do	do	T, E			Estimated yield 25 gpm.
7-4-32cc	SW SW sec. 32	J. Heilman	D	GI do	do	Cy, W	112 30	8-17 54	
8-1-4ddb	T. 8 S., R. 1 E.	P. F. Johnson	D	GI Sandstone	Dakota Formation	N	1293 5	14 51	Abandoned stork well.
8-1-5bb	NW NW sec. 5	Geo. Blackwood	D	GI do	do	Cy, W	1367 8	9 2 54	
8-1-7dd	SE SE sec. 7	School District	D	GI do	do	Cy, H	30 41	9 2 54	
8-1-9dde	SE SE sec. 9	M. McNeil	D	GI do	do	Cy, W	32 48	9 3 54	
8-1-12aa	NE NE sec. 12	J. W. Hay	D	GI do	do	J, E	22 95	9 30 54	Well reached blue shale at 40 feet.
8-1-13cc	NE SW sec. 13	City of Idana	D	S Sand	Terrace deposits	J, E	1239 3	22	
8-1-26cc	SW SW sec. 26	School District	D	GI Sandstone	Dakota Formation	Cy, H	1257 2	9 3 54	School is abandoned.
8-1-30ad	SE NE sec. 30	do	D	GI do	do	N	1382 3	15 90	9-2 54
8-2-1ab	T. 8 S., R. 2 E.	G. Dittmar	D	S Sand and gravel	Terrace deposits	T, G	1205 2	23 70	10-26-54
8-2-1de	SW SE sec. 1	U. S. Geol. Survey	A	N Sand	do	N	1204 8	25 60	1500 gpm. Did not reach bedrock
8-2-1dd	SE SE sec. 1	do	A	N do	do	N	1201 8	26 40	8-27-51
8-2-2cc	SW SW sec. 2	G. Dittmar	D	S Sand and gravel	Alluvium	T, G	1193 3	11 90	9-30-54

TABLE 10.—Records of wells, springs, and test holes in Clay County—Continued

Well number	Location	Owner or tenant	Type of well (1)	Depth of well, feet (2)	Diameter of well, inches (3)	Principal water-bearing bed		Method of lift (4)	Use of water (5)	Measuring point			Depth to water level below land surface, feet (6)	REMARKS (Yield given in gallons a minute; drawdown in feet)
						Character of material	Geologic source			Description	Distance above land surface, feet	Height above mean sea level, feet		
8-2-41a	NE SE, sec. 4	J. B. Kamphaus	D	81.0	6	GI	Limestone	Winfield Limestone	Cy, W	S	Base of pump	1.2	1270.6	67.65
8-2-5aa	NE NE, sec. 5	Eric Swenson	D	76.3	6	GI	do.	do.	Cy, W	S	do.	0.3	1250.1	21.04
8-2-3cd	SE SW, sec. 9	F. N. Schlitz	D	60.9	6	GI	Sand	Terrace deposits	Cy, W	S	do.	0.5	1234.1	26.25
8-2-10aa1	NE NE, sec. 10	U. S. Geol. Survey	A	19.0	4	N	do.	Alluvium	N	N	Land surface	0.0	1198.2	16.10
8-2-10aa2	NE NE, sec. 10	do.	D	40	4	N	do.	do.	N	N	do.	0.0	1194.6	11.80
8-2-10dd	SE SE, sec. 10	do.	A	75.0	4	N	do.	Terrace deposits	N	N	do.	0.0	1216.1	40.00
8-2-11aa1	NE NE, sec. 11	do.	A	29.0	4	N	do.	do.	N	N	do.	0.0	1205.5	20.10
8-2-11aa2	NE NE, sec. 11	H. Walker	D	57.0	16	S	Sand and gravel	do.	T, G	I	Top of casing	1.0	1206.4	25
8-2-11ad	SE NE, sec. 11	do.	D	40.0	16	GI	do.	Alluvium	Ce, G	I	do.	1.0	1193.2	12
8-2-11da	NE SE, sec. 11	do.	D	42.0	16	GI	do.	do.	Ce, G	I	do.	1.0	1193.8	12
8-2-23dc	SW SE, sec. 23	J. D. Shepherd	Du	58.0	48	R	Limestone	Wellington Formation	Cy, W	S	Base of pump	0.3	1242.6	42.00
8-2-27ba	NE NW, sec. 27	Leona Sylvester	D	86.7	6	GI	do.	do.	Cy, W	S	do.	0.4	1303.5	61.60
8-2-30dd	SE SE, sec. 30	School District	D	66	6	GI	Sandstone	Dakota Formation	Cy, H	D	do.	0.4	1303.5	61.60
8-2-33dd	SE SE, sec. 33	Mathew Hill	D	95.4	6	GI	Limestone	Winfield Limestone	Cy, W	S	Base of pump	0.6	1291.1	63.60
8-3-4ab	T. 9 S., R. 5 E. NW NE, sec. 4	W. A. Piper	D	118.2	6	GI	Limestone	Winfield Limestone	Cy, W	S	Base of pump	0.5	.....	78.85
8-3-6bb1	NW NW, sec. 8	Swift and Co.	D	65	8	S	Gravel	Terrace deposits	T, E	Ind.	Land surface	.....	.....	25
8-3-6bb2	NW NW, sec. 8	City of Clay Center	D	55	18	S	do.	do.	T, E	P	Land surface	0.0	1183.9	22
8-3-6bb3	NW NW, sec. 8	do.	D	55	18	S	do.	do.	T, E	P	Land surface	0.0	1183.9	22

Well ID	Section	Owner	D	55	18	S	do	do	T. E.	P	do	0 0	1187 4	22	Notes
8-3-8hh4	NW NW sec. 8	do	D	55	18	S	do	do			do	0 0	1187 4	22	South Utility Park well. 14 ft. drawn down after 8 hrs at 500 gpm. Yields about 50 gpm.
8-3-8bc	SW NW sec. 8	Coca Cola Co.	D	56	7	S	do	do	J. E.	Ind.	Top of curb	0 0	1188 0	26	Yields about 50 gpm.
8-3-8hd	SE NW sec. 8	City of Clay Center	D	60	16	S	do	do	T. E.	P	Land surface	0 0	1197 0	29	Dexter Park well. Reported drawn down of 4 ft. after 8 hrs at 500 gpm.
8-3-8cc	SW SW sec. 8	U. S. Geol. Survey	A	29	4	N	do	do	N	N	do	0 0	1195 0	26 90	At center of 4th and Main St.
8-3-8db	NW SE sec. 8	City of Clay Center	D	55	16	S	Sand and gravel	do	T. E.	P	do	0 0		28	Atch and McBrainery St.
8-3-11dd	SE SE sec. 11	Ella Schlitz	D	124 1	6	GI	Limestone	Barnston Formation	Cy. W	S	Base of pump	0 3		112 90	8-19-54
8-3-16dc	SE SW sec. 16	Steinbach	D	22 2	6	GI	Sand and gravel	Alluvium	Ce. B	I	Top of casing west well	1 0	1161 9	8 50	10-28 55
8-3-17ha	NE NW sec. 17	U. S. Geol. Survey	D	44	4	N	do	Terrace deposits	N	N	Land surface	0 0	1179 3	9 10	6-30-54
8-3-17hb	NW NW sec. 17	do	A	14 0	4	N	Sand	Alluvium	N	N	do	0 0	1179 3	13 20	Test hole. Did not reach bedrock.
8-3-17cc	SW SW sec. 17	do	A	19 0	4	N	do	do	N	N	do	0 0	1179 2	13 00	do
8-3-19hd	SE SE sec. 19	do	A	34 0	4	N	Clay	Terrace deposits	N	N	do	0 0	1196 7	26 20	do
8-3-21bd	NW NW sec. 21	Marshall	D	42 3	12	S	Gravel	do	Ce. T	I	Top of casing	-3 6	1160 7	12 40	10-28-55
8-3-25aa	NE NE sec. 25	H. A. Moeburg	D	112 3	6	GI	Limestone	Barnston Limestone	Cy. W. H	S	Base of pump	0 3		92 72	8-19-54
8-3-26hd	SE NW sec. 26	U. S. Geol. Survey	A	35 0	4	N	Sand	Terrace deposits	N	N	Land surface	0 0	1185 3	29 80	5-29-54
8-3-27ca	NE SW sec. 27	do	A	14 0	4	N	do	Alluvium	N	N	do	0 0	1168 4	10 50	5-27-54
8-3-27da	SE SE sec. 27	do	A	19 0	4	N	do	do	N	N	do	0 0	1173 9	15 00	do
8-3-28ac	SW NE sec. 28	do	A	19 0	4	N	do	Terrace deposits	N	N	do	0 0	1162 4	15 50	5-27-54
8-3-28bc	NW NW sec. 28	do	A	19 0	4	N	do	do	N	N	do	0 0	1168 1	15 30	do
8-3-28hc	SE NW sec. 28	do	A	45 0	18	GI	Gravel	Alluvium	T. B	I	Top of casing	2 3	1161 9	8 00	10-28-55
8-3-28hd	NE SE sec. 28	Albert Parry	D	45 0	18	GI	do	do	T. B	I	do	1 0	1159 4	10 90	10-28-55
8-3-28da	SE SE sec. 28	Norquist	D	37 6	18	GI	do	do	T. B	I	do	1 0	1159 4	10 90	Yields about 500 gpm.
8-3-28da	SE SE sec. 28	U. S. Geol. Survey	A	14 0	4	N	Sand	do	N	N	Land surface	0 0	1159 2	13 20	Did not reach bedrock.
8-3-34cd	SE SW sec. 34	do	A	19 0	4	N	do	Terrace deposits	N	N	do	0 0	1168 1	18 40	do
8-3-35aac	NE NE sec. 35	Bob Bowers	D	33 9	18	GI	Sand and gravel	do	T. B	I	Hole in pump base	0 5	1168 1	12 60	10-28-55
8-3-35da1	NE SE sec. 35	U. S. Geol. Survey	A	19 0	4	N	Sand	Alluvium	N	N	Land surface	0 0	1164 7	10 10	Yields about 350 gpm.
8-3-35da2	NE SE sec. 35	do	D	37	4	N	do	do	N	N	do	0 0	1161 5	6 40	Did not reach bedrock.
8-3-35dd	SE SE sec. 35	do	D	30	4	N	do	do	N	N	do	0 0	1169 3	6 40	Test hole.

TABLE 10.—Records of wells, springs, and test holes in Clay County—Continued

Well number	Location	Owner or tenant	Type of well (1)	Depth of well, feet (2)	Diameter of well, inches (3)	Principal water-bearing bed		Method of lift (4)	Use of water (5)	Measuring point			Depth to water level below land surface, feet (6)	Date of measurement	REMARKS (Yield given in gallons per minute; drawdown in feet)
						Character of material	Geologic source			Description	Distance above land surface, feet	Height above mean sea level, feet			
8-3-36hb. 8-3-36bbe.	NW NW sec. 38. NW NW sec. 38.	U. S. Geol. Survey do.	D A	50 31.0	N N	Sand do.	Terrace deposits. do.	N N	N N	Land surface do.	0.0 0.0	1181.9 1182.5	22.00 26.20	7-12-54 5-27-54	Test hole Did not reach bedrock.
8-4-4dd.	T. & S. R. & E. SE SE sec. 4.	School District.	D	132.5	6	GI	Limestone	Cy, H	D	Base of pump	0.3		85.30	8-16-54	
8-4-15lc. 8-4-17cd. 8-4-20ab. 8-4-21ac.	SW SE sec. 15. SE SW sec. 17. NW NE sec. 20. SW NE sec. 21.	Nellie Senn. Baxter. Samuel Goodin. H. Mall.	D D D D	118.3 130 100 62	6 12 12 12	GI S S S	do. do. do. do.	Cy, W T, B T, G T, T	S I I I	do. Land surface Base of pump do.	0.1 0.0 0.1 0.1	1276.8 1217.4 1223.7	61.62 55.90	8-17-54 10-28-55	Yields 50 gpm. Yields 200 gpm. Floxed when drilled, yields 280 gpm.
8-4-21dc. 8-4-34cba.	SW SE sec. 21. NW SW sec. 34.	do. Robt. Lloyd.	D D	72.0 87.3	12 6	S GI	do. do.	T, T Cy, W	I S	do. do.	0.1 1.0	1220.4	14.90 63.42	10-28-55 10-26-54	Yields 125 gpm.
9-1-1dd. 9-1-4dd. 9-1-7dcd. 9-1-265b. 9-1-277ba. 9-1-32bb.	T. & S. R. & E. SE SE sec. 1. SW SE sec. 4. SW SE sec. 7. NW NW sec. 28. NW NW sec. 27. NW NW sec. 32.	Ernest Landin. O. Sterrett. Abigail Bull. School District. Mathew Hill. O. M. Panton.	D Du D Du Du Du	89.8 41.7 64.0 15.0 49.0 50.2	6 48 6 46 56 30	GI R GI R R R	Sandstone do. do. do. do. do.	Cy, W Cy, W Cy, W Cy, H Cy, H Cy, H	S S S D S S	Base of pump do. do. do. Top of curb. do.	0.8 1.6 1.0 0.1 0.8 0.5	1372.1 1418.2 1319.7 1302.1 1361.7 1366.0	62.40 19.77 19.32 11.00 30.51 42.23	9-30-54 9-3-54 9-2-54 9-3-54 9-3-54 9-2-54	
9-2-11cb. 9-2-16ab. 9-2-17bb. 9-2-24cc.	T. & S. R. & E. NW SW sec. 11. NW NE sec. 16. NW NW sec. 17. SW SW sec. 24.	School District. U. S. Geol. Survey Fed. Land Bank. Geo. Lenhart.	A A Du D	70.1 50.0 61.0 75.7	6 4 40 6	GI N R GI	Limestone Sandstone do. Wellington Formation	Cy, H Cy, W Cy, W Cy, W	D N S S	Base of pump Land surface Base of pump do.	0.2 0.0 1.3 0.3	1323.3	46.00 37.00 47.20 43.18	9-30-54 6-7-54 9-30-54 10-21-51	Wellington shale at 38 feet.





TABLE 10.—Records of wells, springs, and test holes in Clay County—Concluded

Well number	Location	Owner or tenant	Type of well (1)	Depth of well, feet (2)	Diameter of well, inches (3)	Principal water-bearing bed		Method of lift (4)	Use of water (5)	Measuring point			Depth to water level below land surface, feet (6)	Date of measurement	REMARKS (Yield given in gallons a minute; drawdown in feet)
						Character of material	Geologic source			Description	Dis-tance above land surface, feet	Height above mean sea level, feet			
10-1-1da	T. 10 S., R. 1 E., NE SE sec. 1	School District	D	47.4	6	Sandstone	Dakota Formation	Cy, H	N	Base of pump	0.4	1288.3	30.85	9-30-54	School abandoned.
10-1-3cc	SW SW sec. 3	do.	D	39.0	6	(?)	Wellington	N	N	Top of casing	0.1	1280.5	21.80	9-3-54	do
10-1-5chb	NW SW sec. 5	G. C. Stevens	Du	16.2	48	Sand	Terrace deposits	Cy, W	S	Base of pump	1.0	1277.8	9.53	9-2-54	East well.
*10-1-17de1	SW SE sec. 17	City of Longford	D	100	8	Sandstone	Dakota Formation	Cy, E	P	Land surface	0.0	1378.3			West well, yields 100 gpm.
10-1-17de2	SW SE sec. 17	do.	D	110	8	do.	do.	L, E	P	do.	0.0	1392.5			
10-1-22bba	NW NW sec. 22	J. M. Schneider	D	99.9	6	do.	do.	Cy, W	S	Base of pump	0.6	1330.5	41.88	9-3-54	Abandoned domestic well.
10-1-29bc	SW NW sec. 29	do.	Sp			do.	do.	S	S	Land surface	0.0	1297.1	42.00	9-2-54	
10-1-30cbe	NW SW sec. 30	Geo. Emrich	Du	42.3	48	do.	do.	N	N	Top of platform	0.3	1403.7			
10-2-12cc	T. 10 S., R. 2 E., SW SW sec. 12	F. Roderick	Du	48.7	36	Limestone	Nolan Limestone	Cy, W	S	Base of pump	0.3		48.40	9-30-54	Abandoned stock well.
10-2-17aa1	NE NE sec. 17	Henry Youse	Du	37.3	48	do.	Wellington	Cy, W	N	do.	0.8		30.90	9-30-54	do
10-2-17aa2	NE NE sec. 17	U. S. Geol. Survey	A	15.0	4	Clay	Terrace deposits	N	N	Land surface	0.0	1213.7	7.20	6-8-54	Did not reach bedrock.
10-2-17cc	SW SW sec. 17	do.	A	14.0	4	do.	do.	N	N	do.	0.0	1207.3	9.20	6-8-54	Did not reach bedrock.
*10-2-17cd1	SE SW sec. 17	Lawrence Falen	D	55.0	8	Sand	do.	Cy, W	S	Base of pump	1.0	1209.6	12.50	6-9-54	Did not reach bedrock.
10-2-17cd2	SE SW sec. 17	U. S. Geol. Survey	A	19.0	4	do.	do.	N	N	Land surface	0.0	1206.8	11.20	6-8-54	Did not reach bedrock.
10-2-18ada	SE NE sec. 18	do.	A	19.0	4	Clay	do.	N	N	do.	0.0	1213.3	9.40	6-8-54	do
10-2-20ba	NE NE sec. 20	do.	A	19.0	4	Sand	do.	N	N	do.	0.0	1203.7	16.00	6-8-54	do
*10-2-23bb	NW NW sec. 23	School District	D	65.0	6	Limestone	Nolan Limestone	Cy, H	D	Base of pump	0.6		49.90	6-8-54	
10-2-27ad	SE NE sec. 27	A. H. Green	D	51.0	6	do.	Winfield Limestone	Cy, W	S	do.	0.2		34.15	9-30-54	
10-2-31cc	SW SW sec. 31	J. W. Chestnut	D	65.4	6	(?)	Wellington	Cy, W	S	do.	0.2		31.83	9-30-54	
10-2-31aa	NE NE sec. 31	U. S. Geol. Survey	A	19.0	4	Sand	Terrace deposits	N	N	Land surface	0.0	1192.3	17.30	6-8-54	Did not reach bedrock.



## LOGS OF TEST HOLES

On the following pages are given the logs of 28 test holes drilled in Clay County (Pl. 2). The holes were drilled chiefly in areas of thick alluvial deposits, but several were drilled to determine the character, thickness, and water-bearing properties of the Dakota Formation. One test hole was drilled through the gypsum deposits in the northeastern part of the county. The geologic cross sections shown in Plate 3 are based principally on data obtained from test-hole logs. Many holes were drilled with a power auger to determine the depth to water level or as a guide in preparing the geologic map. Logs of these auger holes are not given, as most of them were very shallow. The test holes were drilled by the State Geological Survey. Samples were collected and studied in the field and later examined microscopically in the laboratory.

6-1-2bac.—*Drillers log of test hole in SW¼ NE¼ sec. 2, T. 6 S., R. 1 E. Near the site of Clifton city wells. Static water level 21 feet below land surface. Drilled by Layne-Western Co.*

	Thickness, feet	Depth, feet
Soil .....	4	4
Clay .....	24	28
Sand, fine, gray .....	2	30
Sand, coarse, gray .....	3	33
Sand and gravel, gray .....	10	43
Clay .....	1	44
Sand, fine .....	2	46
Sand, coarse, gray, and gravel on shale .....	16	62

6-1-2bb.—*Sample log of test hole in NW¼ NW¼ sec. 2, T. 6 S., R. 1 E. On east side of road south of Vining, 0.2 mile south of section corner. Drilled June 17, 1954. Surface altitude, 1,272.0 feet.*

	Thickness, feet	Depth, feet
QUATERNARY—Pleistocene		
Wisconsinan Stage—Terrace deposits		
Silt, black .....	3	3
Silt, tan .....	3	6
Clay, sticky, tan to gray .....	7	13
Clay, slightly sandy, dark gray .....	7	20
Clay, sandy, soft, light gray .....	3	23
Sand, fine to coarse, and fine to medium gravel .....	14	37
Clay, greenish gray .....	23	60
PERMIAN—Leonardian		
Wellington Formation .....	9	69
Dolomite, light gray .....	1	70

6-1-2c.—Sample log of test hole near SW cor. sec. 2, T. 6 S., R. 1 E. On east side of road, 200 feet north of section corner. Drilled June 17, 1954. Surface altitude, 1,256.8 feet; depth to water, 7.10 feet, June 18, 1954.

	Thickness, feet	Depth, feet
QUATERNARY—Pleistocene		
Recent and Wisconsinan Stages—Alluvium and terrace deposits		
Silt, black	3	3
Sand, fine to coarse, and fine to medium gravel	13	16
Clay, blue gray	4	20
Gravel, medium, and coarse arkosic sand	37	57
Kansan Stage—unnamed formation		
Sand, medium to coarse, and fine to medium limestone and arkosic gravel	6	63
Gravel, fine to medium; mostly sandstone and limestone pebbles	4	67
PERMIAN—Leonardian		
Wellington Formation		
Shale, platy, black and gray	3	70

6-1-11b.—Sample log of test hole in SW¼ NW¼ sec. 11, T. 6 S., R. 1 E. On east side of road 0.3 mile south of section corner, at north end of row of large cottonwood trees. Drilled June 18, 1954. Surface altitude, 1,258.3 feet.

	Thickness, feet	Depth, feet
QUATERNARY—Pleistocene		
Recent and Wisconsinan Stages—Alluvium and terrace deposits		
Silt, black	2	2
Silt, tan	3	5
Sand, fine to coarse, and fine to medium arkosic gravel,	12	17
Clay, sandy, blue	3	20
Sand, fine to coarse, and fine to medium arkosic gravel.	39	59
Kansan Stage—unnamed formation		
Clay, yellow	1	60
Gravel, fine to medium; many sandstone and limestone pebbles	8	68
Clay, yellow; contains some sandstone and limestone gravel	1	69
PERMIAN—Leonardian		
Wellington Formation		
Shale, platy, blue black	1	70

6-1-14bb.—Sample log of test hole in NW¼ NW¼ sec. 14, T. 6 S., R. 1 E. On east side of road, 25 feet south of point where road turns west. Drilled June 18, 1954. Surface altitude, 1,267.0 feet.

	Thickness, feet	Depth, feet
QUATERNARY—Pleistocene		
Recent Stage—Dune sand		
Sand, very fine	10	10
Recent and Wisconsinan Stages—Alluvium and terrace deposits		
Silt, light tan, very soft	2	12
Clay, gray	3	15

	Thickness, feet	Depth, feet
Silt, black .....	3	18
Silt, light tan .....	8	26
Sand, very fine .....	4	30
Gravel, medium to coarse, arkosic, and coarse sand ..	26	56
Sand and gravel, arkosic, mixed with green clay .....	4	60
Gravel, fine to medium, arkosic, and fine to coarse sand .....	10	70
Kausan Stage—unnamed formation		
Gravel, fine to medium, chiefly limestone and sandstone pebbles, and coarse sand .....	6	76
PERMIAN—Leonardian		
Wellington Formation		
Shale, platy, blue black .....	1	77
6-1-15da.—Sample log of test hole in NE¼ SE¼ sec. 15, T. 6 S., R. 1 E. On west side of road, about 250 feet north of quarter-mile hedge row. Drilled June 18, 1954. Surface altitude, 1,261.2 feet; depth to water, 22.00 feet June 18, 1954.		
QUATERNARY—Pleistocene		
Wisconsinan Stage—Terrace deposits		
Clay and silt, black .....	2	2
Clay, silty, tan .....	3	5
Clay, tough, black .....	2	7
Silt, sandy, tan .....	13	20
Sand, fine to coarse, arkosic, and fine gravel .....	10	30
Gravel, fine to medium, arkosic, and fine to coarse sand .....	30	60
PERMIAN—Leonardian		
Wellington Formation		
Limestone, gray .....	3	63
6-1-22aad.—Sample log of test hole in NE¼ NE¼ sec. 22, T. 6 S., R. 1 E. On west side of road, midway between abandoned farm house on east and house on west. Drilled June 22, 1954. Surface altitude, 1,263.0 feet; depth to water, 27.40 feet June 25, 1954.		
QUATERNARY—Pleistocene		
Wisconsinan Stage—Terrace deposits		
Silt, clayey, black .....	2	2
Silt, tan .....	6	8
Clay, gray .....	10	18
Clay, very sandy, light gray .....	3	21
Clay, sandy, pinkish tan .....	4	25
Sand, fine to coarse, and fine to medium gravel .....	35	60
Sand, fine to coarse, fine to medium gravel, and yellow clay .....	3	63
PERMIAN—Leonardian		
Wellington Formation		
Limestone, hard, dark gray .....	0.5	63.5

6-2-31cd.—Sample log of test hole in SE¼ SW¼ sec. 31, T. 6 S., R. 2 E. On north side of road 15 feet west of drive to house on hill south of road. Drilled June 22, 1954. Surface altitude, 1,253.8 feet; depth to water, 29.30 feet June 25, 1954.

QUATERNARY—Pleistocene		
Wisconsinan Stage—Terrace deposits	Thickness, feet	Depth, feet
Silt, black	2	2
Silt, sandy, gray to tan	9	11
Clay, silty; contains some sand and ironstone	15	26
Sand, fine	7	33
Clay, sandy, tan	3	36
Sand, fine to coarse	8	44
Clay, yellow	3	47

PERMIAN—Leonardian

Wellington Formation

Shale, or soft limestone, gray to black	1	48
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6-2-31dc.—Sample log of test hole in SW¼ SE¼ sec. 31, T. 6 S., R. 2 E. On north side of road about 500 feet east of drive to house on north side of road. Drilled June 22, 1954. Surface altitude, 1,232.9 feet; depth to water, 18.62 feet June 22, 1954.

QUATERNARY—Pleistocene		
Wisconsinan Stage—Terrace deposits	Thickness, feet	Depth, feet
Silt, black	7	7
Silt, sandy, tan	5	12
Clay, blue gray	3	15
Sand, fine to coarse, and fine to medium arkosic gravel,	29	44
Clay, greenish gray	2	46
Sand, fine to coarse, and fine to medium gravel	14	60
Sand, fine to medium, arkosic	5	65

PERMIAN—Leonardian

Wellington Formation

Shale, dark gray, grades downward into harder shale and limestone	5	70
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6-2-32ccc.—Sample log of test hole in SW cor. sec. 32, T. 6 S., R. 2 E. In center of old roadbed, 35 feet north of section corner. Drilled June 22, 1954. Surface altitude, 1,225.6 feet; depth to water, 5.01 feet June 25, 1954.

QUATERNARY—Pleistocene		
Recent and Wisconsinan Stages—Alluvium and terrace deposits	Thickness, feet	Depth, feet
Silt, tan	2	2
Sand, fine to coarse, and fine to medium arkosic gravel,	48	50
Sand, fine to medium; contains some green clay and sandstone gravel	10	60
Kansan Stage—unnamed formation		
Sand, fine to medium; contains much locally derived gravel and shale	9	69

PERMIAN—Leonardian

Wellington Formation

Shale, black	11	80
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6-2-32cd.—Sample log of test hole in SE¼ SW¼ sec. 32, T. 6 S., R. 2 E. At end of trail leading east along section line toward Republican River. Drilled June 23, 1954. Surface altitude, 1,233.7 feet.

	Thickness, feet	Depth feet
QUATERNARY—Pleistocene		
Recent and Wisconsinan Stages—Alluvium and terrace deposits		
Silt, black	2	2
Silt, tan	4	6
Sand, fine to medium	4	10
Sand, fine to coarse, and fine to medium arkosic gravel,	29	39
Clay, blocky, green	10	49
Sand and gravel, fine to medium	12	61
Clay, silty and sandy, dark gray	2	63
Kansas Stage—unnamed formation		
Gravel, fine to medium, mostly locally derived; some thin clay streaks	6	69
PERMIAN—Leonardian		
Wellington Formation		
Shale, micaceous, sandy, light gray	1	70

6-2-32dd.—Sample log of test hole in SE¼ SE¼ sec. 32, T. 6 S., R. 2 E. On north side of road, 175 feet east of creek bridge. Drilled June 23, 1954. Surface altitude, 1,225.2 feet; depth to water, 7.40 feet June 28, 1954.

	Thickness, feet	Depth feet
QUATERNARY—Pleistocene		
Recent and Wisconsinan Stages—Alluvium and terrace deposits		
Silt, black	3	3
Silt, tan	6	9
Silt, sandy, black	5	14
Sand, fine to coarse, and fine to medium arkosic gravel,	36	50
Clay, tan to gray	4	54
Kansas Stage—unnamed formation		
Gravel, locally derived, fine to medium; some yellow clay	12	66
PERMIAN—Leonardian		
Wellington Formation		
Shale, blue gray	4	70

6-2-33cd.—Sample log of test hole in SE¼ SW¼ sec. 33, T. 6 S., R. 2 E. On south side of road, 40 feet east of railroad track. Drilled June 24, 1954. Surface altitude, 1,236.9 feet; depth to water, 21.30 feet June 24, 1954.

	Thickness, feet	Depth feet
QUATERNARY—Pleistocene		
Wisconsinan Stage—Terrace deposits		
Silt, black	2	2
Silt, tan	5	7
Silt, sandy, tan; contains snail shells	8	15
Sand, fine to medium, arkosic	19	34
Sand, medium to coarse, and fine to medium arkosic gravel	36	70



Kansas Stage—unnamed formation		Thickness, feet	Depth, feet
Sand, fine to coarse, mostly locally derived		5	75
PERMIAN—Leonardian			
Wellington Formation			
Shale, blocky, light gray		5	80
6-2-34cd.—Sample log of test hole in SE cor. SW¼ sec. 34, T. 6 S., R. 2 E. On north side of road, 100 feet west of bridge across drainage ditch. Drilled June 22, 1954. Surface altitude, 1,239.2 feet; depth to water, 24.50 feet June 28, 1954.			
QUATERNARY—Pleistocene			
Wisconsinan Stage—Terrace deposits		Thickness, feet	Depth, feet
Silt, black		2	2
Silt, tan		9	11
Clay, hard, gray		3	14
Clay, sandy, tan		14	28
Sand, fine to coarse, and fine arkosic gravel		16	44
PERMIAN—Leonardian			
Wellington Formation			
Shale, slightly calcareous, blue gray and tan		2	46
6-4-19cb.—Sample log of test hole in NW¼ SW¼ sec. 19, T. 6 S., R. 4 E. On top of second hill north of corner, west of bend in creek where creek is nearest to road. Drilled October 14, 1954.			
PERMIAN—Leonardian			
Wellington Formation		Thickness, feet	Depth feet
Shale, greenish buff; contains some thin dolomitic streaks		4	4
Shale, noncalcareous, yellowish green		3	7
Shale, noncalcareous, bright green		1	8
Shale, silty, noncalcareous, chocolate brown		6.5	14.5
Shale, mottled red brown and green		2.5	17
Shale, silty, noncalcareous, red brown		1	18
Shale, greenish gray		3	21
Shale, noncalcareous, light gray		1	22
Shale, noncalcareous, greenish gray		3.5	25.5
Shale, noncalcareous, red brown		1	26.5
Shale, noncalcareous, greenish gray		1.5	28
Shale, silty, noncalcareous, red brown		1	29
Shale, noncalcareous, green; contains some gypsum stringers		2.5	31.5
Shale, platy, greenish gray		1.5	33
Gypsum		1	34
Shale, green		1.5	55.5
Shale, greenish gray; contains gypsum		1	36.5
Gypsum; contains 6-inch shale layer at 38.5 to 39		3.5	40
Gypsum, very clean		8	48
Gypsum, shaly		0.5	48.5
Gypsum		0.5	49
PERMIAN—Wolfcampian			
Nolans Limestone—Herington Limestone member			
Limestone, hard, dolomitic		3	52

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7-1-6bb.—*Sample log of test hole in NW cor. sec. 6, T. 7 S., R. 1 E. About 60 feet east and 8 feet south of section corner. Drilled October 7, 1954. Surface altitude, 1,390.0 feet.*

QUATERNARY—Pleistocene		
Wisconsinan Stage—Peoria Formation		
	Thickness, feet	Depth, feet
Silt, dark gray to brown . . . . .	2	2
Silt, tan . . . . .	2	4
Clay, buff; contains a few ironstone pebbles . . . . .	5	9
Clay, sandy, brown to gray . . . . .	2	11
CRETACEOUS—Gulfian		
Dakota Formation		
Sandstone, fine grained, brown . . . . .	13.5	24.5
Clay, light gray to brown . . . . .	4.5	29
Clay, sandy, buff to brown . . . . .	16	45
Clay, sandy, green gray . . . . .	4	49
Clay, sandy, red brown to buff . . . . .	5	54
Clay, greenish gray . . . . .	5	59
Clay and sandstone . . . . .	5	64
Shale, clayey, dark gray . . . . .	10	74
Sandstone, very hard; contains pyrite . . . . .	3	77
Shale, clayey, sandy, dark gray . . . . .	5	82
Shale, clayey, sandy, light gray . . . . .	9.5	91.5
Clay, white, and sandstone . . . . .	3.5	95
Shale, clayey, dark gray . . . . .	3	98
Lignite . . . . .	3	101
Shale, clayey, sandy, gray . . . . .	27	128
PERMIAN—Leonardian		
Wellington Formation		
Shale, calcareous, very hard, gray, interbedded with limestone . . . . .	5	133
7-2-3bb.— <i>Sample log of test hole in NW cor. sec. 3, T. 7 S., R. 2 E. In edge of field south of road, 40 feet east of section corner. Drilled June 25, 1954. Surface altitude, 1,237.9 feet; depth to water, 23.10 feet June 28, 1954.</i>		
QUATERNARY—Pleistocene		
Wisconsinan Stage—Terrace deposits		
	Thickness, feet	Depth, feet
Silt, black . . . . .	3	3
Silt, tan . . . . .	8	11
Clay, compact, tan . . . . .	5	16
Clay, compact, dark gray . . . . .	2	18
Clay, sandy, tan . . . . .	6	24
Sand, fine to coarse, and fine to medium arkosic gravel, Clay . . . . .	23	47
Clay . . . . .	2	49
Sand, fine to coarse, and fine to medium arkosic gravel, Clay, green . . . . .	8	57
Clay, green . . . . .	4	61
Gravel, fine to medium, arkosic . . . . .	8.5	69.5
PERMIAN—Leonardian		
Wellington Formation		
Shale, dark gray . . . . .	10.5	80

7-2-6bb.—Sample log of test hole in NW¼ NW¼ sec. 6, T. 7 S., R. 2 E. About 25 feet west of west side of old school house foundation. Drilled October 13, 1954. Surface altitude, 1,315.0 feet; depth to water, 82.20 feet October 22, 1954.

<b>QUATERNARY—Pleistocene</b>		
Illinoian Stage—Loveland Formation		
Clay, sandy, reddish tan	Thickness, feet	Depth, feet
Clay, sandy, buff	4	4
Clay, sandy, yellow	11	15
	9	24
<b>CRETACEOUS—Gulfian</b>		
Dakota Formation		
Sandstone, very fine grained, loosely cemented, white to yellow	26	50
Sandstone, or sand, fine, brown; contains some mica and ironstone	38	88
<b>PERMIAN—Leonardian</b>		
Wellington Formation		
Clay, light gray	5	93
Shale, hard, platy, somewhat calcareous, dark greenish gray	9	102
Shale, light greenish gray	11	113
Shale, hard, platy, dark greenish gray	5.5	118.5
Shale, hard, platy, very dark gray to black	13.5	132
Shale, platy, alternately hard and soft, very dark gray to black	8	140

7-3-5bb.—Sample log of test hole near NW cor. sec. 5, T. 7 S., R. 3 E. About 130 feet east of highway and 12 feet south of center of road. Drilled October 14, 1954. Surface altitude, 1,333.7 feet; depth to water, 19.20 feet October 22, 1954.

<b>CRETACEOUS—Gulfian</b>		
Dakota Formation		
Clay, sandy, tan to gray	Thickness, feet	Depth, feet
Clay, sandy, red brown	4	4
Clay, sandy, red and gray	15	19
Clay, sandy, black and gray; contains some lignite	11	30
Clay, very sandy, lignite streaks	5	35
	10.5	45.5
<b>PERMIAN—Leonardian</b>		
Wellington Formation		
Shale, hard, platy, gray	2.5	48
Shale, black	5	53
Shale, dark gray	4	57
Limestone, hard	1	58

8-2-10aa.—Sample log of test hole in NE¼ NE¼ sec. 10, T. 8 S., R. 2 E. On west side of road, about 175 feet south of culvert. Drilled June 28, 1954. Surface altitude, 1,194.6 feet; depth to water, 11.80 feet June 30, 1954.

<b>QUATERNARY—Pleistocene</b>		
Recent Stage—Alluvium		
Silt and clay, tan	Thickness, feet	Depth, feet
Silt, black	5	5
	1	6

	Thickness, feet	Depth, feet
Silt, tan .....	4	10
Sand, fine to coarse, and fine to medium arkosic gravel, .....	10	20
Gravel, medium to coarse, arkosic .....	8	28
PERMIAN—Leonardian		
Wellington Formation		
Shale, blocky, blue gray .....	12	40
8-3-17ba.— <i>Sample log of test hole in NE¼ NW¼ sec. 17, T. 8 S., R. 3 E. In east borrow pit, 20 feet south of entrance to C. E. Mullen farm. Drilled June 28, 1954. Surface altitude, 1,179.3 feet; depth to water, 9.10 feet June 30, 1954.</i>		
QUATERNARY—Pleistocene		
Wisconsinan Stage—Terrace deposits		
Silt, brown .....	6	6
Sand, fine to coarse, and fine to medium arkosic gravel, .....	24	30
Sand, fine to coarse, and fine to arkosic gravel .....	12	42
PERMIAN—Wolfcampian		
Nolans Limestone		
Limestone, trace of chert, yellow to gray .....	2	44
8-3-35da.— <i>Sample log of test hole in NE¼ SE¼ sec. 35, T. 8 S., R. 3 E. In west borrow pit. Drilled June 29, 1954. Surface altitude, 1,161.5 feet; depth to water, 6.40 feet June 29, 1954.</i>		
QUATERNARY—Pleistocene		
Recent Stage—Alluvium		
Sand, fine .....	14	14
Sand, fine to coarse, and fine gravel .....	5	19
Sand, medium to coarse, and fine to coarse arkosic gravel .....	10.5	29.5
PERMIAN—Wolfcampian		
Doyle Shale		
Shale, red and gray .....	5.5	35
Shale, limy, white .....	1	36
Limestone, hard .....	1	37
8-3-35dd.— <i>Sample log of test hole in SE¼ SE¼ sec. 35, T. 8 S., R. 3 E. In south side of road, 40 feet west of center of old black-top road. Drilled June 29, 1954. Surface altitude, 1,159.3 feet.</i>		
QUATERNARY—Pleistocene		
Recent Stage—Alluvium		
Silt, black .....	2	2
Silt, tan .....	3	5
Sand, fine to coarse .....	7	12
Sand, fine to coarse, and fine to medium arkosic gravel .....	15	27
PERMIAN—Wolfcampian		
Doyle Shale		
Shale, chalky, red and green .....	3	30

8-3-36bb.—Sample log of test hole near NW cor. sec. 36, T. 8 S., R. 3 E. On east side of road across from Broughton store and service station. Drilled June 29, 1954. Surface altitude, 1,181.9 feet; depth to water, 22.00 feet July 12, 1954.

QUATERNARY—Pleistocene		
Wisconsinan Stage—Terrace deposits	Thickness, feet	Depth, feet
Silt, brown	3	3
Silt, sandy, black	2	5
Clay, brown to tan	6	11
Clay, sandy, tan	11	22
Sand, fine, and light-gray clay	8	30
Sand, fine to coarse, and fine to medium arkosic gravel	18	48

PERMIAN—Wolfcampian		
Doyle Shale		
Limestone, brown	2	50

9-4-29ac.—Sample log of test hole in SW¼ NE¼ sec. 29, T. 9 S., R. 4 E. South of fence, 40 feet west, 13 feet south of power pole. Drilled June 30, 1954. Surface altitude, 1,149.2 feet.

QUATERNARY—Pleistocene		
Recent Stage—Alluvium	Thickness, feet	Depth, feet
Silt, sandy, tan	15	15
Sand, fine to coarse, and fine to medium arkosic gravel	25	40

PERMIAN—Wolfcampian		
Matfield Shale		
Limestone, or limy shale, greenish gray and black	1	41

9-4-29bc.—Sample log of test hole in SW¼ NW¼ sec. 29, T. 9 S., R. 4 E. In field, 20 feet north of center of road. Drilled June 30, 1954. Surface altitude, 1,164.0 feet; depth to water, 29.90 feet July 12, 1954.

QUATERNARY—Pleistocene		
Wisconsinan Stage—Terrace deposits	Thickness, feet	Depth, feet
Silt, black	2	2
Silt, brown	3	5
Clay, tan to brown	7	12
Clay, dark gray	6	18
Clay, sandy, light gray	13	31
Sand, fine to coarse, and fine to medium arkosic gravel	22	53

PERMIAN—Wolfcampian		
Barneston Limestone		
Shale, green and red	4	57
Limestone, greenish gray; contains some chert	1	58

9-4-29bd.—Sample log of test hole in SE¼ NW¼ sec. 29, T. 9 S., R. 4 E. On east side of road, 35 feet north of point where road turns north. Drilled June 30, 1954. Surface altitude, 1,152.2 feet; depth to water, 18.20 feet July 12, 1954.

QUATERNARY—Pleistocene		
Recent Stage—Alluvium	Thickness, feet	Depth, feet
Silt, sandy, tan	10	10
Sand, medium to coarse, and fine to medium arkosic gravel	23	33

## PERMIAN—Wolfcampian

	Thickness, feet	Depth, feet
Barneston Limestone		
Limestone, greenish tan and black	1	34

9-4-32adc.—*Sample log of test hole in SW¼ SE¼ NE¼ sec. 32, T. 9 S., R. 4 E. On west side of road, 30 feet west and 10 feet north of windmill where highway turns east. Drilled July 1, 1954. Surface altitude, 1,136.0 feet; depth to water, 11.50 feet July 12, 1954.*

## QUATERNARY—Pleistocene

	Thickness, feet	Depth, feet
Wisconsinan Stage—Terrace deposits		
Silt, sandy, tan	4	4
Sand, fine	9	13
Gravel, fine to medium, arkosic	10	23
Clay, blue to black	2	25
Sand, fine to coarse, and fine to medium arkosic gravel	13	38

## PERMIAN—Wolfcampian

Matfield Shale		
Shale, red	2	40

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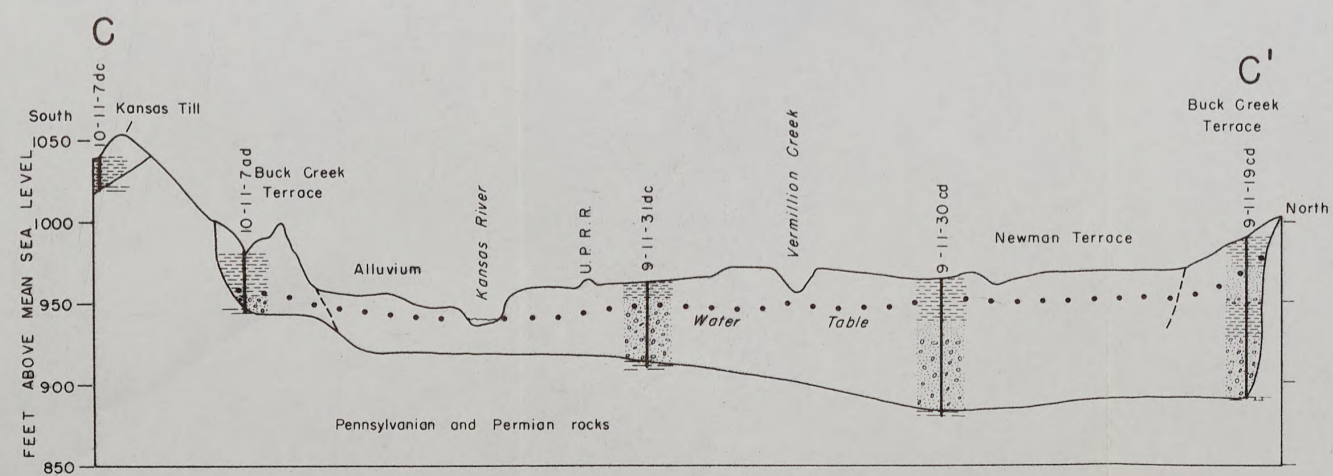
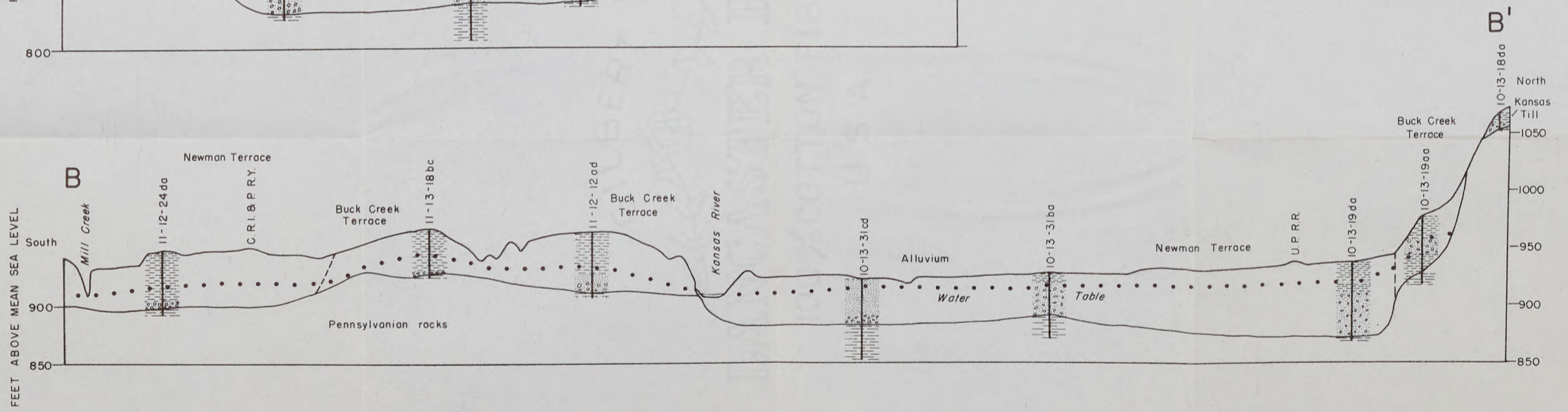
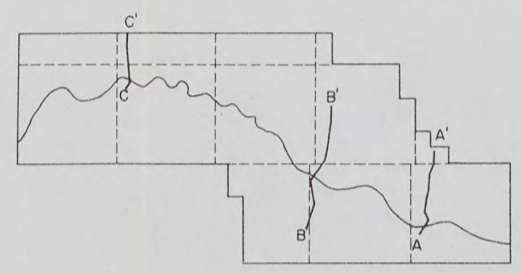
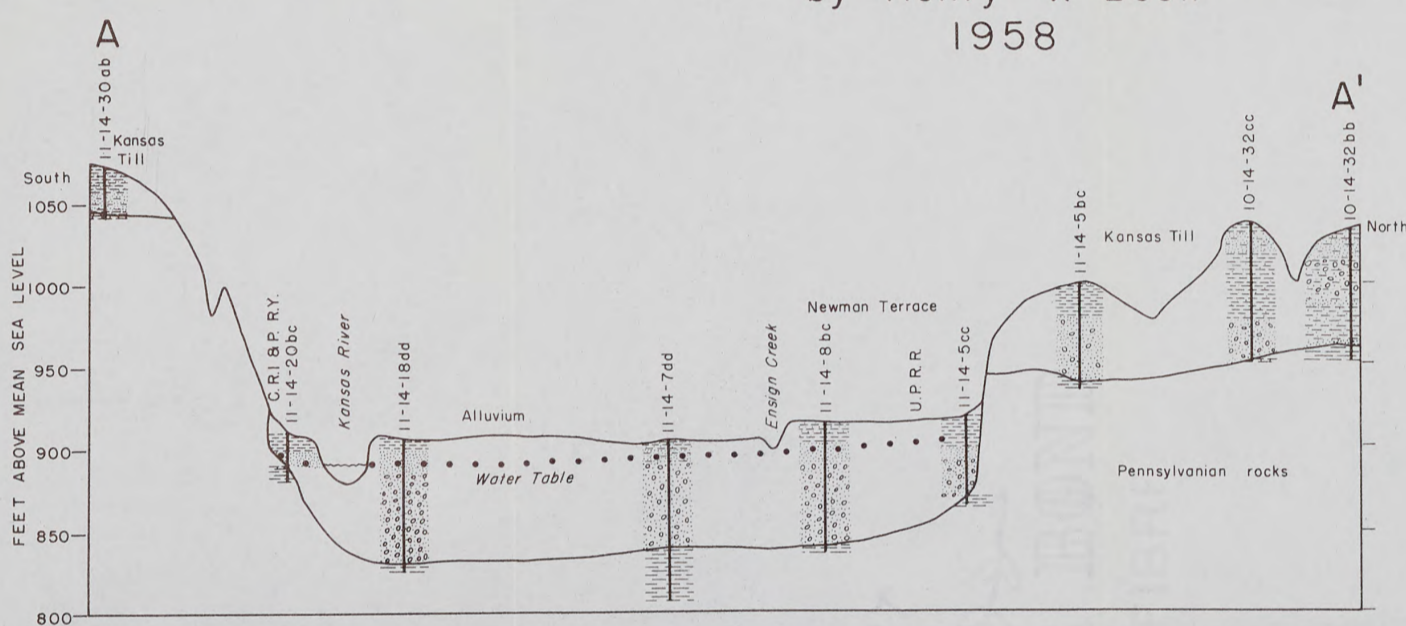
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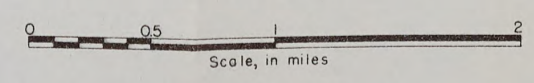
# GEOLOGIC SECTIONS OF KANSAS RIVER VALLEY BETWEEN WAMEGO AND TOPEKA VICINITY

by Henry V. Beck  
1958



### EXPLANATION

- Silt and clay
- Sand
- Gravel
- Shale
- Limestone



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### MAP OF KANSAS RIVER VALLEY BETWEEN WAMEGO AND TOPEKA VICINITY

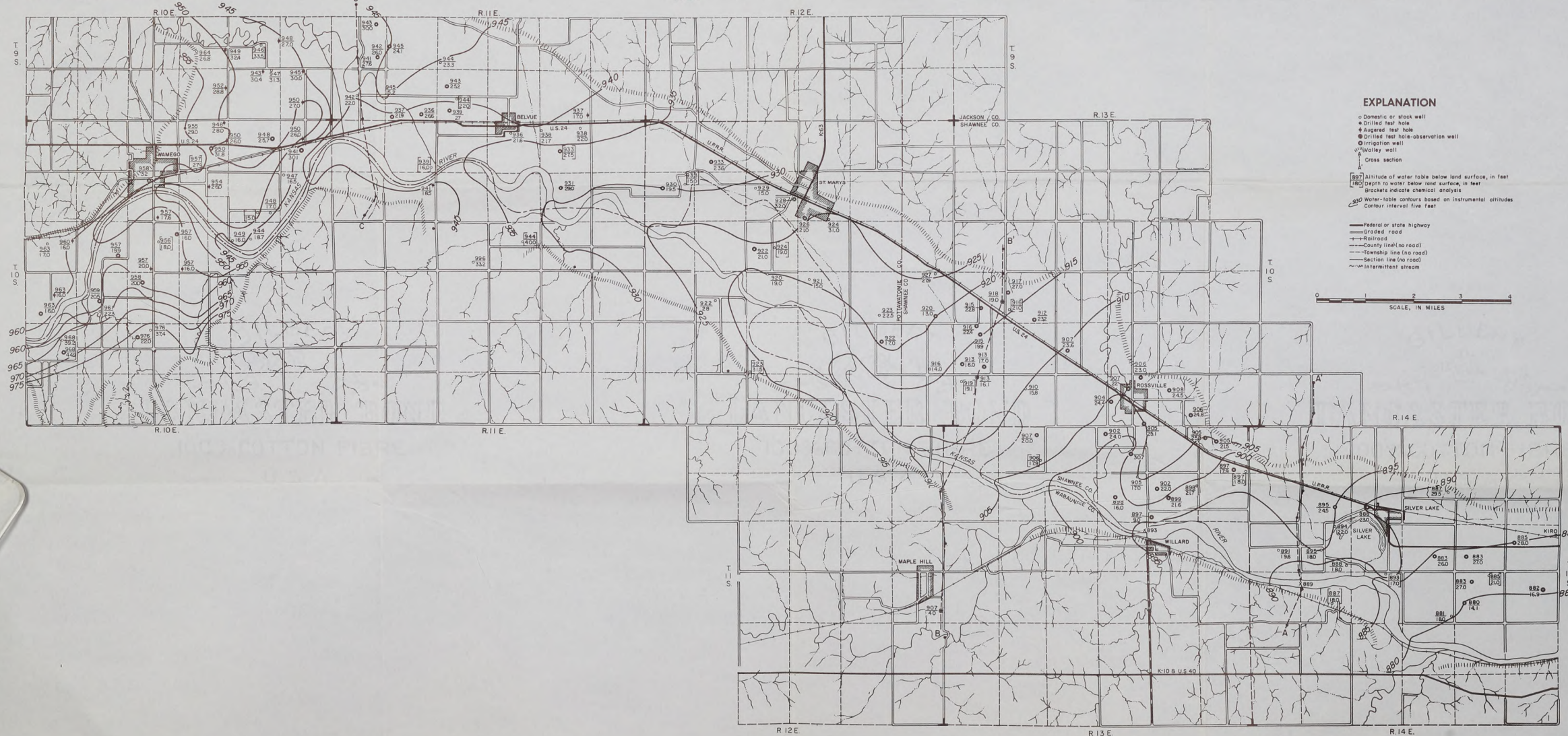
showing water-table contours and locations of wells and test holes

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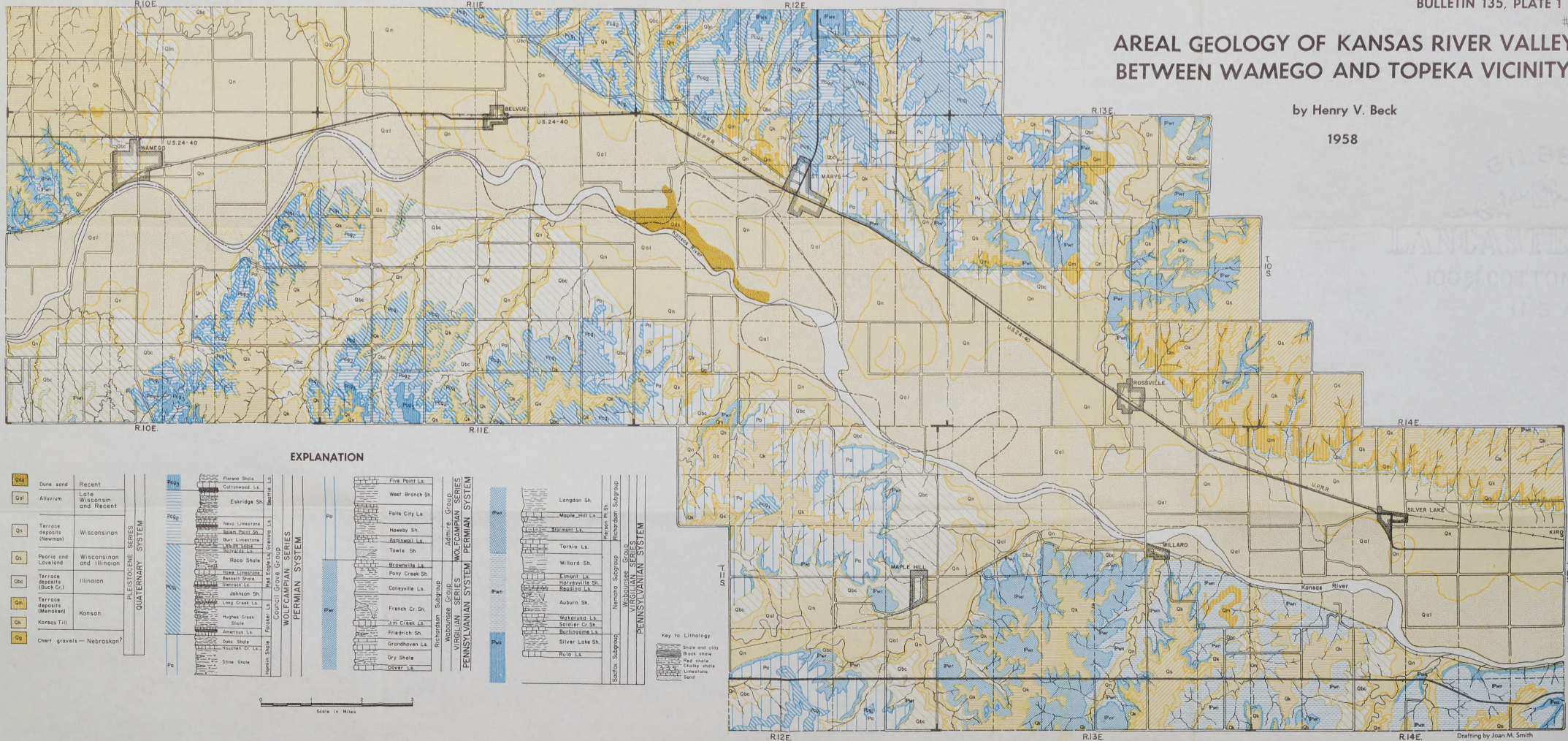
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### AREAL GEOLOGY OF KANSAS RIVER VALLEY BETWEEN WAMEGO AND TOPEKA VICINITY

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1958

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Plate 1

GILBEA  
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QUATERNARY SYSTEM	
Recent	Dune sand
Wisconsinan and Recent	Alfium
Wisconsinan	Terrace deposits (Wamego)
Wisconsinan and Illinoian	Pointe and Leeward
Illinoian	Terrace deposits (Black Cr.)
Kansan	Kansas Till
Nebraskan	Clay gravels - Nebraskan

GLACIOFLUVIAL SERIES	
Recent	Pointe and Leeward
Wisconsinan	Pointe and Leeward
Illinoian	Pointe and Leeward
Kansan	Pointe and Leeward
Nebraskan	Pointe and Leeward

PENNSYLVANIAN SYSTEM	
Recent	Pointe and Leeward
Wisconsinan	Pointe and Leeward
Illinoian	Pointe and Leeward
Kansan	Pointe and Leeward
Nebraskan	Pointe and Leeward

PERMIAN SYSTEM	
Recent	Pointe and Leeward
Wisconsinan	Pointe and Leeward
Illinoian	Pointe and Leeward
Kansan	Pointe and Leeward
Nebraskan	Pointe and Leeward

TRIASSIC SYSTEM	
Recent	Pointe and Leeward
Wisconsinan	Pointe and Leeward
Illinoian	Pointe and Leeward
Kansan	Pointe and Leeward
Nebraskan	Pointe and Leeward

CRETACEOUS SYSTEM	
Recent	Pointe and Leeward
Wisconsinan	Pointe and Leeward
Illinoian	Pointe and Leeward
Kansan	Pointe and Leeward
Nebraskan	Pointe and Leeward

PALEOZOIC SYSTEM	
Recent	Pointe and Leeward
Wisconsinan	Pointe and Leeward
Illinoian	Pointe and Leeward
Kansan	Pointe and Leeward
Nebraskan	Pointe and Leeward

# AREAL GEOLOGY OF CLAY COUNTY, KANSAS

By

Kenneth L. Walters and Charles K. Bayne

Bulletin 136

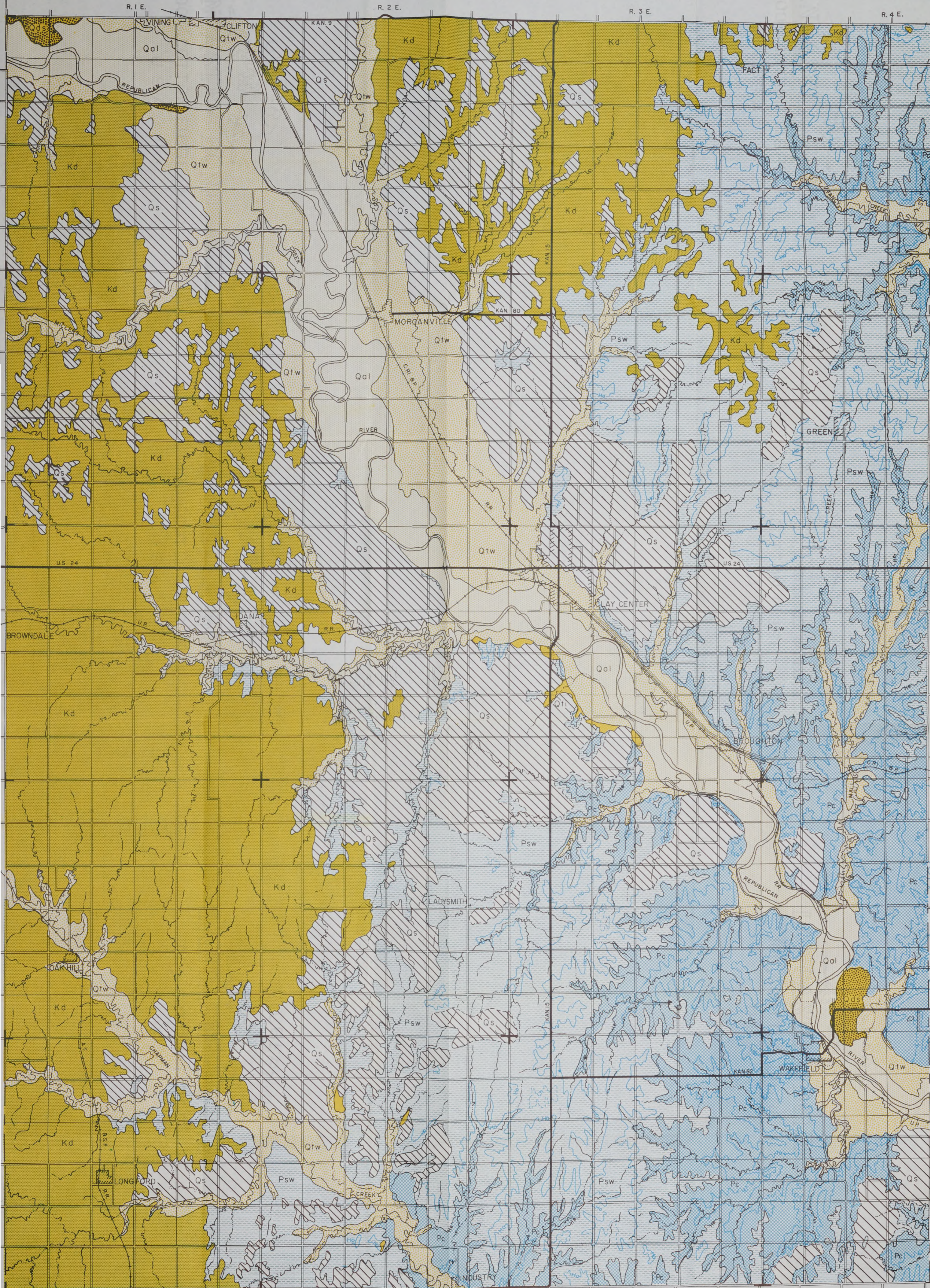
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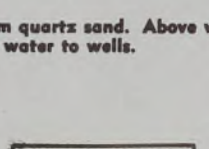
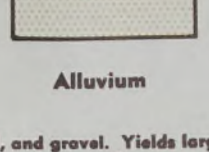
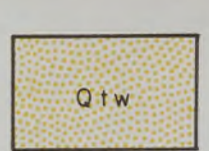
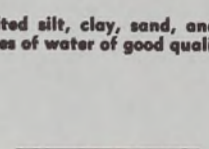
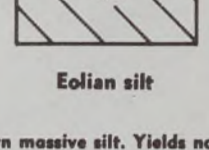
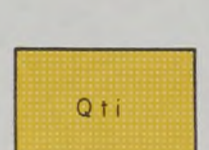
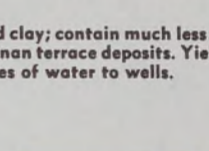
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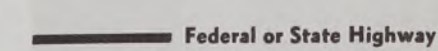
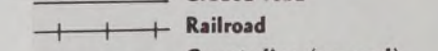
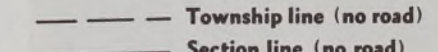
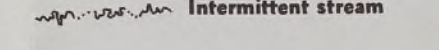

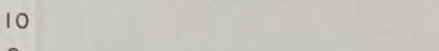

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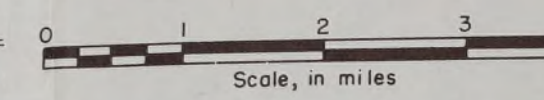


## EXPLANATION

-  **Q<sub>1</sub>s**  
Dune sand
-  **Q<sub>al</sub>**  
Alluvium  
Clay, silt, sand, and gravel. Yields large quantities of water of good quality to wells.
-  **Q<sub>1</sub>w**  
Wisconsinan terrace deposits  
Stream-deposited silt, clay, sand, and gravel. Yield large quantities of water of good quality to wells.
-  **Q<sub>s</sub>**  
Eolian silt  
Reddish-brown massive silt. Yields no water to wells.
-  **Q<sub>1</sub>i**  
Illinoian terrace deposits  
Chiefly silt and clay; contain much less sand and gravel than Wisconsinan terrace deposits. Yield small to moderate quantities of water to wells.
-  **K<sub>d</sub>**  
Dakota Formation  
Varicolored silt to sandy shale and massive sandstone; as mapped includes Kiowa Shale. Yields moderate amounts of water of good quality to wells.
-  **Ch**  
**P<sub>sw</sub>**  
Wellington Formation  
Chiefly gray shale containing a chert bed, Ch, and the Hellenberg Limestone member, Ho. Gypsum occurs locally at base of formation. Yields minor quantities of mineralized water to wells.
-  **H<sub>e</sub>**  
**P<sub>c</sub>**  
Chase Group  
Upper boundary marked by Herington Limestone member, He, of Niobona Limestone; includes as mapped, Cresswell Limestone member, Cr, of Winfield Limestone; Towanda Limestone member, To, of Doyla Shale; and Fort Riley Limestone member, Fr, of Barneson Limestone. Herington Limestone, Cresswell Limestone, and Towanda Limestone members locally yield small quantities of water to wells. Fort Riley Limestone member yields moderate quantities of water to wells.

QUATERNARY  
PLEISTOCENE  
CRETACEOUS  
GULFIAN  
LEONARDIAN  
PERMIAN  
WOLF CAMPIAN

-  Federal or State Highway
-  Graded road
-  Railroad
-  County line (no road)
-  Township line (no road)
-  Section line (no road)
-  Intermittent stream



Base compiled from maps prepared by the Soil Conservation Service

Drainage from map prepared by U. S. Dept. of Agriculture

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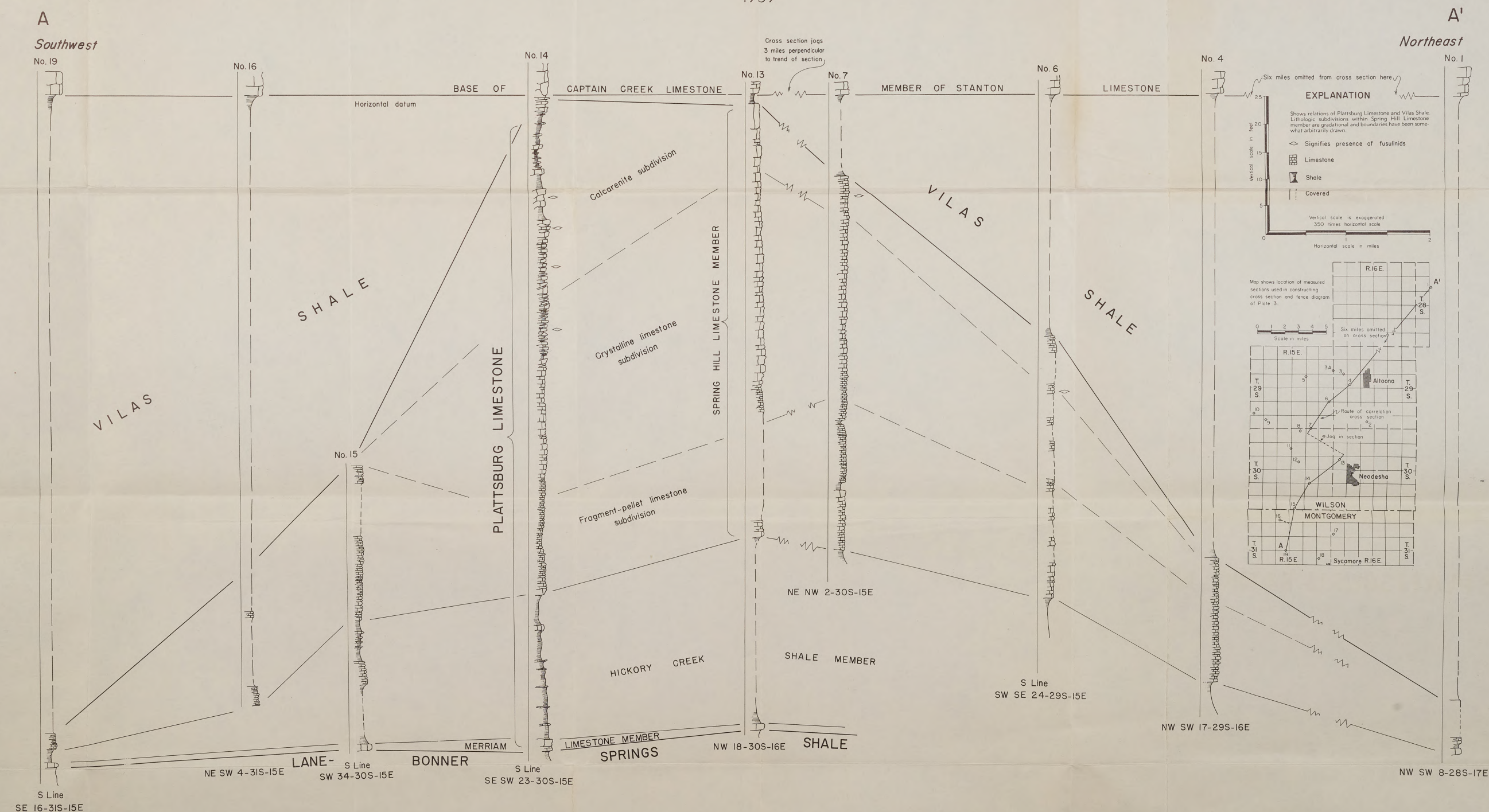
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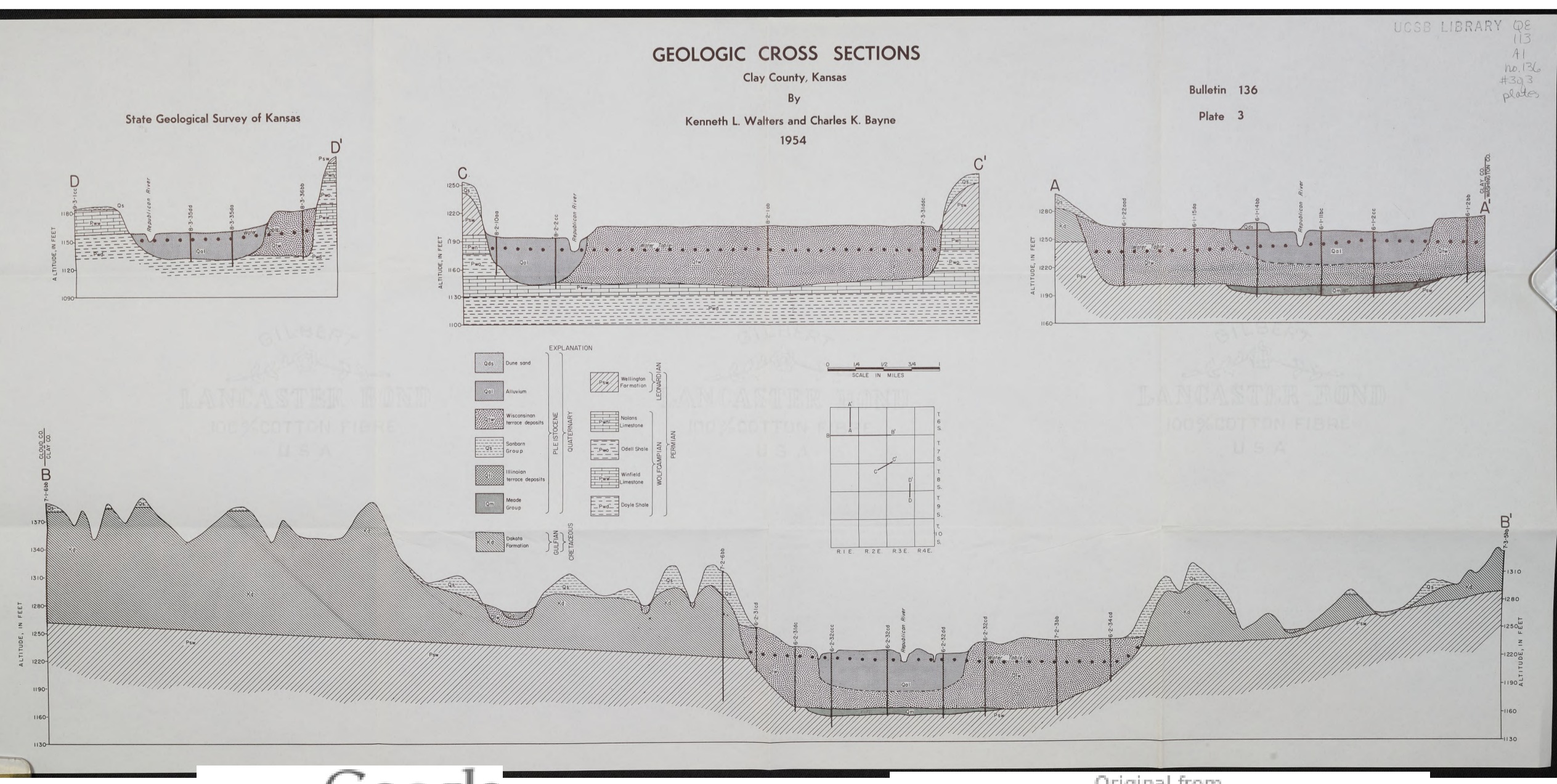
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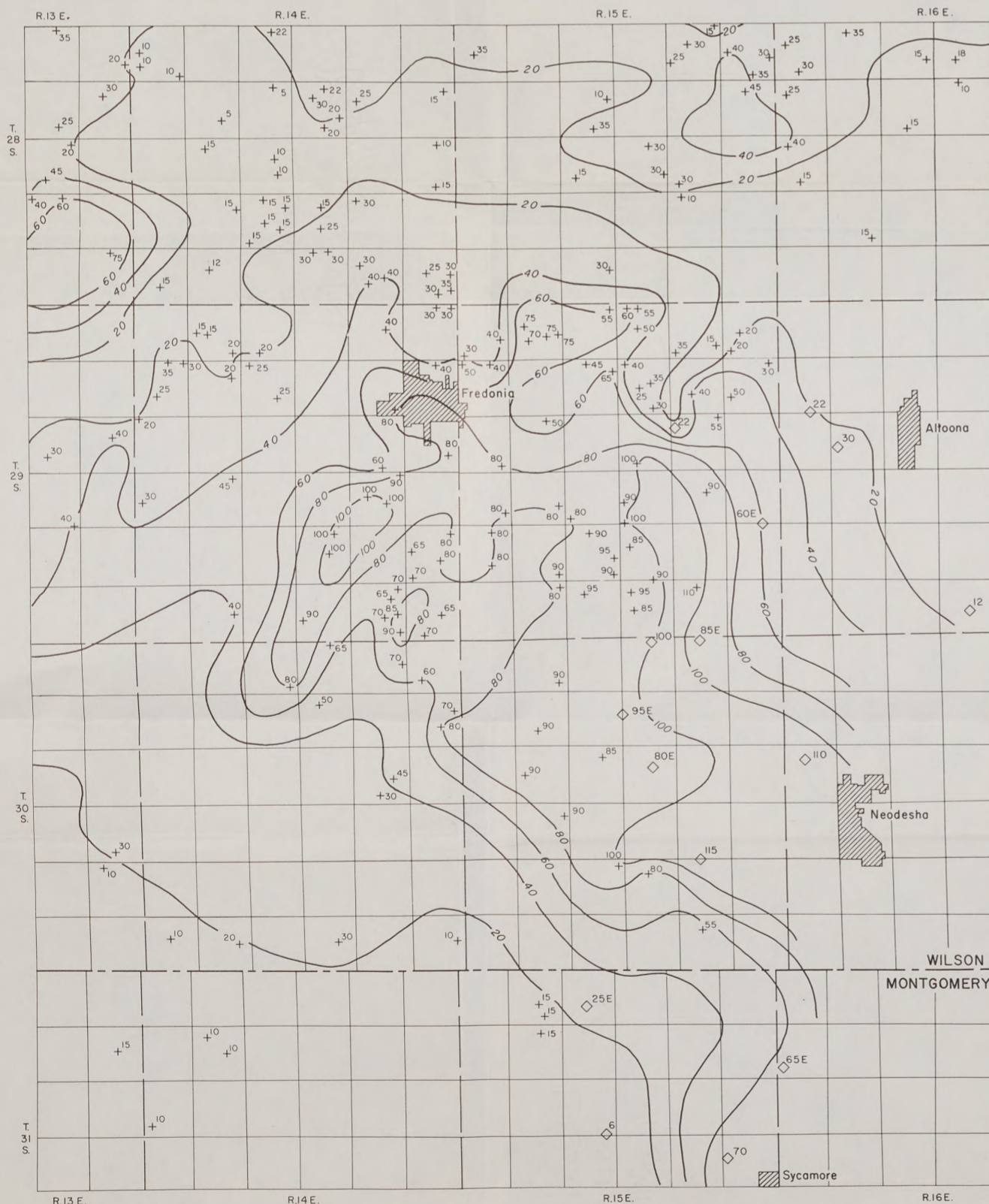
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no. 136  
#333  
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# ISOPACHOUS MAPS OF PLATTSBURG LIMESTONE AND VILAS SHALE, NEODESHA-FREDONIA AREA

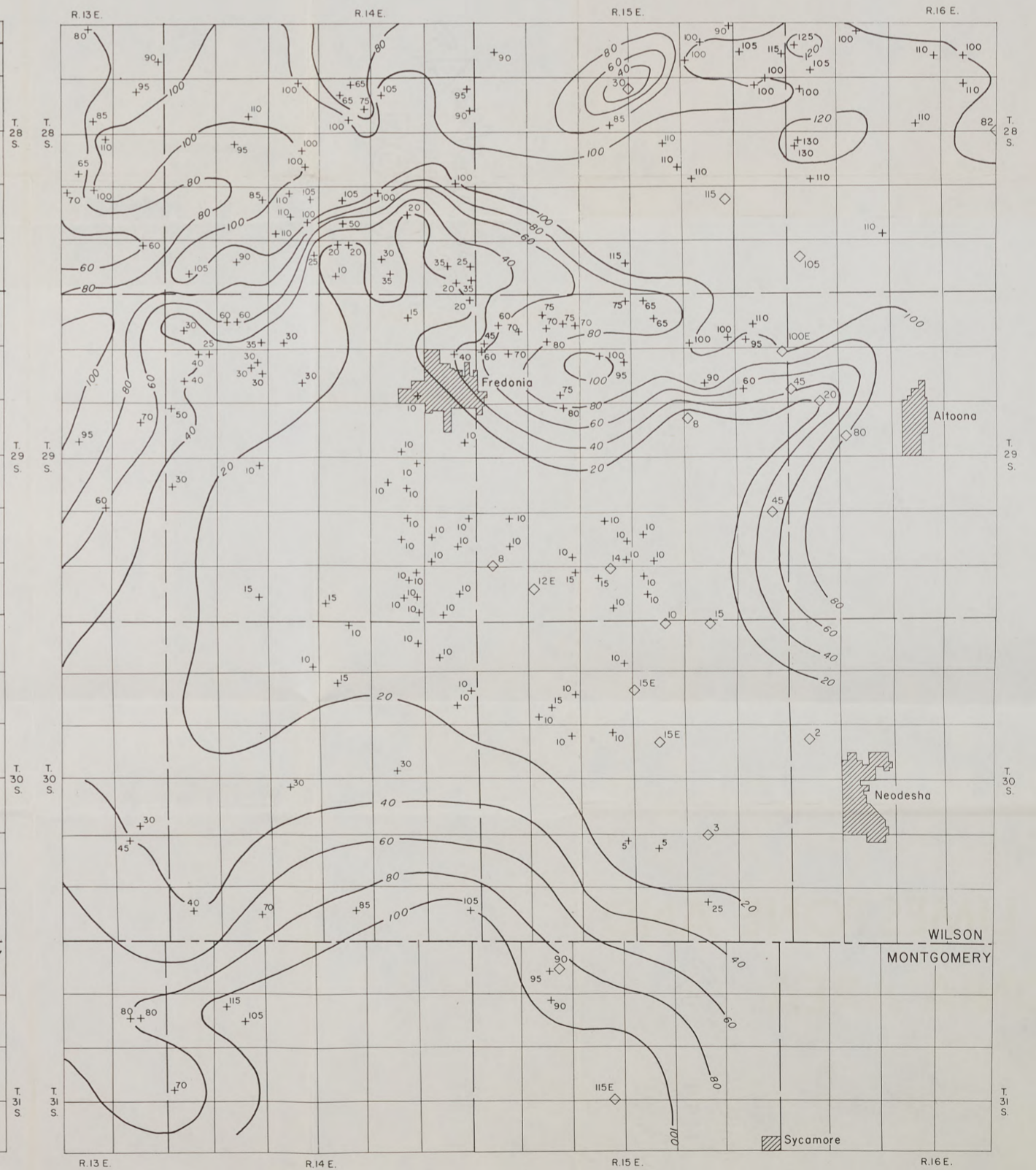
by JOHN W. HARBAUGH  
1959

STATE GEOLOGICAL SURVEY OF KANSAS

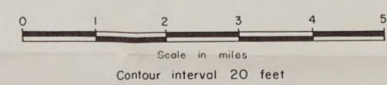
BULLETIN 134, PART 8 PLATE 1



A. Isopachous map of Plattsburg Limestone



B. Isopachous map of Vilas Shale



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#203  
plates

# MAP OF CLAY COUNTY, KANSAS

Showing Depth to Water, Location of Wells, Springs, and Test Holes for which  
Records are Given, and Water-Table Contours

By

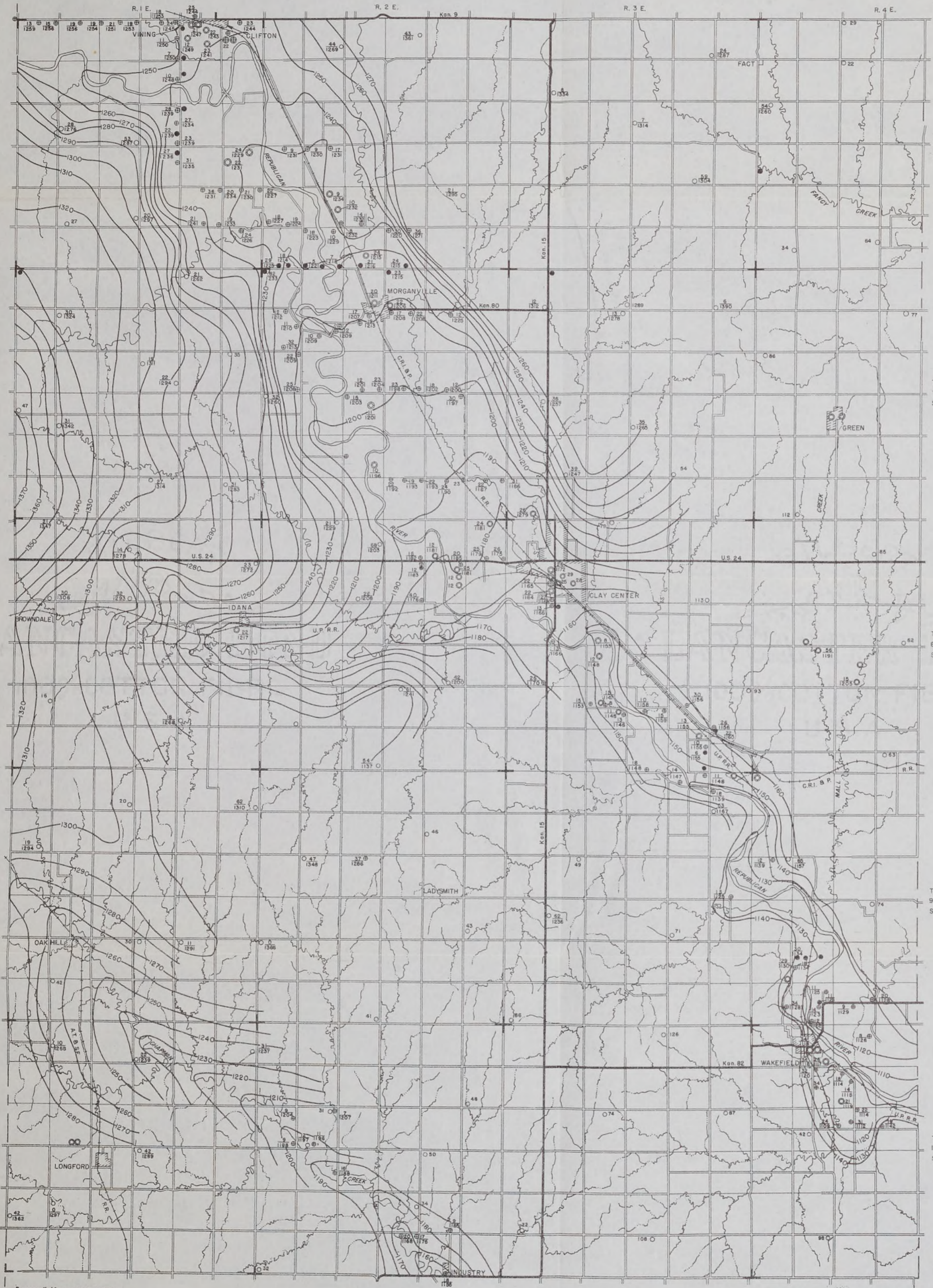
Kenneth L. Walters and Charles K. Bayne

1954

Bulletin 136

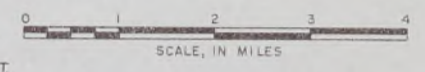
Plate 2

State Geological Survey of Kansas



## EXPLANATION

- Domestic or stock well
  - Drilled test hole
  - ⊕ Augered test hole
  - ⊙ Irrigation well
  - ⊙ Public supply well
  - ⊙ Industrial well
  - ⊙ Spring
- $\frac{12}{1312}$  Upper number refers to depth to water level; lower number refers to altitude of water level.
- 1250 — Water-table contours
- Federal or State Highway
  - Graded road
  - Railroad
  - - - County line (no road)
  - - - Township line (no road)
  - - - Section line (no road)
  - - - Intermittent stream



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# LITHOLOGIC DIVISIONS IN EXPOSED EASTERN PART OF PLATTSBURG BANK, NEODESHA-FREDONIA AREA

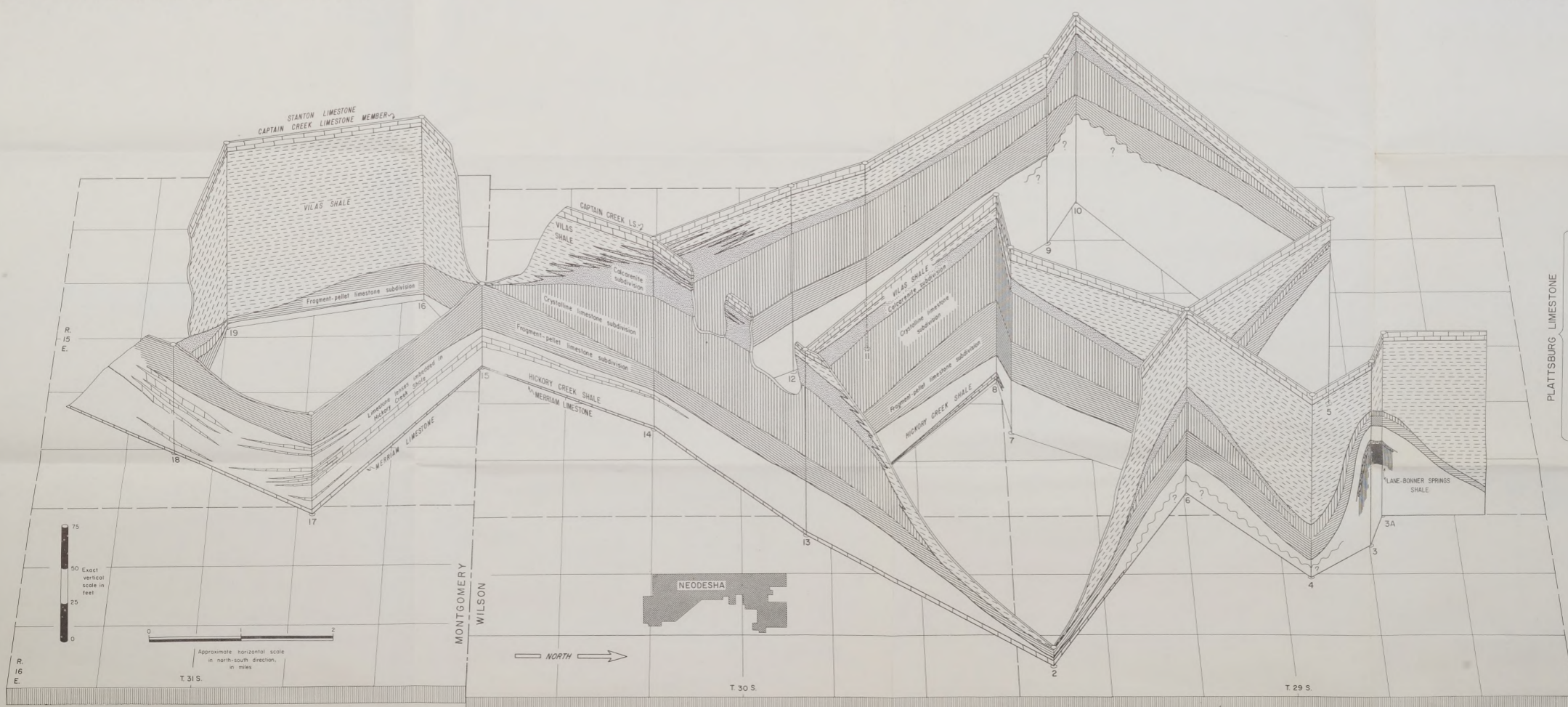
by JOHN W. HARBAUGH  
1959

STATE GEOLOGICAL SURVEY OF KANSAS

BULLETIN 134, PART 8 PLATE 3

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92  
113  
A1  
No. 134  
pt. 8  
#343  
plates



## EXPLANATION

- STANTON LIMESTONE  
CAPTAIN CREEK LIMESTONE  
MEMBER (lower part only)
- VLAS SHALE
- Calcareite subdivision
- Crystalline limestone  
subdivision
- Fragment-pellet limestone  
subdivision
- HICKORY CREEK MEMBER  
(shale with limestone beds)
- MERRIAM LIMESTONE  
MEMBER
- LANE-BONNER SPRINGS  
SHALE

Vertical lines on some measured sections indicate that  
bar in each measured section refers to location on map  
and not to distance. Symbols for Spring Hill Limestone  
Member in geological and historical maps have been  
used where appropriate.

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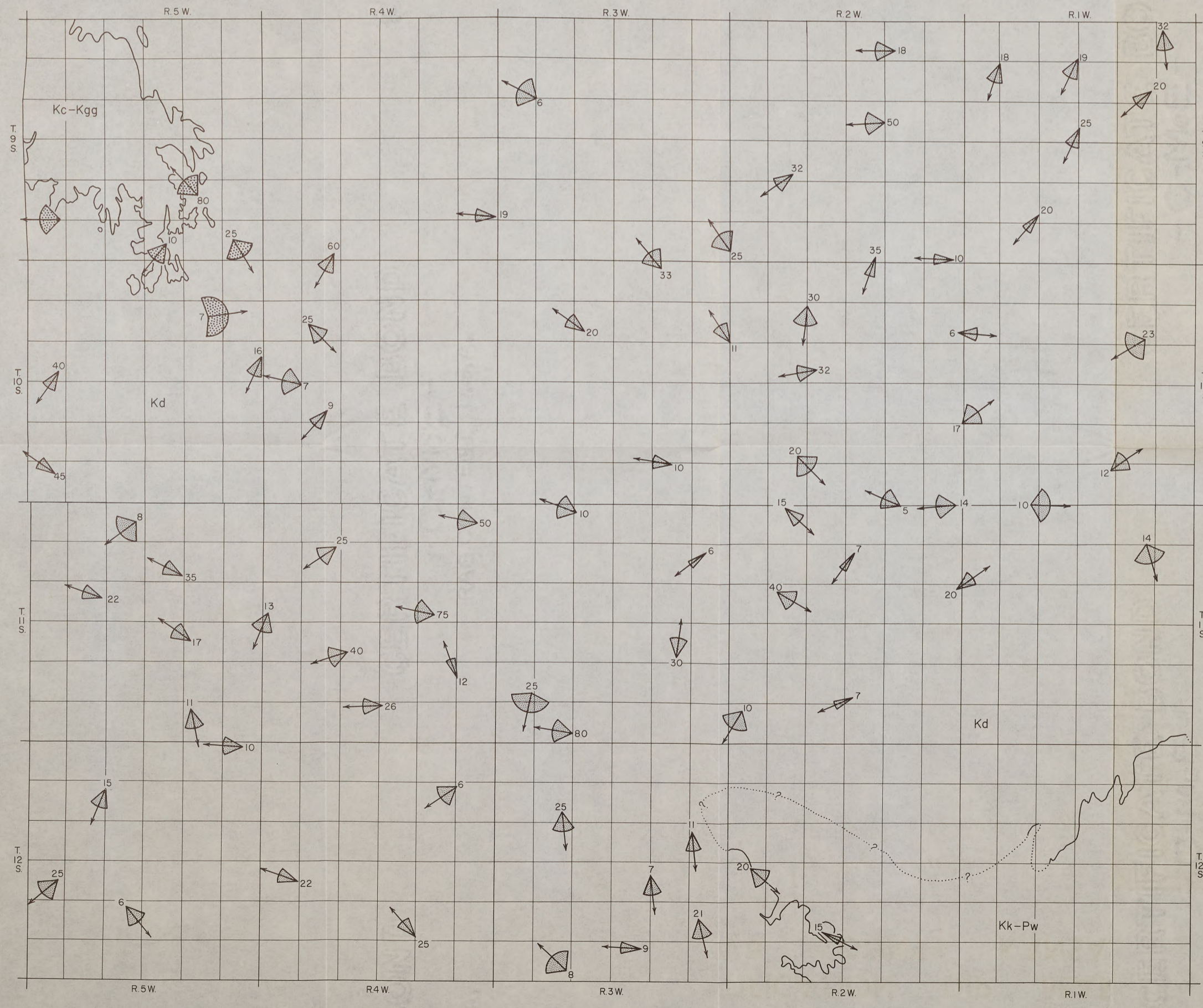
UCSB LIBRARY QE  
113  
A1  
No. 134  
pt 6  
#181 plate

# VECTOR RESULTANTS OF CROSS-STRATIFICATION DIP BEARINGS IN DAKOTA SANDSTONE, OTTAWA COUNTY, KANSAS

State Geological Survey of Kansas

By Paul C. Franks, George L. Coleman, Norman Plummer, W. Kenneth Hamblin, 1959

Bulletin 134 Part 6, Plate 1

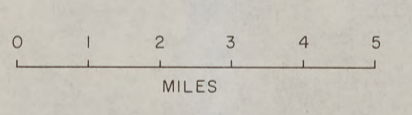


### EXPLANATION

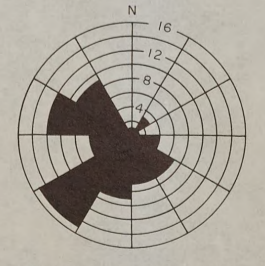
Vector resultant of dip bearings of cross-stratification. Arc shows one standard deviation in degrees plotted on each side of vector resultant. Numerical gives number of measurements. Coarse dots, Janssen Clay member of Dakota Formation; small dots, Terra Cotta Clay member of Dakota Formation.

Kc-Kgg Undifferentiated Carlile, Greenhorn, and Graneros formations  
Kd Dakota Formation  
Kk-Pw Undifferentiated Kiowa Shale and Wellington Formation

Contact  
(dotted and queried where concealed)



TRUE NORTH  
MAGNETIC NORTH



CIRCULAR HISTOGRAM  
showing vector resultant dip bearings plotted in 30° class intervals. Outside circumference represents 16 vector resultants.

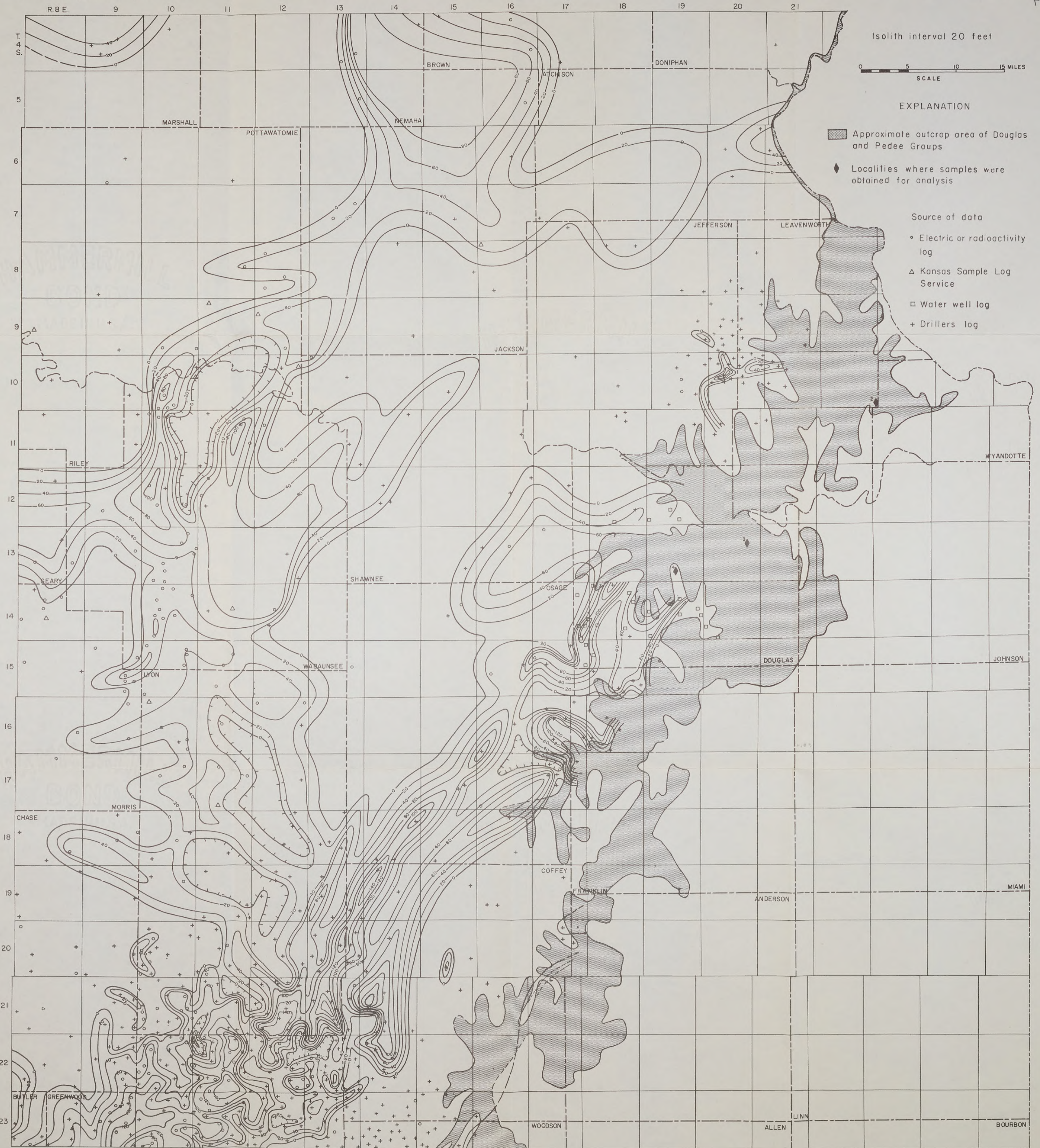
Geologic contacts adapted from Mack, report in preparation.

# NET THICKNESS OF SANDSTONE BELOW HASKELL LIMESTONE

State Geological Survey of Kansas

by Donald T. Sanders, 1957

Bulletin 134 Part 3, Plate 2



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# NET THICKNESS OF SANDSTONE ABOVE HASKELL LIMESTONE

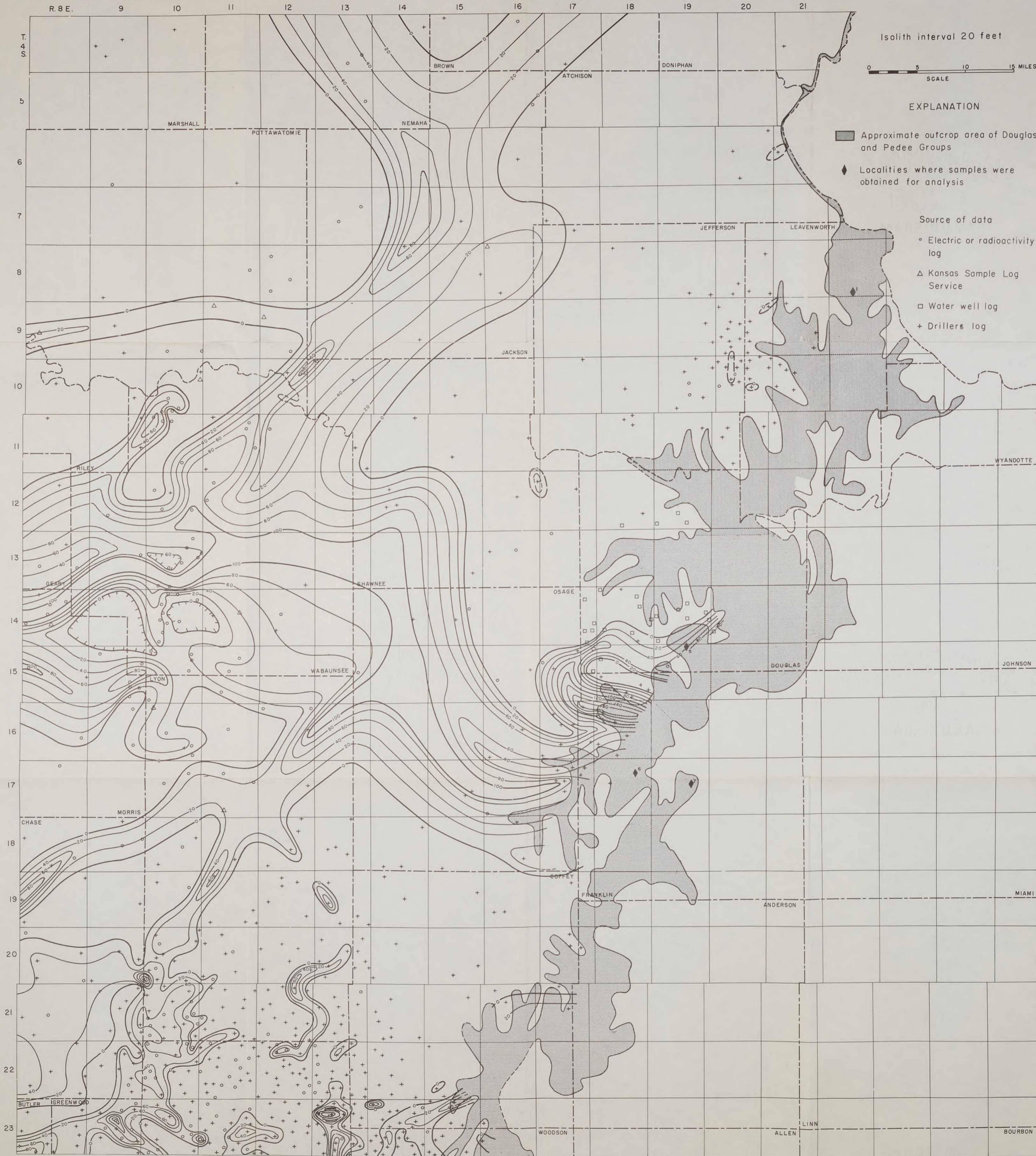
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113  
A1  
no. 134  
pt. 3  
#102  
plates

State Geological Survey of Kansas

by Donald T. Sanders, 1957

Bulletin 134 Part 3, Plate



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QE  
113  
A1  
no. 134  
pt. 5  
#101 plate

# Map of Bourbon and Crawford Counties, Kansas

Showing location of Mulky coal mines, reserve coal  
areas, and trace of the basal Fort Scott limestone

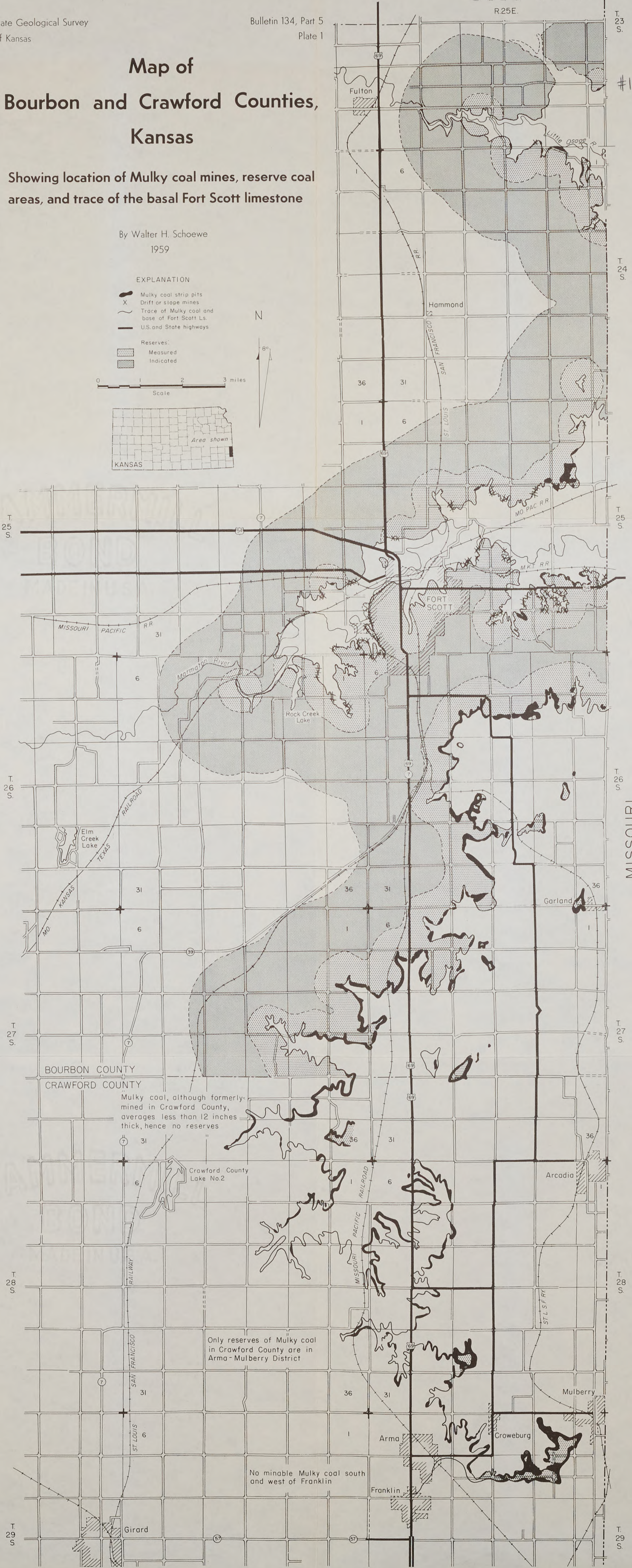
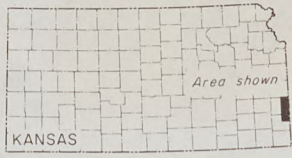
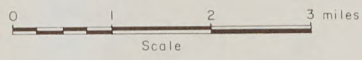
By Walter H. Schoewe  
1959

### EXPLANATION

- Mulky coal strip pits
- Drift or slope mines
- Trace of Mulky coal and base of Fort Scott Ls.
- U.S. and State highways

### Reserves:

- Measured
- Indicated



BOURBON COUNTY  
CRAWFORD COUNTY

Mulky coal, although formerly mined in Crawford County, averages less than 12 inches thick, hence no reserves

Crawford County Lake No. 2

Only reserves of Mulky coal in Crawford County are in Arma-Mulberry District

No minable Mulky coal south and west of Franklin

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QE 113 A1 no. 133 #44 plates

State Geological Survey of Kansas

DATA ON SECONDARY RECOVERY PROJECTS IN KANSAS 1957

Bulletin 133, Table 1

Table with columns: No., Field, Operator, Project, Location, Commenced or unitized, Year started, Total developed acres, Production formation, Thickness of producing zone, Average depth of zone, Flowing, Pumping, Total, No. producing wells, No. producing wells in 1957, No. injection wells, No. injection wells drilled in 1957, Medium of injection, Source of water, Average production per well, Cumulative secondary oil recovery, Production attributable to secondary recovery, No.

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Table with columns: No., Field, Operator, Project, Location, Cooperative on unit, Year started, Total developed acres, Production formation, Millions of cubic feet, Average depth, Average number of wells, Active wells (Flowing, Pumping, Total), No. producing in 1957, No. active injection wells, No. injection wells drilled in 1957, Method of injection, Source of water, Average daily production, Cumulative production, Production to secondary recovery, Production in 1957, No.







\* Production not to be revealed; included in county total.

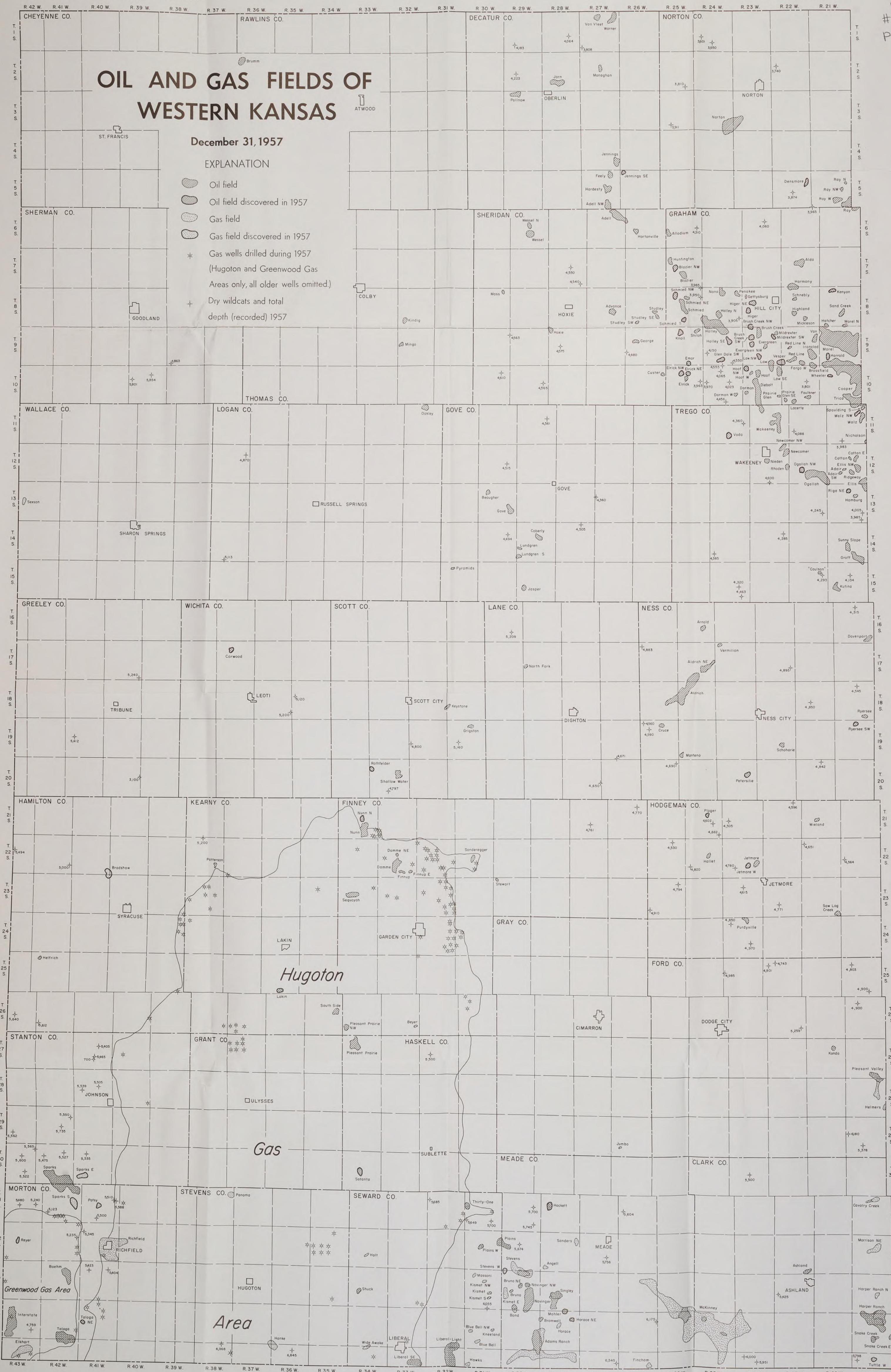
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# OIL AND GAS FIELDS OF WESTERN KANSAS

December 31, 1957

## EXPLANATION

-  Oil field
-  Oil field discovered in 1957
-  Gas field
-  Gas field discovered in 1957
-  Gas wells drilled during 1957 (Hugoton and Greenwood Gas Areas only, all older wells omitted.)
-  Dry wildcats and total depth (recorded) 1957








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QE 113 A1 no. 133 #204 plates

# OIL AND GAS FIELDS OF CENTRAL KANSAS

December 31, 1957

## EXPLANATION

-  Oil field
-  Oil field discovered in 1957
-  Gas field
-  Gas field discovered in 1957
-  Dry wildcats and total depth (recorded) 1957

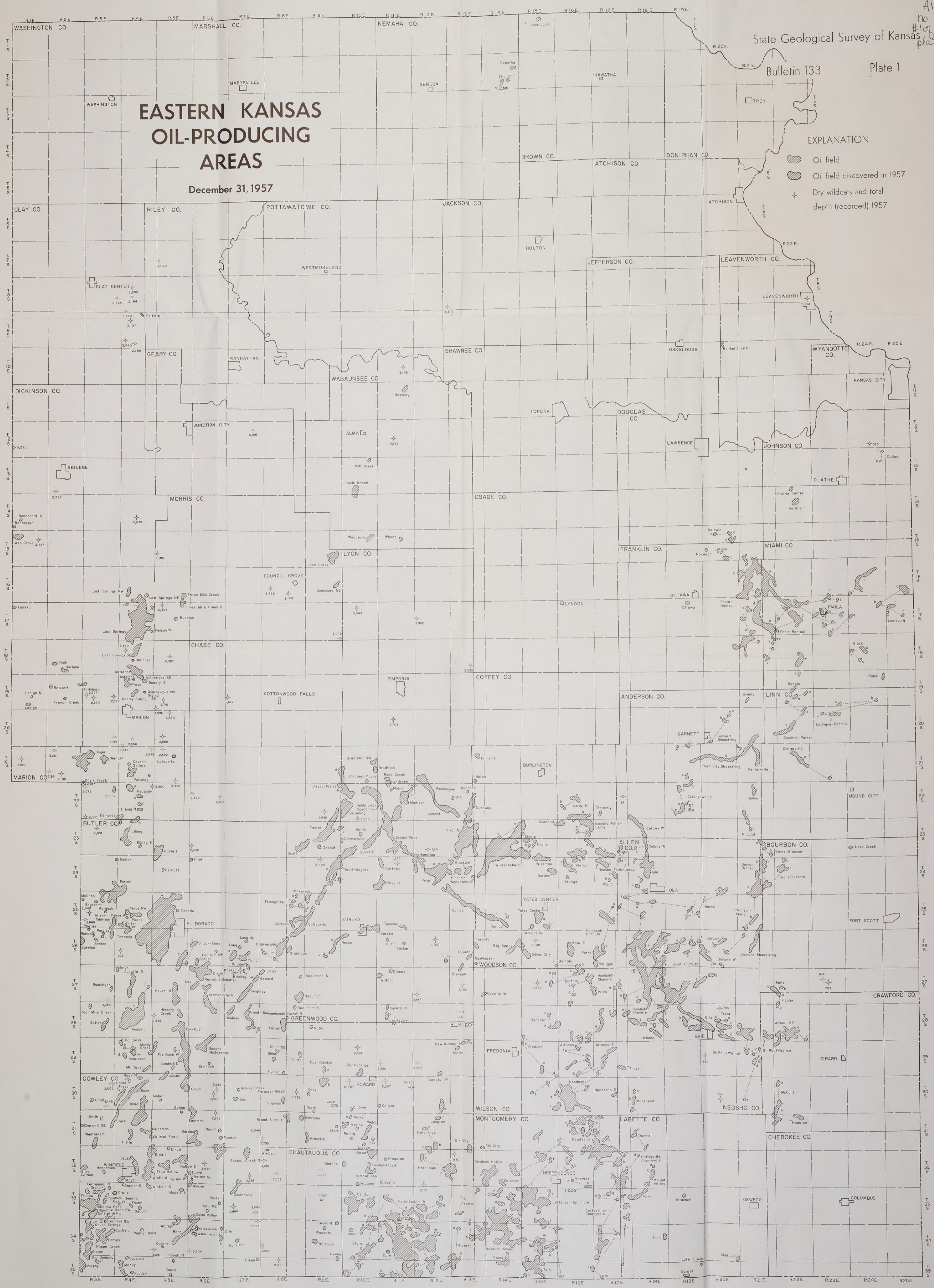


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# EASTERN KANSAS OIL-PRODUCING AREAS

December 31, 1957

EXPLANATION  
Oil field  
Oil field discovered in 1957  
Dry wildcats and total depth (recorded) 1957



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