# Geology and Ground-Water<br>Resources of Trego County, Kansas

By Warren G. Hodson



STATE **GEOLOGICAL** SURVEY KANSAS

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# Geology and Ground-Water Resources of Trego County, Kansas

By Warren G. Hodson

Prepared by the United States Geological Survey<br>and the State Geological Survey of Kansas with<br>the cooperation of the Environmental Health<br>Services of the Kansas State Department of<br>Health, and the Division of Water Resour

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(in pocket ) sas

#### FIGURE



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# Geology and Ground-Water Resources of Trego County, Kansas

#### ABSTRACT

Trego County is a 900-square-mile area in the High Dy II<br>Plans section of west-central Kansas. The altitude ranges sion, 0-square-mile area in the High  $t_{\text{max}}$  about 2,000 to 2,600 feet above sea level. the annual zero to zhood feel above the definition of the rarm. The purpose of this investigation was  $\frac{1}{2}$  $\pi$ , net propulation of the County in 1960 was 5,473. Farming and livestock raising are the chief occupations.

The Carlile Shale of late Cretaceous age is the oldest  $f(x)$  formation cropping out in the County. EXA formation cropping out in the County. The Carlile acter, Neogene (Pliocene) are unconformably overlies Creta. This study was made n the County. The Ogallala Formation of  $\alpha$  is rocks in much of the rocks in much of the County and consists principally erative ground-wate.<br>
Raviatile deposits of sand, gravel, and silt. Terrace the State Geological **Krosits of Pleistocene age occur along the principal val-** HS Geological Survey in cooperation with the es. Folian silts that mantle the uplands and alluvium U.S. Geological Survey,  $\alpha$  stream valleys constitute the are late Pleistocene <mark>in age.</mark>

The Ogallala Formation is frand vields water to many domestic and stock wells. Alluvium and terrace deposits yield moderate quantities If water, but these deposits are not of wide areal extent. a areas of Cretaceous outcrops, small quantities of ground water are obtained from the Dakota Formation of early ( ?)soae Member of the Carlile Shale of late Cretaceous age .

Hydrologic data are given in tables and include rec- Trego Count rds of  $280$  wells, logs of 176 test holes and wells, chemical analyses of 59 samples of water from wells .

# PURPOSE AND SCOPE OF INVESTIGATION

This report gives the results of a study of geology and ground-water resources of Trego Gounty, Kansas. Ground water, taken for granted, is one of the most vital natural resources of Kansas. water and for information concerning its ability has greatly increased in recent years be-

cause of an increasing population, accompanied ly municipal, industrial, and irrigation expanand new water uses in the home and on  $\frac{1}{140}$  the farm. The purpose of o determine the occurrence, availability, move ment, and chemical quality of ground water, and o relate the occurrence of the water to the char thickness, and distribution of the geologic formations.

as apart of the coop program begun in 1937 by I<sup>errace</sup> the State Geological Survey of in cooperation with the youngest deposits and Environmental Health Services of the Kansa State Department of Health and the Division of the most widespread aqui-<br>domestic and steek walls of the program is shown in Figure 1.

#### LOCATION AND EXTENT OF AREA

is in west-central Kansas, in and the third tier of counties south of the Nebrask border and in the fourth tier of counties east of the Colorado border (Fig. 1). It is almost square, extending approximately 30 )**INTRODUCTION** square, extending approximately 50 miles east<br>west and 30 miles north-south and has an area of about  $900$  square miles. It contains 25 town from T 11 S to T 15 S, and from R 21 W to <sup>R</sup><sup>25</sup> W.

#### PREVIOUS INVESTIGATIONS

although often Cretaceous rocks of Kansas have been <sup>a</sup> sub of study since the railway explorations and The demand for the Meek and Hayden surveys of the 1860's. avail- Adams (1898) has given an historical summar of the early studies of the Cretaceous rocks in

Kansas. Haworth contributed reports on the physiography ( 1897a ) , the physical properties of Tertiary rocks (1897b), and the geology of underground waters (1897c) in western Kansas. Logan (1897) discussed Upper Cretaceous rocks in western Kansas, and Williston (1897a) described the Niobrara Chalk and discussed the Pleistocene deposits of Kansas (1897b). Johnson (1901, 1902) discussed the Tertiary rocks of the High Plains and ground water in western Kan-High Plains and ground water in western Kan<br>sas. Darton (1905) discussed the geology and

ground-water resources of the central Grea Plains, in which he treated the loess of western **Kansas as a separate formation from the under** lying Tertiary deposits. In a special report on well waters in Kansas, Haworth (1913) dis cussed the Tertiary deposits of western Kansa and their water-bearing characteristics. Russe (1929 ) studied the stratigraphy and structure of the Smoky Hill Chalk Member in Logan, Gove, and Trego counties and reported on the useful ness of bentonite layers in interpreting the strati



FIGURE 1. Map of Kansas showing area discussed in this report, and other areas for which ground-water reports have been published or are in preparation,



graphy. Elias (1931) studied the geology of western Kansas with emphasis on Shale, the Ogallala Formation, and the late Samples Pleistocene deposits. Landes and Keroher leum resources of Logan, Gove, and Trego coun- gineering Laborator 1939 ) briefly discussed the ties. A report by Byrne, *et al.* (1947) describes ment the occurrence of construction materials in the Cedar Bluff area in Trego County. Leonard (1952) discussed the Pleistocene geolbgy of Kansas. Ground-water studies in which border Trego County, or include part of the County, were made by Prescott (1955), Hodson and Wahl (1960), and Leonard and Order: Berry (1961). Areas in Nansas in ground-water studies have been made, which are published or are in preparation, shown in Figure <sup>1</sup>

The writer spent about three months during the summer and fall of 1959 and about three the same months during the summer and fall of 1960 in the field gathering data upon which this report they were inventoried. is largely based . The areal geology was mapped from field observations and from stereoscopic ACKNOWLEDGMENTS study of areal photographs obtained from the  $U.S.$  Department of Agriculture. County maps prepared by the State Highway Commission of Kansas at a scale of one inch to used to record field data.

Data on 280 wells are given in Table 6 and include the depth to water and the depth of the well . Measurements were made with <sup>a</sup> well. Measurements were made with a steel tape who provided data concerning city water supgraduated to hundredths of afoot. Measure n a few wells could not be made and ments in a few wents could not be made and were obtained from the owner or driller. Information concerning yield, adequacy of supply, and quality of water was obtained, if possible, trom well owners. Drillers logs of wells and test holes were obtained, it available, from well owners and well drillers .

Logs of 176 test holes and wells are given. Included are 145 sample logs of test holes, 19 logs of test holes obtained from the Bureau of Reclamation, and 4 logs obtained from the State Highway Department . Well drillers and well owners provided logs of 8wells and test holes .

 Locations of wells and test holes within the sections were determined by means of an odo meter and from aerial photographs. The altitudes of measuring points of wells and test holes were determined with <sup>a</sup> plane table and alidade . The base map used for Plate <sup>1</sup>

from maps prepared by the Soil Conservation the Pierre Service, U.S. Department of Agriculture.

of water from representative wells Landes and Keroher were collected and were analyzed by Howard A. geology and petro- Stoltenberg, Chief Chemist, in the Sanitary Enof the Kansas State Depart ment of Health.

# WELL-NUMBERING SYSTEM

are thous, quarter-quarter sections, and ro-acte tracture are designated  $a, b, c,$  or  $d$  in a counterclockwise Kansas The locations of wells, test holes, and locations part of features are designated in this report according of wells, test holes, and local o General Land Office surveys in the following township, range, section, quarter section, which quarter-quarter section, and quarter-quarter and quarter section ( 10 -acre tract ) The quarter sec quarter-quarter sections, and 10-acre tracts direction beginning in the north the northeast quarter sec tion. For example, well 11-24-25 act is in METHODS OF INVESTIGATION NW SW NE sec. 25, T 11 S, R 24 W (Fig. 2). If more than one well or test hole is located in )U-acre tract, the location numbers are tollowed by serial numbers in the order in which

Thanks and appreciation are expressed to the many residents who gave permission to in the mile were ventory their wells and who allowed access to their property for the study of rock exposures, o those who permitted aquiter tests to be made using their wells, and to the municipal officials



FIGURE 2.—Map of Trego County, Kansas, illustrating **was compiled** the well-numbering system used in this report,

plies. The Trego County office of the Soil Conservation Service, U.S. Department of ture, located at WaKeeney, gave helpful infor- steep and mation.

E. L. and Carrie Reavis and Melvin Franz , of the U.S. Geological Survey and the State Geological Survey of Kansas , gave assistance sides tend during field work. The illustrations were drafted by Judy Crissler and Larry Shelton.

The manuscript for this report has been reviewed by members of vey and the State Geological Survey of Kansas; y Robert V. Smrha, Chief Engineer and Harris the Smoky Hill Valle L. Mackey, Engineer, of the Division of Water surface Resources of the Kansas State Board of Agricul- leve ture; and by J. Lee Mayes, Chief Engineer, Bruce F. Latta, Geologist, of Health Services of the Kansas State Departmen of Health.

# TOPOGRAPHY AND DRAINAGE TOPOGRAPHY AND DRAINAGE

Most of Trego County lies within the High Pediment slopes are characteristic Plains section of the Great Plains physiographic outer valleys province (Fig. 3) as designated by (1949, p. 276). The southeastern corner of County is in the Blue Hills division of sected High Plains section. Gently rolling uplands that <mark>are moderately dissected by s</mark>malle

ie County. Agricul- Near the major streams, the slopes are quite ne topography is rough. Slopes along the north sides of the valleys of the larger streams are relatively more gentle and longer than along the south sides, whereas slopes along the southern to be cut into steep canyons and bad I'm Fort Hays Limestone Member of the Niobrara Chalk forms prominent escarp ments in the southeastern part of the Count the U.S. Geological Sur- and along the Smoky Hill Valley. One of the more prominent physiographic features along is a high terrace. The of the terrace is about 50 feet above the of the flood plain and about 200 feet below and the general level of the upland. Loess and slope the Environmental wash obscure the outer margin of the terrace in most places.

**Three perennial streams cross the County, all** flowing in an easterly direction , and are part of the Kansas River drainage system. The Smok GEOGRAPHY Hill River flows across the southern part of the County, Big Creek across the central part, and the Saline River along the northern edge of

> of the of the major streams . Along the Schoewe Smoky Hill River valley, long pediment slopes the mask the prominence of the Fort Hays escar the Dis- ment and grade into the high terrace . Along the Gently rolling up- Saline River Valley, the pediments are developed n the soft Cretac<mark>eous</mark> chalk bed<mark>s a</mark>nd ble<mark>nd</mark> int<sup>1</sup>



FIGURE 3.—Map of Kansas showing physiographic regions. (From Schoewe, 1949, fig. 22.)



te terrace surface so imperceptibly that the upland edge of the terrace cannot be mapped pre cisely without subsurface information.

I he general slope of the land surface in Trego County is eastward. The total topograph ic relief is about our reet. The lowest elevation, about  $2,000$  feet above sea level, is in the chanel of the Smoky Hill River on the easter county line in the NE sec .1, T 15 S, R21W. The highest elevations are in the uplands in e west-central and northwestern parts of the County and are about 2,600 feet above sea level.

# CLIMATE

I rego County has a semiarid continenta climate. Low to moderate precipitation, a high rate of evaporation, reasonably mild winters, and fairly hot summers are typical. The hot FIGURE weather in summer is moderated by brisk winds (1931-1955) and low humidity. As a rule the winters are moderate with only short periods of severe cold weather and relatively little snowfall . Apercentage of the winter days are clear, snow flurries are common, and occasionally the communities are Collyer, area is subjected to blizzard conditions . In I rego County, the amount of precipitation and ts seasonal distribution are the controlling factors in crop growth. Rainfall is erratic, coming ometimes as storms of 4 inches or more, at other times, in no appreciable amount for several weeks .

I'm climatic data in this report were compiled sedimentary trom published records of the U.S. Weather Cretaceou Bureau (U.S. Department of Commerce ) .Approximately 70 percent of the annual precipitation in Trego County falls during the growing Figure season of about  $5\frac{1}{2}$  months (Fig. 4) . annual precipitation at WaKeeney is 21.40 given inches (based on period of record 1931-55 ) The greatest annual precipitation recorded was  $37$  inches in 1961; the least was 11.15 County. inches in 1888. The annual precipitation for Hill Valley east e period of record and the cumulative deearture from normal precipitation are shown in Figure 5. The average length of the growing County. eason is 167 days, the average date of alling frost in spring being April 27, and the everage date of the first killing frost in <sup>1</sup>eing October II.

#### POPULATION

According to the U.S. Census Bureau, Trego and County had a population of 5,473 in 1960, of \*hich approximately 50 percent was urban and  $\alpha$  percent rural. The County had an average of



at WaKcene

6.1 persons per square mile as compared to alarge State average of 26.6 persons. WaKeeney, the but county seat, had a population of  $2,808$ . Other population 233, and Ogalian, unincorporated.

# GENERAL GEOLOGY

# SUMMARY OF STRATIGRAPHY<sup>1</sup>

The rocks that crop out in Trego County are in origin and range in age from o Recent. I neir areal distribution is shown on Plate <sup>1</sup> ; their stratigraphic relation is illustrated by cross sections on Plate 2 and in 6. Ageneralized section of the rock The mean units and their water-bearing properties are  $\sum_{n=1}^{\infty}$ 

 The Carlile Shale of late Cretaceous age is in Trego County. The Carlile crops out along the Smoky of the junction of the Smoky Hill River and Hackberry Creek and is buried beneath rocks of younger age in the rest of the

the last **The Niobrara Chalk of late Cretaceous age** overlies the Carlile Shale. The Niobrara crops fall out along the outer valleys of the Smoky Hill and Saline rivers and their tributaries and along Big Creek valley in the eastern part of the The Niobrara Chalk is divided into two members, the Fort Hays Limestone below ie Smoky Hill Chalk above.



in this report follow that of the StateGeolorical Survey of kinsas the U.S. Geological Survey. <sup>1</sup> The classification and nomenclature of rock units described and differ somewhat from the classification and nemenclature of

ous rocks in much of the upland areas of the County. Erosion during the Pleistocene Epoch valleys. Smoky Hill valleys and in the eastern part of the County along Big Creek valley. Along the formation Smoky Hill Valley, the Ogallala has been removed along a belt as much as wide, exposing Cretaceous chaik and shale.

Unconsolidated deposits of and eolian origin represent the Pleistocene Series. Fluviatile deposits fill

The Ogallala Formation of Neogene age Cretaceous bedrock along the principal streams. (Pliocene) unconformably overlies the Cretace-<br>
ous rocks in much of the upland areas of the<br>
places overlie older terrace deposits along stream Deposits belonging to the Pleistocer has removed the Ogallala along the Saline and Series are the Grand Island and Sappa forma tions of Kansan age, the Crete and Loveland of Illinoisan age, the Peoria Forma tion of Wisconsinan age, and alluvium of Wisconsinan and Recent age.

> In this report the Grand Island, Sappa, and both fluviatile Crete formations are shown in the illustrations is terrace deposits, and the Peori and Loveland formations are discussed together

Table 1.—Geologic formations in Trego County, Kansas, and their water-bearing properties.

System	Series		Stratigraphic unit	Maximum thickness, feet	Physical character	Water supply
Neogene		Alluvium (incl. Recent and late <b>Wisconsinan</b> terrace deposits) Peoria and Loveland formations		70	Stream-deposited silt, sand, and gravel; thick, coarse deposits in major valleys, finer deposits in smaller valleys	Yields moderate to large quantities of water to wells along major val- leys; lesser amounts along smaller valleys
	Pleistocene			30	Silt, mostly colian, sandy in lower part. Mantles most of uplands and masks much of valley walls	Yields small amounts of water to wells locally
		Grand Island, Sappa, and Crete formations (Terrace deposits)		100	Stream-deposited sand, gravel, and silt in a terrace position along major valleys. Locally contains Pearlette ash bed	Yields small to moderate quantities of water to wells along major valleys
	Pliocene	Ogallala Formation		150	Fluviatile deposits of sand, grav- el, silt, and clay; mostly uncon- solidated, but cemented at places to various degrees	Yields moderate quanti- ties of water to wells in much of the County
Cretaceous	<b>Upper Cretaceous</b>	Niobrara Chalk	Smoky Hill Chalk Member	500	Chalk and chalky shale, thin- bedded and platy; bentonite beds throughout. Light gray to dark gray when fresh; weathers to brown and orange at outcrop	Yields no water to wells
			Fort Havs Limestone Member	55	Massive chalk beds separated by thin chalky shale partings, gray- ish-white, but may be yellow or light brown at outcrop	Yields small amounts of water to wells locally
		Carlile Shale		300	Lower part consists of gray cal- careous shale and thin beds of chalky limestone: upper part consists of dark-gray clayey shale with silty fine-grained sandstone lenses in upper part	Yields small amounts of water to wells from Co- dell Sandstone Member
			Greenhorn Limestone	100	Alternating beds of chalky limestone and chalky shale	Yields no water to wells
			Graneros Shale	45	Noncalcareous dark-gray shale; persistent bentonite bed in upper part	D <sub>0</sub>
	Cretaceous Lower(?)	Dakota Formation		300	Clav. shale, siltstone, and sand- stone; interbedded and vari- colored. - Sandstone fine- to medium-grained and lenticular. Lignite and "ironstone" com- mon	Yields small amounts of water to wells

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as loess deposits. Figure <sup>6</sup> illustrates the graphic relation of Pleistocene deposits across the surface the principal valleys. Deposits penetrated in test holes 13-21-5bbc and 13-21-8abb , along the north side of Big Creek valley, may be of Nebraskan age.

# UNEXPOSED ROCKS

Geologic rock units below the Carlile Shale the ocean, that do not crop out in Trego County but occur<br>weretaking in the subsurface within a possible drilling depth vegetation or water wells consist of approximately 500 teet of Greenhorn Limestone, Graneros Shale, and Dakota Formation. The Greenhorn Lime- underlying strata until stone and Graneros Shale are not considered to e aquiters. The Dakota Formation is an portant source of ground water for domestic and **fultration** and transpiration stock purposes in parts of Trego County. formations will be discussed further in the sec tion on geologic formations and their water bearing properties.

# GROUND WATER

Figure <sup>7</sup> is adiagram of the hydrologic cycle adapted from Piper (<sup>1953</sup> ) showing the part of

strati- ground water in the circulation of water near of the earth. The earth's water is constantly circulating from air to ground to air in a process called the hydrologic cycle. Wate falls to the ground in the form of precipitation such as rain, snow, sleet,  $cr$  hail. Part of this water vaporizes and returns to the air by evapa part is removed as runoff via rivers to a part percolates into the ground by some of which is used by trees and d is returned to the air through a process called transpiration, and a very small part percolates downward through the soil and it reaches the water table , where it becomes ground water . The hydro im-<br>filtration and transpiration is continuous. It is of rainfail, evaporation, runoff, in-These this cycle which provides the source of our ground water. Figure 8 shows the generalize divisions of subsurface water.

Where a water-bearing formation is confined between relatively impermeable beds and water is supplied to it from an adjacent area of ECONOMIC GEOLOGY **and SECONOMIC GEOLOGY** altitude, the water table is absent and the water is said to be confined or under artesian pressure . When an aquifer under artesian pressure is SOURCE penetrated by a well, water will rise in the well o a height equal to the hydraulic head. The imaginary surface connecting this level in wells





is called the *piezometric surface*, which may or may not be above the land surface. Ground water in the Dakota Formation in Trego Coun ty is under artesian pressure, the water rising in some wells as much as 300 to 400 feet above the top of the Dakota. Water in the Codell Sandstone Member rises nearly 200 feet above the top of the Codell in some of the deeper wells in the County. Ground water in the Dakota and in the Codell is discussed further in the section on geologic formations and their water -bearing properties.

THE WATER TABLE AND

MOVEMENT OF GROUND WATER

The water table fluctuates up and down in response to additions to or withdrawals from an aquifer. Fluctuations of the water table can be determined by periodic measurements of the depth to water in wells. The depth to water has been measured periodically since 1949 in several wells in the Smoky Hill Valley below the Cedar Bluff dam. These measurements have been published annually by the U.S. Geological Survey ( 1936-1959 ) and thereafter by the State Geologi



FIGURE  $6$  – Stratigraphic relation of Pleistocene deposits across Saline River  $(A-A')$ , Big Creek (B-B'), and Smok Hill River (C-C') valleys.

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al Survey of Kansas (1958-1963). n Figure 9 show fluctuations of evels below the Cedar Bluff dam. The hydro- Ground water graph of well 14-22-36aab , about <sup>a</sup> le dam, reflects seepage from the reservoir with ground water which was filled in 1951.

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The configuration of the water table in Ogallala Formation in the northern part of County is shown on Plate 1 by means of cor tour lines. The direction of ground-water move-water table ment is at right angles to the contours and in adownslope direction. The movement is general-<br>and from there into Big Creek (Pl. 1). ly eastward, but along the outer edges of upland the contours bend back reflecting dis-<br>charge along the edge of the upland. Local available. At the present anomalies in the water-table contours along the water level edge of the upland indicate areas of greater recharge due to the predominance of gravel near the land surface where the upland

Hydrographs loess has been removed by erosion , exposing the ground -water coarser clastic deposits of the Ogallala .

in the alluvial valleys of the mile below Smoky Hill and Saline rivers is not continuous in the Ogallala Formation in the uplands but is separated from it by imper the meable Cretaceous rocks . Along Big Creek in the the central and western parts of the County, however, the water table is continuous with the in the Ogallala , and water -table con tours show that ground water from the Ogallala is discharging into terrace deposits and alluvium  $g^{eneral}$  and from there into Big Creek (Pl. 1)

the No long-term record of observation wells in in Trego County is Local available. At the present rate of withdrawal, the is fairly stable and ground -water re charge and movement into the County is in equilibrium with discharge and movement from it .



FIGURE 7.—Diagram of the hydrologic cycle. Solid flow lines indicate movement of water as a liquid and broken lines movement as vapor. Heavy flow lines indicate man's principal interruption of the natural cycle (after Piper, 1953 ) .

FIGURE 8. Diagram showing generalized divisions of subsurface water (after Meinzer, 1923, fig. 2).

## ADDITIONS OF GROUND WATER SUBSURFACE INFLOW

Additions to the ground-water reservoir may The quantity occur in several ways: by direct infiltration of precipitation, by seepage of water from streams, by subsurface inflow of areas, and by infiltration of irrigation water applied to the land.

#### INFILTRATION OF PRECIPITATION

Approximately ou percent of nual precipitation in Trego County falls during centers the Count the months of climate is characterized by brisk wind move- Because ment, high temperatures, and relatively low valley deposits, humidity. Consequently, evaporation is rapid, and much of the precipitation is returned to the atmosphere. Because much of during the growing season, a of the precipitation is returned to phere through transpiration of plants.

I'm topography, vegetative cover, and na- areas. ture of the subsoil greatly affect the amount of water that will infiltrate below the land surface. The upland areas underlain by the Ogallala Irrigation Formation are conducive to infiltration of raintall in areas where sand and gravel crop out. Where thick loess deposits cover the Ogallala, infiltration is retarded .

The alluvium and terrace deposits along stream valleys are very conducive to because of the sandy nature of the deposits, low relief, and the shallow depth to water. These deposits, however, are of relatively small sfiltration areal extent. The annual quantity of recharge

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a is not known, but it is believed to be a small per centage of the total precipitation that falls in the County. Nevertheless, this recharge prob ably amounts to at least 10,000 acre-feet per year.

 $S$ moky Hill River, Saline River, and Big Creek are effluent streams in Trego County . Their channels are cut below the water table and thus generally are receiving water from, rathe than adding water to, the zone of saturation (Fig. 10). Deposits adjacent to the channels are recharged during flood stages, but the wate represents temporary bank storage and returns to the stream soon after the flood. Smalle tributaries in the County lie above the wate table and during periods of stream flow con tribute water to the underlying strata where the channel is underlain by deposits that are con ducive to minitation.

of subsurface inflow to the County is not known, but it is estimated that from 3 to 4 acre-feet of ground water per day water from adjacent moves into Trego County from the west through the Ogaliala Formation (1 acre-foot equal 325,850 gallons). A small amount of grounc water enters the County along the norther chiefly in valleys tributary to the the mean an-<br>
Saline River valley. Some ground water also by movement in alluvium and terrace deposits along the major stream valleys. of the small cross-sectional area of the the amount of ground-wate movement in them is relatively small.

#### INFILTRATION OF IRRIGATION WATER

considerable part . Recharge from water applied to the land for the atmos- irrigation has been estimated to approach 25 per cent of the applied water in certain irrigate In these areas, the irrigated water is distributed in relatively long ditches over soils and subsoils that are very permeable.

> in Trego County is confined chiefl o the Big Creek and Saline River valleys where the soils are sandy loam. The percentage of applied water that recharges the alluvium of these valleys is not known but probably is consider in small local areas . Because irrigation is not practiced extensively and because the areal the lextent of the alluvial valleys is relatively small, the amount of recharge to the aquifer from in of irrigation water is not significant .

In Trego County, much of the upland plair

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is thickly mantled by loess and has <sup>a</sup> relatively dense soil, and recharge probably does not exceed an average of 10 percent of the applied irrigation water. Because only a few irrigation

wells obtain water from the Ogallala in Trego County, the amount of recharge to the aquifer from infiltration of irrigation water is con sidered to be negligible.



FIGURE 9.—Hydrographs showing fluctuations of water levels in the Smoky Hill Valley, and graphs showing monthly and normal precipitation at WaKeeney,



FIGURE 10.-Diagrammat c sections showing inducnt and effluent streams.  $A$ , Longitudinal section showing how a stream gains water (right) and how it loses water (left).  $B$ , Transverse section across influent part of a stream.<br> $C$ , Transverse section across effluent part of a stream. (After Meinzer, 1923, fig. 26).

# **GROUND-WATER DISCHARGE**

Subtractions from the ground-water reservoir n Trego County include discharge by effluen seepage into streams, discharge of springs, spiration and evaporation, discharge from wells, and subsurface outflow.

#### SEEPS AND SPRINGS

Streamhow is manualned at e perennial streams in the County by water discharge into the streams, mainly as along stream channels. Discharge is chiefly from alfalfa,<br>allumed and terrore denosite but along Pix alluvial and terrace deposits, but along  $Big|$  salt cedar. Creek, in the western and central parts of County, the Ogaliala Formation is in with the valley deposits and discharge occurs from the Ogallala into the valley deposits. Small springs occur at places where the contact of the Ogallala Formation and underlying Cretaceous rocks crop out. Springs also occur along the edges of the high terrace deposits in the Smoky Hill Valley. The amount of discharge of water by springs and seeps in Trego County is not known, but it is believed to be considerable.

#### TRANSPIRATION AND EVAPORATION

Shallow ground water along the major streams supports a neavy growth of vegetation, and during the growing season <sup>a</sup> amount of ground water is transpired by plants. The quantity n this manner is difficult to determine, but it probably exceeds the discharge by all other means combined. Transpiration reduces the the County. ground-water discharge into a number of streams, and flow in these streams during the growing season is largely runoff from storms .

After the first killing frost in the fall,

nial streams show an appreciable increase in flow due to a decrease in plant transpiration. In the upland areas transpiration does not occur from the water table because of the greater depth of the water table. Soil moisture lost by tran spiration on the uplands, however, must be replaced before any water can infiltrate downward to the water table.

Where the water table is within a few feet of the land surface, ground water may evaporate from the capillary fringe overlying the zone of saturation, or, if the water table is within a few inches of the land surface, water may evaporate directly from the zone of saturation. The evaporation of water from the ground-water reservoir in the County is restricted to the alluvial valleys and the amount is probably small compared to other means of ground-water discharge.

#### **PHREATOPHYTES**

reservoir Phreatophytes are plants that characteristica ly grow where they can send their roots down to tran- the water table, or to the capillary fringe, and thus obtain their water supply. Most phreato phytes are of low economic value, and the wate they transpire to the atmosphere is defined as "consumptive waste." More than /0 plan low stages along species have been classified as ground-<br>
(Robinson, 1958, p. 1), some of the more com-<br>
as seens sceps<br>mon being pickleweed, rabbitbrush, salt grass, cottonwood, willow, greasewood, and Salt cedar (T*amarix pentandra)* is a the serious problem in many stream valleys in the contact southwestern and western United States. In salt cedar has developed a jungle-like Small growth, invading and choking overflow channel of streams , and producing serious flood hazards . or control of salt cedar is very Springs also occur along the difficult and expensive . The annual use of water by salt cedar in areas of heavy growth may be several acre-feet per acre.

> Salt cedar has invaded some stream valleys in Kansas and is now found along the Cimar ron and Arkansas valleys in southwestern Kan is, Smoky Hill and Saline valleys in west central Kansas, and South F<mark>ork and Nor</mark>th Fork significant Solomon valleys in north -central Kansas ( To these – manek and Ziegler, 1963).

of ground water discharged Salt cedar has been used extensively as a shelter belt in Trego County and is also found s an ornamental shrub in many farm yards in A number of places around the Cedar Bluff reservoir have become heavily in fested with salt cedar. Ecological studies of sal cedar in the Uedar Bluff area (Tomanek and peren- Legier, 1965) show that the plant benefits from

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a drought or conditions that keep the water level The coefficient of transmissibility (T) may falling or low for a long period of time.

mestic, stock, and irrigation use in Trego Couny. Municipal water departments in reported <sup>a</sup> total use of about 450 acre -feet of ground water in 1960. Records of the Division presse of Water Resources of of Agriculture show that 3,487 acre -feet of ground water per year were appropriated for irrigation in the County as amount of ground water reported as us of ground water reported as used for amount may vary considerably, depending upon 1 rainfall. Most rural residents obtain water for foot per foot, domestic and livestock use from small-diameter drilled wells equipped with cylinder pumps powered by windmills. The yields of these well are small and probably average about one gal lon a minute. Because of their great number and long pumping periods, however, such well represent an important withdrawal of water, estimated to be about 800 acre-feet per wells, vear for both household and stock purposes . Withdrawal of ground water from wells is increasing somewhat because of new water uses in the home and on t<mark>he farm. There is also a small</mark>htable annual increase in the number of n the County, mostly in the alluvial valleys.

Ground-water contours on Plate 1 show the subsurface flow of ground water to be in a erally eastward direction in the northern part formula, of the County. In most areas the westward along the edge of the uplands, cating discharge from the uplands toward the formulas and their application are discussed by stream valleys. The amount of outflow in alluvial valleys is small because of small cross-sectional area of the alluvial valleys. It is estimated that about 700 acre-feet of ground may water per year leaves Trego County by face outflow.

# HYDROLOGIC PROPERTIES OF

The quantity of ground water that an r will yield to wells depends upon the logic properties of the material forming the aquifer. The hydrologic properties of greatest deposits significance are the ability to transmit and to store water, which are measured by the coeffi- and have a cients of transmissibility and storage. Controlled characteristic aquifer tests in the field provide the data required to compute these coefficients.

be expressed as the rate of flow of water, in gallons per day, through <sup>a</sup> vertical strip <sup>1</sup> foot DISCHARGE FROM WELLS<br>Ground water is pumped for municipal, do-<br>Ground water is pumped for municipal, do-<br>rated thickness of the aquifer, under a hydraurated thickness of the aquifer, under a hydrauic gradient of 1 foot per foot, at the prevailin the County ground-water temperature.

The coefficient of storage (S) may be exs the change in the stored volume of the Kansas State Board water per unit surface area of the aquifer per unit change in the component of head normal o that surface.

of April, 1963. The The field coefficient of permeability (P) may e expressed as the rate of flow of water, in This gallons per day, through a cross-sectional area of square foot, under a hydraulic gradient of 1 it the prevailing ground-water The field coefficient of permeabil ty can be computed by dividing the coefficient of transmissibility by the aquifer thickness ( m)

#### DETERMINATIONS OF TRANSMISSIBILITY AND PERMEABILITY

Aquifer tests were made using two irrigation 12-23-20ccc , deriving water from alluvium in Big Creek valley, and 12-23-30acc, derivin water from terrace deposits along Big Creek valley. Details of the wells are given in the of well records at the end of this report. Observation wells for drawdown measurements were constructed at appropriate distances from the pumped wells. Figures  $11, 12, 13,$  and  $14$ of these tests. Values of trans-SUBSURFACE OUTFLOW give the results missibility were computed from the test data by gen- the methods generally referred to as the Thiem le Theis nonequilibrium formula, the contours bend Jacob modified nonequilibrium formula, and the indi- Theis recovery formula. The theory of these the Ferris, et al.  $(1962)$ .

the relatively The aquiter tests indicate that a coefficient of transmissibility of about 40,000 gpd per foot e expected in the alluvium at localitie subsur- where the saturated thickness is adequate and where reasonably good sand and gravel beds are in the saturated zone . Test drilling should dis close the more favorable well sites.

WATER-BEARING MATERIALS **Although the aquifer test** in the terrace deaqui- posits shows a coefficient of transmissibility of hydro- about 120,000 gpd per foot, the transmissibility is too high for most deposits shown as terrace in this report. Deposits in the vicinit of well 12-23-30acc are thicker, better sorted, greater saturated thickness than is of terrace deposits in most of the County. Also, in this area along the central and western reaches of Big Greek, terrace deposits

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are adjacent to, and in many places lithologically silt are characteristic of the upper part of the indistinguishable from, the Ogallala Formation. deposits and that sand and gravel generally oc-<br>The water table in the ter Ogallala Formation into the terrace deposits. When pumping occurs in the terrace deposits, there is increased movement of ground water<br>from the Ogallala into the terrace deposits to-<br>Data on 280 wells in Trego C ward the area of drawdown. In contrast, along the Saline River and Smoky Hill River valleys and along Big Creek valley in of the County, the terrace deposits are separated the time from the Ogallala by wide areas of impermeable Cretaceous shale.

I'm coefficients of storage calculated in aqui  $er$  test  $12-23-30$ acc (Fig. 13) show a change One from artesian to water-table conditions as ing continued. Logs of test holes in the alluvium needs. and terrace deposits show that silt

The water table in the terrace deposits is con-<br>sinusus with the water of the control of the control in the lower part. The silt beds cause a tinuous with the water table in the Ogallala slight artesian effect because of confinement of Formation and ground water moves from the ground water below these less permeable silt ground water below these less permeable silt beds.

Data on 280 wells in Trego County are given<br>in Table 6. Only part of the domestic and stock the Saline River and Smoky Hill River valleys wells were visited, but records were made of all the eastern part municipal and irrigation wells in the County at of this investigation. The principal uses of ground water in the County are listed below .

#### DOMESTIC AND STOCK SUPPLIES

of the chief uses of ground water in pump- Trego County is to supply domestic and stock Nearly all domestic and stock supplies in and clayey the rural part of Trego County are obtaine



FIGURE 11.—Drawdown of water levels in pumped well and observation wells during aquifer test, plotted agains time since pumping started (well 12-23-20ccc ) .

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rackie 12.—Recovery of water levels in pumped well and observation, wells, during aquifer, test, plotted, against time since pumping stopped (well 12-23-20ccc ) .



Figure 13.—Drawdown of water levels in observation wells during aquifer test at 90 and 1,415 minutes after pump ig started, plotted against distance from nummer well (12-23-20-cc).

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Original from UNIVERSITY OF MINNESOTA small watercourses. Most domestic and stock wells in the County are drilled wells in which casing has been set and which are equipped with Collyer. displacement-type pumps in is below the water level . Most pumps are oper- Ogallala Formation ated by windmills; the others are operated by electric or gasoline motors, or by hand.

#### MUNICIPAL SUPPLIES

At the time of this investigation only the communities tained municipal water -supply systems . Data re communities of WaKeeney and Collyer main- partment. garding individual wells and details of well con struction are given in Table 6.

WaKeeney.—I he city of WaKeeney obtains Count is water supply from seven drilled wells in the alluvial deposits of Big Creek valley about <sup>2</sup> miles south of the city. The wells are equipped batteries with electrically driven turbine pumps; yields range from about 100 to 500 gpm. A 1 well is also maintained at the east edge of which derives water from the Ogallala Formation. Storage is provided by 2

from wells, but in parts of the County where tanks with capacities of 250,000 and 75,000 gal-<br>ground water is difficult to obtain, some stock lons and 2 underground reservoirs with capaciground water is difficult to obtain, some stock lons and 2 underground reservoirs with capaci-<br>supplies are provided by dams constructed across ties of 200,000 and 100,000 gallons. Water used supplies are provided by dams constructed across ties of 200,000 and 100,000 gallons. Water used<br>small watercourses. Most domestic and stock in 1960 was about 136 million gallons, as reported by the WaKeeney Water Department.

> which the cylinder water supply from two drilled wells in the The city of Collyer obtains its at the southwest edge of the city. The wells are equipped with electrical ly driven turbine pumps; yields are about 100 gpm. Storage is provided by an elevated 55,000 gallon tank. An annual use of about 10 million gallons was reported by the Collyer Water De

#### IRRIGATION SUPPLIES

There were  $22$  irrigation plants in Trego by the fall of 1960 ,most of which were in the alluvial valleys. Nearly all were single wells, but a few plants were pumping from of two or more wells . Yields of the irri yields gation wells ranged from about 150 gpm to reserve about 1,000 gpm. According to records of the town Division of Water Resources of the Kansas State of Agriculture, 1,725 acres in Trego elevated storage County were covered by ground -water rights , or



Figure 14.—Drawdown of water levels in observation wells during aquifer test, plotted against time sin<mark>ce pum</mark>pin started (well 12-23-30acc ) .

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applications for rights, for which 3,487 acre-feet of ground water per year was appropriated, as of April 1963. Surface-water rights, or tion for rights, totaled 1,105 acres, 2.119 acre-feet of surface water per year was struction materials, appropriated. This does not include water released from Cedar Bluff reservoir for irrigation districts below the reservoir.

#### POSSIBILITIES OF DEVELOPING SEC. ADDITIONAL SUPPLIES

Most of the land irrigated with ground water pool,<br>From County is in Big Crook valley. On the pool, in Trego County is in Big Creek valley. assumption that the alluvium and Wisconsinan terrace deposits of Big Creek valley have an average width of a half mile, a length of 50 miles, an average thickness of satur mies, an average incriness of saturated material County,<br>of 20 feet, and a specific yield (ratio of the duction volume of water the saturated part of r will yield, by gravity, to its own volume) of percent, there would be feet of water in storage, which is the amount of water pumped for irrigation in no new pools were discovered<br>Big Creek valley in 1962 (as reported to the until 1941, when the Ogallah pool was opened<br>Division of Water Becauses of the Kenses State 1962 ( as reported to Division of Water Resources of Board of Agriculture ) If pumping were in creased enough to lower the water table below stream level, the ground-water reservoir would be recharged from the stream instead of charging into the stream as it does now. It is estimated that about equal amounts of water are in storage in each of  $\frac{1}{10}$  valleys in the County—Smoky Hill Valley, Saline Valley, and Big Creek valley.

About 300 square miles in the northern and central parts of Trego County are underlain by the Ogallala Formation in which the estimated cal Survey average saturated thickness is about 25 feet . As  $suming$  a specific yield of 15 percent, the ground water 700,000 acre -feet . Not all this water could be In storage in this area would be about fields, pumped from the aquiter, and yields from wells  $\frac{1}{\text{Year}}$  Programs Production, are only moderate to low, but it is an important ducing thousands ducing ducing thousands ducing thousands ducing thousands ducing the state of barrels. source of water in the area.

#### **CEDAR BLUFF RESERVOIR**

Cedar Bluff Reservoir on the Smoky Hill <sup>1941</sup> <sup>9</sup> River was filled in 1951 soon after construction  $\frac{1942}{1043}$   $\frac{8}{6}$ of the dam was completed. The height of the m is 134 feet above the river; its length is  $12,560$  feet. The reservoir is approximately 9 miles long and nearly 5 miles wide at its widest  $\frac{1948}{1948}$ point. It has a shoreline of about 54 miles at irrigation-pool level. The reservoir has a floodstorage capacity of 191,800 acre-feet and an irrigation -storage capacity of 176,800 acre -feet .

#### MINERAL RESOURCES

applica-<br>which than soil and ground water, include oil, conof Trego County, other for which than soil and ground water, includ<br>year was struction materials, and volcanic ash.

#### OIL

I rego County has been explored for oil since 1925. Oil was first discovered in May, 1929, in  $0, 113$  S, R  $21$  W. Production was from rate of Pennsylvanian age at a depth of  $3,960$  to  $3,972$  feet. Named the Rega ool, it was soon abandoned after producin On the about 15,000 barrels of oil. The pool was revived with a producer in 1947, but was abandoned again in 1950. The WaKeeney pool, discovered in 1934 in the northern part of the was the second pool discovered. Pro the duction was from the Lansing Group of Penn the aqui-<br>lume) of sylvanian age. The Gugler pool was discovered In 1930 about 4 miles northeast of the Rega pool. about 30,000 acre-<br>ich is 1.500 times 1,<sup>500</sup> times age and from Arbuckle rocks of Cambrian and Ordovician age. No new pools were discovered the Kansas State about <sup>4</sup>miles north of the Gugler pool . Further development of the oil industry in Trego Count  $\mathbf{r}$  is summarized in Table 2.

the ground-water reservoir would<br>trom the stream instead of discussion of the 1963 production from 331 wells in 29  $\frac{dis}{dis}$  fields was 1,582,000 barrels of oil. Cumulativ production to the end of 1963 was almost 19 ground million barrels. Gas in commercial quantitie the three prin-<br>Ly Hill Valley has not been produced in Trego County. Yearl data on production of oil and gas in Kansas are available in the files of the Oil and Gas Divisio of the State Geological Survey of Kansas and are published as bulletins of the State Geologi of Kansas .

Table 2. —Number of producing wells and producing and annual production of oil from 1938 to 1963 in Trego County, Kansas.

Year	Pro- ducing wells	Pro- ducing fields	Production, thousands of barrels	Year	Pro- ducing wells	Pro- ducing fields	Production. thousands of barrels
1938	13	2	$326*$	1951	51	10	264
1939	13	2	43	1952	110	14	802
1940	8	2	48	1953	139	20	1,032
1941	9	3	45	1954	160	23	1.028
1942	8	2	23	1955	182	25	1.077
1943	6		42	1956	243	29	1,470
1944	9	2	33	1957	278	26	1.928
1945	11	3	64	1958	276	29	1.877
1946	11	3	74	1959	287	31	1.781
1947	14	5	88	1960	294	31	1.584
1948	17 <sup>°</sup>	5	97	1961	293	30	1,537
1949	18	7	107	1962	295	32	1,590
1950	18	7	90	1963	331	29	1.582

• Cumulative production to the end of 1938.<br>Data from Bulletins of the State Geological Survey of Kansas, annual oil and gas development reports.



#### MIXED AGGREGATE

Deposits of sand and gravel suitable as road metal for light-traffic roads and in the construc tion of black -top roads are abundant in Trego County. The deposits are the Smoky Hill Valley as terrace deposits and in the Ogaliala Formation in the upland areas. Terrace deposits along the Smoky Hill Valley are predominantly sand and gravel and are more suitable for mixed aggregate than are terrace deposits along other streams in the County, which in places contain considerably more silt Chef Chemist, with little or no sand and gravel .

#### CONCRETE AGGREGATE

Aggregate for concrete should be free from per million adherent coatings or silt and clay particles that lents per million are given<br>analyses show only the dis would interfere with bonding. Arkosic sand and gravel in terrace deposits along the Smoky Hill stituents and Valley are generally suitable for concrete aggre gate, although if silicious materials such as line or chalcedonic chert is present in significant shown quantities, a low-alkali cement may be necessary.

#### STRUCTURAL STONE

The Fort Hays Limestone Member of the Niobrara Chalk has been quarried for stone at numerous localities in Trego County. Calcium C.<br>stone at numerous localities in Trego County. Magnesium M. I he Fort Hays is relatively soft, although it hardens upon weathering. The Fort Hays also  $P_{\text{Gauss}}$  and  $P_{\text{Gauss}}$  .  $(225, 235)$ tends to absorb water and thus to deteriorate  $\begin{array}{ccc}\n\text{Carbonate} & \text{CO}_3 & \text{CO}_3\n\end{array}$ through freeze-and-thaw action and from spalling. Many farm buildings, city dwellings, business houses in the area constructed of Fort Hays Limestone seem to stand up well for many years, however. In a report by Risser (1960 ) the sources and characteristics of (1960) the sources and characteristics of build-<br>ing stone in Kansas are discussed. The RELATION TO LISE

volcanic ash in Trego County consists pre- constituent dominantly of minute, platy or curved frag ments of volcanic glass. The ash is white to light pearly-gray, but occasionally may Kansas show tints of yellow or red The most common uses of volcanic ash are as an ingredient of DISSOLVED SOLIDS ceramic glazes, as an additive to certain types of concrete, and as surfaced highways.

Asmall pit of n the SE sec. 36, T 14 S, R 21 W. About  $10$ feet of ash is present, with the lower 6 tively free of impurities. The deposit becomes generally thinner laterally from the pit. Only a amount of ash has been taken from the pit

CONSTRUCTION MATERIALS Two other ash deposits were noted —one in the SE sec. 28, T 14 S, R 21 W, and another in the SW sec. 30, T 14 S, R 22 W. These deposits are small and have little commercial value.

# QUALITY OF GROUND WATER

I'm chemical character of the ground water in Trego County is indicated by analyses of samples from wells deriving water from the principal aquifers (Table <sup>3</sup> ) . The analyses of water were made by Howard A. Stoltenber<br>Chief Chemist, in the Sanitary Engineerin y Howard A. Stoltenberg, Laboratory of the Kansas State Department of Health. The results of the analyses are given in parts per million . Factors for converting parts of mineral constituents to equiva in Table 4. Inc. analyses show only the dissolved mineral conlo not indicate the sanitary con dition of the water. Representative analyses of ground water from the principal aquifers are in Figure 15 .

TABLE 4.-Factors for converting parts per million to equivalents per million.

Mineral constituent	Chemical symbol	Conversion factor 0.0499	
Calcium	$Ca^{**}$		
Magnesium	$Mg^+$	.0822	
Sodium	Na:	,0435	
Potassium	К÷	.0256	
Carbonate	CO <sub>3</sub>	.0333	
Bicarbonate	HCO <sub>3</sub>	.0164	
Sulfate	SO <sub>2</sub>	,0208	
Chloride	Cl	.0282	
Fluoride	F	.0526	
Nitrate	NO <sub>3</sub>	.0161	

# IN RELATION TO USE

VOLCANIC ASH The following discussion of the chemical of ground water has been adapted  $\sum$  part from publications of the U.S. Geological generally Survey and the State Geological Survey of

The residue that is left after a sample of a mineral filler for bituminous- water has evaporated consists mainly of the dissolved minerals in the original sample, but may also include some organic material and water of crystallization . Water containing less than feet rela- 500 ppm (parts per million) of dissolved solid is satisfactory for domestic and many small industrial purposes . Water containing more than 1,000 ppm of dissolved solids is likely to Generated at University of Kansas on 2023-10-04 19:52 GMT / https://hdl.handle.net/2027/umn.31951000881966a<br>Public Domain in the United States, Google-digitized / http://www.hathitrust.org/access\_use#pd-us-google



(Samples analyzed

by  $\equiv$ 

A.Stoltenberg.)



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of water.or 8.33 pounds per million gallons of water of substance per million pounds to one pound is equivalent t \*One part per million  $\bullet$  One  $\uparrow$ 

> Original from UNIVERSITY OF MINNESOTA

a noticeable taste or to make the water unsuit-<br>able in other respects.

collected ranged from 229 to 2, 280 ppm (Table carbonate hardness. Carbonate hardness in-<br>3) Most samples contained less than 500 ppm cludes that portion of the calcium and magnesi-3). Most samples contained less than 500 ppm cludes that portion of the calcium and magnesi-<br>of discolued solids. Eight samples contained um that would combine with the bicarbonate of dissolved solids. Eight samples contained more than 1,000 ppm of dissolved solids.

Hardness of water is recognized most com monly by the amount of soap needed to a lather or suds and by an insoluble scum that forms during the washing process. Calcium and magnesium cause almost all the hardness of water and are the constituents that contribute to

contain enough of certain constituents to cause the incrustation that may develop when water

The total hardness of water may be divided<br>into two types—carbonate hardness and non-The dissolved solids in the samples of water into two types—carbonate hardness and non-<br>ected ranged from 229 to 2.280 ppm (Table carbonate hardness. Carbonate hardness inand the small amount of carbonate that are present. Carbonate hardness can be virtually re moved by boiling the water, thereby causing HARDNESS precipitation of magnesium and calcium car bonate . Noncarbonate hardness is the difference between the total and the carbonate hardness and is caused by that portion of calcium and mag nesium that would combine with the sulfate, chloride, and nitrate ions that are present, plus the slight hardness effect of other minor con-



FIGURE 15.—Graphic representation of chemical constituents in samples of water from wells in principal aquifers in Lrego County, Kansas.

stituents. Noncarbonate hardness cannot be re-<br>moved by boiling.<br>collected ranged from less than 1 ppm to 348

Water that has a hardness of less than 60 ppm (Table 3). Most samples were low in ppm is classified as soft. Hardness of 60 to 120 nitrate content although eight samples ranged ppm is classified as soft. Hardness of 60 to 120 nitrate content although eight samples ranged<br>ppm will cause an increase in the amount of between 45 and 90 ppm and three exceeded 150 scap required for washing but will not interfere with the use of the water for most purposes—although water in the upper part of this range<br>will cause considerable scale in steam boilers. will cause considerable scale in steam boilers. Fluoride generally is present only in small<br>Hardness of 120 to 180 ppm will cause a hard- amounts in ground water. However the fluor-Hardness of 120 to 180 ppm will cause a hard-<br>hess that is quite noticeable. Water that has a side content of drinking water should be known ness that is quite noticeable. Water that has a ide content of drinking water shculd be known<br>hardness of more than 180 ppm is considered because if children drink water containing too hardness of more than 180 ppm is considered because if children drink water containing too<br>very hard. Where municipal water supplies are much fluoride during the formation of permavery hard. Where municipal water supplies are much fluoride during the formation of perma-<br>settened, the hardness is generally reduced to ment teeth, mottling of the enamel may result. softened, the hardness is generally reduced to ment teeth, mottling of the enamel may result.<br>about 100 ppm. The fluoride content is as much as 4 ppm.

Most samples of the water collected were about 90 percent of the children using the water<br>high in total hardness (carbonate and noncar- may develop mottled tooth enamel (Dean, high in total hardness (carbonate and noncar- may develop mottled tooth enamel (Dean, bonate). Most of them ranged from 100 to 400 (1936). Although too much fluoride has a detribonate). Most of them ranged from 100 to 400 1936). Although too much fluoride has a detri-<br>ppm, with the greatest proportion ranging from mental effect, a smaller amount in drinking ppm, with the greatest proportion ranging from mental effect, a smaller amount in drinking<br>200 to 300 ppm. Only four samples had a total water, about 1 ppm. lessens the incidence of hardness of less than  $100$  ppm (Table 3).

The nitrate content of vary greatly, and in many ground waters nitrates carriers. may seem unrelated to any geologic formation . Although some nitrate may be derived from The fluoride content nitrate-bearing rocks and minerals in bearing formation, strong concentrations of nitrate probably are due to other sources . trates are dissolved readily from soils that contain nitrate concentrations derived from plants, nitrate fertilizer, animal waste, or nitrifying bacteria . High nitrate concentrations in water may be due to the flow of surface water into a well. In an area where privies, cesspools, barnyards are sources of organic nitrogen, a socks-contain small to large amounts of emotion large amount of nitrate in cate harmful bacteria or pollution .

n the last two decades, investigations into water the effects of nitrate on the human system have concentration shown that too much nitrate in water may cause cyanosis in infants (so-called blue babies ) when the water is taken in directly or used in the formula for feeding. the Kansas State Department of Health and the U.S. Public Health Service regard 45 ppm as the safe limit of nitrate (as  $\text{NO}_3$ ). of nitrate is equivalent to 10 ppm of nitrogen . Water containing as much as 90 ppm of generally is considered very dangerous to inlants, and water containing as much as 150 ppm may cause severe cyanosis. Moderate nitrate concentrations seemingly are not harmful to older children or adults. Nitrate cannot be removed from water by boiling.

ved by boiling. collected ranged from less than 1 ppm to 348<br>Water that has a hardness of less than 60 ppm (Table 3). Most samples were low in between 45 and 90 ppm and three exceeded 150 ppm.

#### FLUORIDE

and 100 ppm.<br>Most samples of the water collected were about 90 percent of the children using the water water, about 1 ppm, lessens the incidence of tooth decay (Dean, et al., 1941). The U.S. Public Health Service ( 1962 ) recommends the NITRATE STAND Standards for content of mineral constituents in natural water may drinking water that are to be used on interstate carriers. The recommended maximum content tor fluoride is 1.5 ppm.

of the samples of water the water-<br>
collected ranged from 0.06 to 5.2 ppm (Table 3) Although most samples contained less than 1.0 N<sub>i-</sub> ppm, seven samples exceeded 1.5 ppm. Of five<br>anglyces of water from the Dakota Formation of water from the Dakota Formation , four exceeded 1.5 ppm of fluoride .

#### CHLORIDE

Chloride is abundant in nature and many and rocks contain small to large amounts of salts which may be dissolved by ground water. well water may indi-<br>Ilution e<mark>r ordinary use, unless present in s</mark>uch  $\mathbf s$  to make the water nonpotable or corrosive. Water that contains less than 150 ppm of chloride is satisfactory for most pur poses . Water containing more than 250 ppm Both generally is objectionable for municipal supplies, of Health and the and water containing more than <sup>350</sup> ppm is ob jectionable for most irrigation or industrial uses ; **This amount** water containing 500 ppm has a disagreeable taste. However, animals can tolerate water with  $\frac{a}{c}$  a much greater chloride concentration (e.g., con centrations of as much as  $4,000$  to  $5,000$  ppm can be tolerated by cattle . )

**Moderate nitrate The chloride content of the water samples** collected ranged from <sup>7</sup> ppm to 368 ppm (Table 3). Most samples were low in chloride; all but three samples contained less than 100 ppm. )



IRON

state, but upon exposure to air most of the iron divided Iron and manganese in quantities that exceed a rew tenths of a part per million are undesirable, as they stain fabrics and plumbing fixtures and produce an objectionable coloration The sodium content and taste in the water. Water in the ground gation may contain considerable fron in the ferrou is oxidized and precipitated as reddish -brown terric hydroxide. Iron can be removed fron most water by aeration and filtration , but some water requires additional treatment. Drinking water standards recommended by the U.S. Department Public Health Service are that the iron content should not exceed 0.3 ppm and that the manganese content should not exceed 0.05 ppm .

The iron content of the water samples  $col$ - and the Ogallala Formation were used<br>the material content of the Content of the content of the formation of the content lected ranged from 0.03 to 4.7 ppm ( Most samples contained less than  $0.3$  ppm of iron, but three samples contained more than U.S. Department 1.0 ppm.

#### SULFATE

Suitate ( $SO_4$ ) in ground water principally from gypsum or sulfate) and from the oxidation of pyrite  $($ disulfide). Magnesium sulfate (Epsom salt) sodium sulfate (Glauber's salt), if present in sufficient quantities, impart a bitter taste to the water, and the water may act as a laxative on people not accustomed to drinking it.More than 250 ppm of sulfate in drinking water gen erally is undesirable .

Most water samples collected were low in sulfate, although nine samples exceeded 250 ppm and one exceeded 1,000 ppm of sulfate  $($ l able  $3)$ 

#### SILICA

Silicon combined with oxygen in the form of  $SiO<sub>2</sub>$  is called silica. Silica is a mineral constituent in most ground waters. Except for the scale it may form, silica has little effect on the use of water for most purposes . Silica may be deposited as scale with other incrustants, ger erally in the form of calcium or magnesium sili cate. The silica content of the water samples collected ranged from 8 to 45 ppm (Table 3) Analyses of water from both the Dakota Forma tion and the Codell Sandstone Member of the Carlile Shale were characteristically low in silica content.

# **BICARBONATE**

of ground water . The concentration of bicar bonate, the predominant anion in the samples of water from Trego County, ranged from 194  $\sigma$  +15 ppm (Table 5)

# SODIUM

of water used for irri is important because <sup>a</sup> large percentage of sodium (equivalents per million of sodium by total equivalents per million of sodi m, potassium, calcium, and magnesium) has n adverse effect on soil, especially where leach ing is not adequate . The effect of sodium in Drinking Ingation water is discussed in detail in O.S. of Agriculture Handbook 60 (U.S. 1*7)* 17

> of water samples from 10 wells de riving water from alluvium, terrace deposits, to illus<br>T trate the suitability of water for irrigation. The procedure is based upon methods outlined in of Agriculture Handbook 60.

Figure 16 shows sodium -adsorption ratios determined by plotting the sodium content (equivalents per million ) on the left scale ( A), and the calcium plus magnesium content anhydrite (calcium ) (equivalents per million) on the right scale  $(B)$ . iron The point at which <sup>a</sup> line connecting these two points intersects the sodium-adsorption-ratio scale  $(C)$  indicates the sodium-adsorption ration



FIGURE 16.-Nomogram for determining the sodiumadsorption ratio <mark>of</mark> water.

Original from UNIVERSITY OF MINNESOTA



TABLE 5.—Sodium -adsorption ratios (SAR), conductivities, sodium content, and calcium plus magnesium content of water samples from selected wells

of the water. **Table 5 gives sodium-adsorption** ratios, electrical conductivities, and values for sodium and for calcium plus magnesium of the water samples for which analyses were plotted.

The specific conductance of a water sample n be measured directly in the laboratory, or it can be approximated by multiplying the total equivalents per million of the cations (calcium, magnesium, sodium, and potassium) by 100, or bydividing the dissolved -solids content in parts  $r$  million by 0.64.

Sodium-adsorption ratios and electrical conductivities are plotted in Figure 17 to provide <sup>a</sup> classification of waters for irrigation use. Lowsodium water (S1) can be used for irrigation on most soils with little danger of development of harmful levels of exchangeable sodium. Medi m-sodium water (S2) may be used safely on coarse -textured or organic soils having good permeability, but S2 water will present an appreciable sodium hazard in certain fine-textured soils , especially under poor leaching conditions . High-sodium water (S3) may produce harmful levels of exchangeable sodium in most soils and will require special soil-management techniques, such s good drainage, leaching, and addition of organic matter. Very high-sodium wate unless special practices are followed, such as  $\Theta(\cdot)$  generally is unsatisfactory for irrigatio addition of gypsum to the soil.

Low-salinity water (C1) can be used for irri gation of most crops on most soils with little iikelihood that soil salinity will develop. Medin-salinity water (C2) can be used it a mod erate amount of leaching occurs . Crops that olerate moderate amounts of salt, such as potagated with C2 water without special practices. corn, wheat, oats, and alfalfa, can be irri  $High-salinity$  water  $(C3)$  cannot be used on soils having restricted drainage. Very highsalinity water (C4) can be used only on certain crops and then only if special practices are fol

lowed. In Figure 17, all the waters were classined as low-sodium water (51) and either ) medium -salinity water ( C2 or high -salinity )water  $(C3)$ .

# SANITARY CONSIDERATIONS

I'm analyses of water in Table 3 give only the dissolved -solids content of the water and do not indicate the sanitary quality of the water , although <sup>a</sup> large amount of certain mineral con stituents, such as nitrate or chloride, may indicate pollution. Water containing mineral mater that imparts an objectionable taste or odor may be free from narmful bacteria and quite



safe for drinking. Conversely, water that is member is dominantly gray and red mottled clear and pleasant to the taste may contain clay. A zone of siderite, limonite, and hematite clear and pleasant to the taste may contain<br>harmful bacteria. Great care should be taken to protect domestic and public water supplies Gray clay containing siderite pellets and yellow-<br>from pollution. To guard against contamina- orange coloring matter dominates the upper from pollution. To guard against contamina- orange coloring matter dominates the upper<br>tion, a well must be properly sealed to keep out member, with lignite common in the upper part. tion, a well must be properly sealed to keep out member, with lignite common in the upper part.<br>dust, insects, vermin, debris, and surface water. Lenticular beds of sandstone occur throughout dust, insects, vermin, debris, and surface water. Lenticular beds of sand<br>Wells should not be placed where barnyards, the Dakota Formation. Wells should not be placed where barnyards, the Dakota Formation.<br>
privies, or cesspools are possible sources of pol-<br>
The Dakota Formation is present in the privies, or cesspools are possible sources of pollution.

# RELATION TO GROUND WATER

# LOWER(?) CRETACEOUS SERIES

#### DAKOTA FORMATION

The Dakota Formation [Early(?) Cretace-<br>ous does not crop out in Trego County, but it  $\frac{300}{200}$  feet. Swineford and Williams (1945) gave ous | does not crop out in Trego County, but it 300 feet. Swineford and Williams (1945) gave<br>contains water within a practical drilling depth. a thickness of 213 to 300 feet for the Dakota contains water within <sup>a</sup> practical drilling depth . <sup>a</sup> thickness of <sup>213</sup> to <sup>300</sup> feet for the Dakota In certain areas where shallow ground waters in southwestern Russell County. Latta (1950)<br>are scarce or not available, the Dakota is an im-<br>indicated the Dakota was 200 to 300 feet thick portant source of ground water.<br>The term "Dakota group" (Meek and Hay-

The term "Dakota group" (Meek and Hay-<br>den, 1862) was first applied to the varicolored near the Mitchell-Cloud county line. Leonard clay, sandstone, and lignite beds underlying the "Benton group" in exposures near Dakota City, Dakota County, Nebraska. The use of grouping has been discontinued (Plummer and Romary, 1942), and the Dakota Formation now Gove County.<br>includes attack hing between the Viewe Shale includes strata lying between the Kiowa Shale below and the Graneros Shale above. The Dakota Formation as defined by the State Geologi al Survey of Kansas has been variously referred to as "Dakota group" (Meek and Hayden,  $1802$ , Dakota sandstone († 1038), 1027),  $16.19941$   $16.19941$ 17akota formation (Twenholel, 1924); "Cockrum sandstone" (Latta, 1941). Rubey o the upper 125 feet of Bass (1925 ) referred the Dakota in Russell County as the "Rocktown" from the Dakot<br>channel sandstone member." Merriam (1957) ical analyses of fical analyses of five water samples from the<br>used the term "Omadi formation" for the Da-<br>Dakota are given in Table 3. Water from the Lakota are given in Table 5. Water from the<br>kota Formation in the subsurface of western Dakota is more mineralized than water in the Kansas .

#### CHARACTER AND SUBDIVISIONS

In Kansas the Dakota Formation has been and Depth divided into two members (Plummer and Romary,  $1942$ ), the lower called the Terra Cotta land surface Clay Member and the upper called the Janssen Clay Member. Plummer and Romary described tion. the Dakota Pormation as consisting emeny of varicolored clay containing irregular, lenticular beds of siltstone and sandstone .

pellets marks the base of the upper member.<br>Gray clay containing siderite pellets and yellow-

subsurface throughout the County. The depth to the Dakota ranges from about <sup>300</sup> feet below the land surface in the southeastern part of the **GEOLOGIC FORMATIONS IN** County to about  $1,000$  feet below the land sur-<br>**REI ATION TO GROUND WATER** face in the northwestern part.

#### THICKNESS

CRETACEOUS SYSTEM-<br>
The thickness of the Dakota Formation is not uniform, probably because of the disconformity at the base. Frye and Brazil (1943,  $table 3)$  gave the thickness of the Dakota Forindicated the Dakota was 200 to 300 feet thick in northern Barton County. Hodson (1959) near the Mitchell-Cloud county line. Leonard and Berry (1961) indicated that in southern Ellis County the thickness of the Dakota was  $\frac{1}{2}$  only 155 feet. Hodson and Wahl (1960) gave a of about 250 feet for the Dakota in A study of drillers' logs indi thickness of 150 to 250 feet for the I he Da- Dakota Formation in Trego County.

# WATER SUPPLY

In the east -central and southern parts of Trego County, a number of wells obtain water and for domestic and stock use from the Dakota Rubey and Formation. Information on eight wells ranging in depth from 500 to 670 feet that derive water is given in Table 6, and cher ical analyses of Dakota are given in Table 3. western Dakota is more mineralized than water in the unconsolidated rocks, but generally the water is soft.

o the water level in wells ranged to about 400 feet below the or wells obtaining water from the in this investiga The sandstone lenses in the Dakota contain ground water under artesian pressure, and, lenticular although locally there are in hydro The lower static pressure from well to well (Leonard and

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Berry,  $1961$ , p.  $26$ ), ground-water investigations in much of Kansas suggest a connected hydrologic system among the sandstone lenses, for most part.

Wells in the Dakota Formation in County yield only quantities of or domestic or stock supplies, although in other areas in Kansas wells penetrating thick Trego County . sandstones in quantities of water, generally in areas nearer tion Dakota outcrops.

# **UPPER CRETACEOUS SERIES**

The Graneros Shale overlies the Dakota Formation and consists chiefly of dark -gray to noncalcareous shale. The Graneros is present in the subsurface of Trego County where drillers' legs indicate it is 40 to 50 feet thick. It yields shown o water to wells in Trego County.

GREENHORN LIMESTONE upper part of the UNESCOPE UPPER THE Creenhorn Limestone overlies the Gra- abundant upward. The Greenhorn Limestone overlies the Graneros Shale and consists principally of and calcareous shale. The Greenhorn Limeston of alternating thin beds of chalky limestone Blue Hill. is divided into four members, which, in ing order, are the Lincoln Limestone, Shale, Jetmore Chalk, and Pfeifer Shale (Rubey - ous clay-ironstone and Bass,  $1925$ ; Bass,  $1926$ ). The chalky lime- locally, stone and shale that comprise the Greenhorn found. are relatively impervious and do not yield water <sup>a</sup> o wells in Trego County.

The Carlile Shale, consisting of the Chalk, the Blue Hill Shale, stone members, is the oldest formation exposed top In Trego County. Only the upper part of Carlile crops out, chiefly along the Smoky Hill brown fine sand, Valley in the southern part of the County.

#### CHARACTER AND SUBDIVISIONS

The Carlile Shale was named by Gilbert (1896 ) from exposures of west of Pueblo, in Pueblo County, Colorado. Logan ( 1897 ) correlated Gilbert's section with Cretaceous rocks in north -central Kansas . and Bass (1925) divided the Carlile Shale into<br>the carlier of the carlier into the late of the carlier of the carlier of the carlier of the carlier of the car<br>of the carlier of the carlier of the carlier of the carlier of two members, the "Fairport chalky shale" and the "Blue Hill shale" above. The Code Sandstone Member (upper part of the Carlile The thickness Shale) was named by Bass (1926) posures along the Saline River in Ellis County, Kansas .

The Fairport Chalk, lowermost member of<br>the Carlile Shale, consists of alternating beds of  $\frac{1}{2}$  calcareous shale and thin, nodular, chalky lime stone. The Fairport does not crop out in Trego County.

water sufficient The Blue Hill Shale, middle member of the some Carlile Shale, is the oldest rock unit exposed in Trego County. The Blue Hill crops out along the Smoky Hill Valley eastward from the juncof the Smoky Hill River and Hackberry Creek . The Blue Hill is well exposed on the steep slopes below the escarpment made by the Niobrara Chalk . Because the shale is soft , how CRETACEOUS SYSTEM ever, good exposures are rare elsewhere.

The Blue Hill Member is a blocky to fissile, GRANEROS SHALE clayey shale that characteristically weathers into brittle flakes. The predominant color of black the shale is dark blue-gray, but locally the shall is dark blue-gray, but locally the shal is . the Blue Hill may be very dark -gray to black as in several logs of test holes at the end of this report. I hin lenses of siltstone, fine-grained sandstone, and sandy shale are common in the of the Blue Hill and become more

about 100 Concretions characterize outcrops of the Most of the concretions are found in zones and become much more abundant in ascend the upper part. Most of the concretions are Hartland - calcareous septarian concretions, but noncalcare concretions are common, and, calcareous sandstone concretions are I'm concretions range in diameter from few inches to as much as 3 feet. Most are ellipsoidal or discoidal, but smaller concretion may be nearly spherical. The concretions con-CARLILE SHALE tain intersecting veins of brown calcite that on Fairport weathering tend to stand out in relief.

The Codell Sandstone Member marks the of the Carlile Shale. In Trego County, out the crops of the Codell consist of a zone of rust sandy silt, and clayey silt. It is generally very argillaceous; shark teeth and bone fragments are common in it. In fresh exposures, the Codell is light gray to gray. Specks of disseminated limonite cause an orange to gray argillaceous shale yellowish -brown coloration in weathered out crops. The contact between the Codell and the Rubey underlying Blue Hill Shale Member is gradational, whereas the contact between the Codel and the overlying Fort Hays Limestone Member of the Niobrara Chalk is generally sharp.

 of the Codell Sandstone Mem from ex- ber changes laterally in Trego County, but usually about 2 or 3 feet of Codell is present. Locally, however, as much as 5 feet of sandy

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It representing the Codell was observed, while stone it a few localities, only a few inches was noted. A study of drillers' logs indicated that in northeastern part of the County , the Codell may be as much as 20 feet in thickness .  $\overline{\phantom{a}}$ northern Ellis County, along the Saline Valley. 1962) reported a thickness of nearly 20 feet in

# DISTRIBUTION AND THICKNESS

The Carlile Shale crops out along the Smoky Hill Valley eastward from the mouth of berry Creek. Drillers logs indicate that the nearly vertical, Carlile is nearly <sup>300</sup> feet thick and that the Blue along the Smoky Hill Valley . Hill Shale Member comprises about two-thirds of the formation .

#### WATER SUPPLY

The Codell Sandstone Member yields water o a number of wells in Trego County, mostly in the eastern and southern parts. Information n 19 wells ranging in depth from 31 feet to feet below the land surface is given in Table 6, and chemical analyses of six water samples from shale, the Codell are given in Table 3. yields only small amounts of water to wells , although the yield is generally adequate for mestic and stock supplies.

n the southern part of the County, where in much of the area only meager amounts of low ground water are available, small amounts crops. of ground water are obtained from shallow wells dug or drilled into colluvium and into the the abundant fossils underlying weathered shale and sandy shale of the upper part of the Blue Hill Shale Member.

The Niobrara Chalk was named by and Hayden (1862) from exposures of calcare- ous ous marl and chalky limestone near the mouth of the Niobrara River in northeastern Nebraska . Logan ( 1897 ) described the Niobrara in central Kansas and divided it into the Fort Hays the chalk . Limestone Member below and the Smoky Hill Chalk Member above.

The Niobrara Chalk consists chiefly of ternating beds of light -gray chalk , stone, and chalky shale. The total thickness of the Niobrara in west-central Kansas is about  $700$  feet, but only the lowermost  $400$  to  $500$  fee is present in Trego County , the upper part hav- The Niobrara Chalk ing been removed by erosion .

The Fort Hays Limestone Member is guished from the Smoky Hill, upper member of the Niobrara Chalk , by the nance of massive beds of

 $\frac{0}{10}$  to  $\frac{60}{100}$  feet in thickness. The chalky limestone beds in the Fort Hays generally range the  $\frac{1}{10}$  to  $\frac{1}{2}$  reet in thickness and are separated by thin partings of chalky shale. These beds contain large *Inoceramus* and small Ostrea shells. The Fort Hays is characteristically grayh-white, but locally may be stained yellow or light brown on the outcrop. The Fort Hays is o erosion than the overlying to form Hack- shoulders along slopes. Steep clifts, in places are typical of the Fort Hays in the southeast of the County, the Fort Hays forms prominent escarpments. The Fort Hays has been quarried at places along the outcrop for building stone and for<br>road metal.

#### SMOKY HILL CHALK MEMBER

The Smoky Hill Chalk Member consists principally of thin -bedded chalk and chalky with thin beds of bentonite throughout . The Codell Fresh exposures are platy and light to dark gray, but the beds weather colorfully to white, do- orange, and brown. Concretions of limonite and pyrite are common, and they account for the bright colors of the Smoky Hill when weathered.<br>The Smooth Color of the Smoky Hill when weathered. shal- Thin veins of gypsum are characteristic of out

is noted for it contains. Vertebrate fossils include bones of aquatic reptiles, such as mosasaurs and plesiosaurs, and numerous fish remains. Sharks' teeth are common. Invert NIOBRARA CHALK brate fossils characteristically include the phyla Meek Mollusca and Echinodermata, the most numer of which are the genera *Inoceramus*, a clam, n oyster. Minute shells of for minifers belonging mainly to the families Globi north- gerinidae and Textulariidae comprise much of

> is also notable for its effec n topography. Soil development is either thii or absent, vegetation is sparse, and rainwash and chalky lime- gully crosion produce <sup>a</sup> badlands type of topog raphy.

#### WATER SUPPLY

is not important as an aquifer in Kansas and carries very little ground water in Trego County . The beds of shaly chalk FORT HAYS LIMESTONE MEMBER are relatively impermeable, and water is transmitted chiefly through fractures and joints and is found only locally and in small quantities. Fractures filled with secondary calcite are indicative that water has circulated through the for-



mation at some time in the past. Where the solidated, although cementation of beds occurs<br>Niobrara was observed to be fractured, however, to some degree throughout the Formation. Cal-Niobrara was observed to be fractured, however, it was generally in the topographically higher parts of the County well above the water table. Fractures and bedding planes are more prevalent n the Fort Hays Limestone Member, may yield small amounts of ground water to a few wells.

The Ogallala Formation was named by Darn in  $1899$  (p. 732-734) from exposures in southwestern Nebraska . Darton (1920 , p6) designated the type locality as being near Ogalala Station in western Nebraska. Since Darton's occasionally seen. work, the most significant studies of the Ogallala and, in western Kansas have been by Elias (1931 ), Smith  $(1940)$ , and Frye, Leonard, and Swine-  $\frac{a_{10}}{D_{10}}$ icra (1956*)*.

#### CHARACTER AND SUBDIVISIONS

The Ogallala Formation in Kansas is into three members which, in ascending order, are the Valentine, Ash Hollow, and Kimball.  $A$ thin, discontinuous pisolitic limestone, I to 3 feet thick , commonly occurs as the topmost bed and clayey of theOgallala . No attempt was made le Ogallala Formation in Trego County, and it is shown on Plate <sup>1</sup> as asingle unit .

The Ogallala Formation constitutes a wide- and thin bed spread mantle of fluvial deposits consisting pre- are common dominantly of sand, gravel, silt, and clay. The Ogallala was deposited upon an erosional sur- found locally face of Upper Cretaceous rocks by trending streams whose source of igneous rocks of the Rocky Mountains and sedi- cludes flat uplands , mentary rocks of eastern Colorado. e comprised chiefly of a series of valley fillings hdeposits, overlapping laterally from the axes of drainageways onto the gentle erosional slopes of the valley sides. Thus, the Ogallala Formation mented consists of a heterogeneous complex of predomi- color. nantly clastic deposits, with textures ranging sistant from very coarse gravel and pebbles to sorting ranging from good to poor . ogy changes sharply both vertically and laterally . Lentils of volcanic ash, marl, or marly limestone, and bentonite contrast with the predominantly stream laid clastics. Throughout this heterogene originally called the " ous assortment of sediments there is virtually no distinctive bed that can be able distance.

The deposits are interbedded and admixed in various proportions and are largely uncon- aerial origin,

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cium carbonate is a common constituent throughout almost all of the Ogallala. It is diss fine material and as stringers of and these caliche and small- to medium-sized nodules. Calcium carbonate in many places binds the deposits so firmly as to produce a series of hard ledges, interbedded with beds that are only NEOGENE SYSTEM—<br>PLIOCENE SERIES suppose that the process is a supported and form roughly weathered slighted cemented. The hard ledges are usually PLIOCENE SERIES<br>benches and cliffs which resemble mortar and<br>meventy cemented and form roughly weathered OGALLALA FORMATION accordingly are often referred to as "mortar Silica also is present as a cementing ma terial in beds of opaline sandstone or as cher deposits, and variously colored chert in the form of noquies or small, irregular lenses and beds is occasionally seen.

> Sand, the principal material within the Ogaliaia Formation, is present at all horizon is typically light-gray or greenish in color. Beds of uniform sand may occur, but generall the sand ranges from fine to coarse and com is mixed with gravel, silt, or clay. Grave divided beds containing lenses of sand , silt , and clay are common, but thick beds of uniform gravel are rare. Beds of sand and gravel with distinct cross-bedding occur in places. Silt, sandy silt, ilt are present throughout the Ogal to divide lala and are greenish-gray, pink, tan, and gray; if the beds contain <sup>a</sup> large amount of calcium carbonate, they are light gray or white. Lenses of white and pinkish limestones in the middle and uppermost parts of the Ogallala . Bluish -gray volcanic ash is in the lower and middle parts of the Formation.

sediment was The Formation's topographic expression in gentle erosional slopes, and The deposits nearly vertical cliffs. In spite of the diversity of the outcrop pattern of the Ogallal the main presents a uniformity of aspect that makes it readily identifiable. Typical outcrops are ce o various degrees and are ash gray in Because the cemented beds are more re le cemented beds are more reclay and form rough benches, hard ledges, and cliffs; **The lithol-** exposed surfaces commonly have a knobby, ifregular aspect.

Opinions differ regarding the origin of the thin, discontinuous bed of pisolitic limestone, algal limestone" by Elia (1931), which marks the stratigraphic top of the traced any appreci- - Ogaliala Formation. - Elias postulated a facustrine origin for the capping limestone. Subse an hypothesis of sub such as a caliche zone. Smith

(1940) discussed the two hypotheses, and more  $r$ ecently Frye, Leonard, and Swineford (1950, p. 13-16) critically discussed the bed, postulating amode of origin in which development of amature to senile lime-accumulating soil was The Ogallala Formatio later modified by solution .

#### DISTRIBUTION AND THICKNESS

The Ogallala Formation, although generally northwestern parts mantled with eolian silts, underlies much of interstream areas. The Ogaliala is most extensive in the central and northwestern which moderate yields parts of the County. It crops out along the bluff of Big Creek valley and locally is well exposed in many of the tributary canyons along the Sa-<br>A consider line River valley. The Ogallala is thin and discontinuous in the southern part of the County been removed due to erosion by the Smoky Hill River and its tributaries. In Trego County, the Formation and Big Creek ( rests on an erosional surface of considerable re  $l$ <sub>1</sub> which slopes generally eastward (Fig. 18)

Logs of test holes show that the thickness of the Ogallala Formation in Trego County ex-<br>seeds 150 feet in some of the thicker sections ceeds 150 feet in some of the thicker sections . The thickness is not uniform, however, of unconformable contacts at the top and bot- County. tom of the Formation. The thickness and char- lala are not large, acter of the Ogallala Formation are shown in the *Logs* of *I est Holes and Wells* at the end of

this report and are illustrated in the cross sections on Plate 2.

#### WATER SUPPLY

is the most wide spread water -bearing formation in Trego Coun y. It supplies water to most domestic and stock lls in the County. In much of the central and of the County, the saturated the thickness of the Ogallala is great enough to store thickest and hairly large quantities of ground water, from which moderate yields of water are obtained. In the southern part of the County, the Ogalial is either missing or thin and contains little

considerable amount of the Ogallala has by erosion along the three prin cipal streams ——the Smoky Hill and Saline Rivers Pl.  $2$ ). Few areas in the Couny have much more than 50 feet of saturated thickness and the saturated thickness of the Ogallala decreases toward the edge of the up land where the Ogallala has been completely of ground water. Nevertheless it is estimated that about 700,000 acre-feet of groun because water is in storage in the Ogallala in Trego Although yields of wells in theOgal it is the most important aqui er in the County because of its wide areal extent.



Figure 18.—Contours on base of Ogallala Formation in northern Trego County, Kansas (contour interval 50 feet)

The classification of Pleistocene deposits by e State Geological Survey of Kansas is n the classification of glacial deposits in Midcontinent Region (Frye and Leonard, 1952).  $\cdot$   $\cdot$   $\cdot$   $\cdot$   $\cdot$   $\cdot$ Aithough the nearest glacial deposits are more<br>than 150 miles portbeast of Trego County corthan 150 miles northeast of Trego County , cor relations between the glaciated and nonglaciated posits, which<br>areas have been made on the basis of continuous Terrace leess deposits, their molluscan fauna and buried the Supposits, their molluscan fauna and buried the Smoky Hill Valley,<br>solis, and petrologically distinctive volcanic ash lesser degree along the  $(0.000a, \text{Reed}, \text{and Gordon}, 1947)$ tord, and Leonard, 1948).

tively thin, are the surficial materials in much of  $\overline{r}$ Trego County, as shown by the geologic map<br> $\frac{m}{2}$  valley in the western part of the County. Along  $(P<sub>1</sub>, 1)$  and cross sections  $(P<sub>1</sub>, 2)$ . - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 100 deposits include the Grand Isand Formation and remnants,<br>Sappa Formation of Kansan age, the Crete For- eroded. sappa normation of Kansan age, the Crete For-<br>mation and Loveland Formation of Illinoisan and western parts of the County, terrace deposite age, the Peoria Formation of Wisconsinan age, and terrace deposits and alluvium of sinan and Recent age. The Grand Island, Sapa, and Crete formations occur as terrace deand Crete formations occur as terrace de-<br>terrace deposits are more<br>soling times and along the Creek walley Alliu. Saline rivers and along Big Creek valley. Alluvium of Figure of the larger streams. thin cover<br>along the inner valleys of the larger streams. loess tend Eolian silts classified as formations cover much of extend along the outer valleys, masking the valley slopes. Pleistocene deposits have been treated n this report as terrace deposits (Crete, Sappa, and Grand Island formations), a and Loveland formations), and alluvium (late Wisconsinan and Recent deposits). Figure 6 shows the stratigraphic relation of Pleistocene sand, deposits along the principal valleys.

I ERRACE I DEPOSITS<br>Terrace deposits shown on the geologic map accompanying this report (PL 1) consist chiefly of sand, gravel, and silt deposited during Kanhas a n and Illinoisan times. m and firmorsan times. Terrace deposits along modern river, underliver and the Smoky Hill Valley are predominantly sand<br>and area with the predominantly sand Formation along much and gravel, whereas terrace deposits along the and gravel, whereas terrace deposits along the valley,<br>Saline River and Big Creek valleys contain a hole in the relatively greater amount of finer clastic ma terial with silt being predominant in the upper part. Included in the terrace deposits (Pl. 1) e the Grand Island Formation and Sappa For-<br>The Grand Island consists predom mation of Kansan age and the Crete Formation The Grand Island consists predominant. of Illinoisan age. The Crete and the Grand sand and gravel containing cobbles Island are lithologically similar, both consisting  $\tau$ chiefly of arkosic sand and gravel.

NEOGENE SYSTEM ically , the Sappa Formation lies between the Crete and the Grand Island. The Sappa consists chiefly of silt and sandy silt but locally con tains the Pearlette ash bed which identifies it as based late Kansan in age. The Crete Formation of  $\lim_{(2)}$  Illinoisan age constitutes most of what are called terrace deposits in this report. e restricted to extensive area than the underlying Kansan de

continuous Terrace deposits crop out extensively along but they also occur to and petrologically distinctive volcanic ash lesser degree along the Saline Valley and along<br>Ital Reed and Gordon 1947: Frve Swine Frye, Swine-<br>Big Creek valley. These deposits are nearly 2 miles wide along much of the Smoky Hill Val-<br>Deposits of Pleistocene age, although rela-<br>ley and in places little dissection has occurred y and in places little dissection has occurred in them, particularly along the south side of in the western part of the County. Along Pleistocene the Saline Valley, terrace deposits are mostly since most of the deposits have been Along Big Creek valley in Illinoisan and western parts of the County, terrace deposits are of narrow width and tend to be obscured Wiscon-<br>because of slope wash from the Ogallala. Along the north side of Big Creek valley in the easter part of the County, in places. where the terrac Allu-<br>ceurs deposits have not been dissected by erosion, a of loess mantles them . Although the loess tends to mask the edge of the terrace dethe Peoria and Loveland posits nearest the upland margin, these terrace the upland areas and<br>eys masking the value deposits are a geomorphologically prominent in the major valleys .

#### GRAND ISLAND AND SAPPA FORMATIONS

loess beds (Peor-<br>ad alluvium (late The oldest Pleistocene deposits definitely recognized in Trego County are stream-deposited gravel, and silt classified as the Grand Island Formation and Sappa Formation of Kan san age, although deposits penetrated in tes holes  $13-21-5$ bbc and  $13-21-8$ abb may be Nebras in age. Along the Smoky Hill Valley, Kan san deposits fill a narrow channel about a quarer of a mile wide. This ancient channel, which meander pattern much like that of the underlies the widespread Crete of the terrace extent of but crops out only locally. In a tes hole in the SE cor. NE sec.  $27, T$  14 S, R  $21$  W, **93 feet of terrace deposits were penetrated, the** lower 54 feet of which are classified as the

> The Grand Island consists predominantly of as much as inches in diameter . The sand and gravel are mostly arkosic but intermixed with the deposits

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are rounded chalk pebbles and shale fragments. material is sufficiently thick, moderately large The Sappa Formation consists mostly of silt sandy silt, but thin lenticular beds of gravel with pebbles of chalk are common . The Sappa is generally very calcareous . It is buff or light gray when fresh but characteristically weathers to very light gray. Locally, a lenticular Chanacter, DISTRIBUTION, AND THICKNE bed of volcanic ash (Pearlette ash bed ) n thickness from a few inches to several feet is  $\frac{a}{b}$  found within the Sappa. Outcrops of the Grand  $\frac{a}{b}$ Island and Sappa formations occur in small areas along the Smoky Hill Valley.

No outcrops of Kansan deposits were found the gentle slope along the Saline Valley, but test drilling has thicke shown that deposits of the Illinoisan terrace surface in a stratigraphi position analogous to that found along the geologic map (i) Smoky Hill Valley. In a test hole drilled in Saline Valley in the SE cor. NE sec. 9, T 11 S, R21 W, 70 feet of terrace deposits were pene trated, the lower part of which is considered to e Kansan in age. Deposits considered to be of Kansan age were penetrated in test holes along mation and separate Big Creek and underlie the Illinoisan terrace at least locally along Big Creek, especially along eolian,<br>diameter of the uple the north side of the valley in the eastern part the upland area of the County .

#### CRETE FORMATION

Deposits classified as the Crete Formation of Illinoisan age constitute most of the deposits that underlie the prominent terraces along the major valleys. The Crete Formation consists the underlying rocks they have been included chiefly of arkosic sand and gravel. A amount of chalk and shale pebbles and frag ments are usually present but are not as com mon as in the Grand Island. Locally, the upper part of the Crete grades upward into yellow and buff silt and sandy silt classified as Formation of late Illinoisan age .

#### WATER SUPPLY

Terrace deposits that border the principal stream valleys are generally continuous with the alluvial deposits that underlie the streams and Ill the inner valleys, and the water table in alluvium is continuous with the water table in the terrace deposits along much of the valleys. Because the terrace deposits overlie bedrock that was not cut as deeply as the incised valley of the present cycle, the saturated thickness of terrace deposits is not as great as the saturated thickness of the alluvium of the inner valleys .

Terrace deposits yield water adequate for domestic and stock supplies at where they occur. In areas where the saturated – water drains from much

and yields from wells can be expected . Test drilling will disclose the more favorable well sites.

## PEORIA AND LOVELAND FORMATIONS

# CHARACTER, DISTRIBUTION,

ranging Eolian silts (loess ) of late Pleistocene age cover <sup>a</sup> considerable part of Trego County with relatively thin mantle which tends to subdue the topography in much of the area. The loes caps the rolling hills and fat uplands and masks of the valleys. It is generall in the northwestern part of the County, Kansan age underlie where test-hole data indicate the loess may be 25 or 30 feet thick . The loess is shown on the Pl. 1) as the Peoria and Lovelan the formations undifferentiated.

The Loveland Formation is a reddish-tan lt, mostly eolian, which characteristically grade into sand in the lower part. The Sangamo buried soil marks the top of the Loveland For it from the overlying Peoria Formation. The Peoria is a massive,  $\frac{1}{\pi}$  to gray silt, which covers much of of Trego County.

Colluvial deposits, consisting chiefly of reworked loess, local bedrock fragments, and, in many places, sand and gravel slope wash derive from the Ogallala Formation, constitute much of the surficial material on the slopes of stream valleys and upland draws. Where these colli vial deposits are of sufficient thickness to small the underlying rocks they have been included<br>frag. with the Peoria and Loveland formations in this report.

## WATER SUPPLY

the Loveland The deposits mapped as Peoria and Love land formations consist mostly of relatively thin, wind deposited silts which generally lie above the water table. In parts of Trego County, how ever, where the Ogallala Formation has been removed by erosion and shallow ground-wate supplies are meager, domestic and stock wells the colluvial and slope deposits, which in this report of ground water from are included with the Peoria and Loveland for mations. In these areas, the best well sites gen to be in creek valleys and draws where thin alluvial and colluvial fill and uncon solidated material overlie the bedrock and serve is local catchment basins. In periods of ample rainfall these wells generally prove adequate for most domestic or stock supplies. During exmost places - tended periods of dry weather, however, the of this material and

wells tend to become dry. Because ground water n colluvial deposits is generally in e underlying Cretaceous bedrock and with penetrated bedrock fragments that are incorporated within the deposits, the water obtained from these wells feet thick. is generally of interior chemical quality.

#### ALLUVIUM

## CHARACTER, DISTRIBUTION,

Recent in age occurs in posits along the principal streams in the County. Thick, coarse alluvial deposits of sand and gravel are restricted to derived mostly by erosion of older alluvial deposits and from the Ogallala Formation. Thin, poorly sorted deposits of alluvium lie in smaller valleys and contain relatively less coarse material , the deposits being predominantly silt d fine sand. Deposits in the grade headward into colluvium and slope degrade headward into colluvium and slope de-<br>posits at the edge of the uplands. The lithology of the finer, less permeable material in the smal posits at the edge of the uplands. The lithology of the finer, less permeable material in the small<br>er valleys. Since the cross-sectional areas of the depends largely upon the type of rock into which the valley has been cut.

The alluvium that underlies the stream chan nels and the narrow floodplains along the principal streams is considered to be Recent in age. Low, relatively narrow terrace deposits, sidered to be late Wisconsinan in age, border the hormal railiant. In addition because of the rela-<br>tively shallow denth to water in the alluvium Hoodplain . Although the narrow floodplain and the low terraces are best developed along smoky<br>Hill Valley, Saline Valley, and Big Creek values water levels. pest developed along Smoky<br>growing se y, they are also found to a lesser extent along ic smaller valleys. The alluvium of Recent age and the alluvial deposits of e underlying the low terraces are lithologicall indistinguishable and are shown together as alluvium on Plate <sup>1</sup> .

 The width and thickness of greatest along Smoky Hill Valley, Saline Valley, and Big Creek valley. It reaches a maximum giver width of about a mile, but in most places it is than a half mile in width. erally has <sup>a</sup> thickness of 60 to 70 feet in

Ill along the Smoky contact with Hill Valley, although one test hole (14-24-35bcc) 95 feet of it Alluvium along the reek valley is 40 to 50 Alluvium in the smaller valleys is thin and of narrow extent, and headward these deposits grade into colluvium and slope wash.

### WATER SUPPLY

AND THICKNESS<br>late Wisconsinan and Alluvial deposits constitute an Alluvium classified as late Wisconsinan and Alluvial deposits constitute an important Alluvian equation of the<br>Source of ground water in Trego County Man source of ground water in Trego County. Man domestic and stock wells obtain water from the alluvium, and most irrigation wells in the County obtain water from alluvium along the the larger valleys and are<br>sion of older alluvial deo moderately large yields of water can be expected from wells in alluvium along the Smok Hill Valley.

> Ground-water yields from wells in alluvium in the smaller valleys can be expected to be con is than in the larger valleys, er valleys . Since the cross -sectional areas of rock into smaller valleys are small, water levels tend to nuctuate more in response to rainfall. Declining p of well: n the smaller valleys and upland draws can be expected during extended periods of below the normal rainfall. In addition because of border the tively shallow depth to water in the alluvium, by deep-rooted plants during the growing season often results in a decline of

## RECORDS OF WELLS

Information pertaining to wells is given in Table 6. Measured depths of wells are given to the alluvium are the nearest tenth of a foot; reported depths are given in feet . Measured depths to water are to the nearest hundredth of a foot; reless ported depths are given in feet. The well-num The alluvium gen- bering system used in Table <sup>6</sup> is described on the upage 7 and illustrated in Figure 2.





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14aaa

1.Littlechild

C<sup>Sand</sup>

,silt Colluvium

 $\frac{1}{2}$ 

27.23 9-21-60 2,309.6



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6. -Records of wells inTrego County ,Kansas .-Concluded

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# LOGS OF TEST HOLES AND WELLS

Listed on the following pages are logs of <sup>176</sup> test holes and wells. Logs designated "sample logs" describe test holes from which samples were collected. The logs are numbered according to the well-numbering system illustrated in Figure 2. Locations of wells and test holes are shown on Plate 1. Plate 2 illustrates the characshown on Plate 1. Plate 2 illustrates the charac-<br>ter of material penetrated by the test holes. Water-level measurements are stated in feet below land surface.

10-24-34ddd . -Sample log of test hole in SE SE SE sec. 34, T 10 S, R 24 W, 50 feet west and 20 feet north of center of crossroads, Graham County (Prescott, 1455, p .92 ); drilled September 1952. Altitude of land surface 2,357.2 feet; depth to water 35.4 feet.



10-25-36ccc.————Sample log of test hole in SW SW SW  $\epsilon$ , 36, T 10 S, R 25 W, 20 feet cast and 15 feet north of center of crossroads ,Graham County (Prescott, 1955 , p.92 ); drilled September 1952. Altitude of land sur face  $2.418.8$  feet; depth to water  $31.5$  feet.

Shale , gray <sup>2</sup>







- Niobrara Chalk Fort Hays Limestone Member Shale , chalky , white <sup>5</sup> 26 Shale , dark -gray <sup>2</sup> 28
- 11-21-9add.——Sample log of test hole in SE SE NE sec. 9, T 11 S, R21 W, 0.4 mile south of NE cor . sec. 9, at edge of road; drilled October 1957. Altitude of land surface 2,099.8 feet; depth to water 19.88 feet.

NEOGENE

Thickness, Depth, feet fee!



11-21-22ccc. ----------Sample log of test hole in SW SW SW sec. 22, T 11 S, R 21 W, 10 feet north and 10 feet east of SW cor. sec. 22; drilled September 1960. Altitude of land surface 2,309.5 feet.

	Thickness, Depth, teet	teer
<b>NEOGENE</b>		
Pleistocene		
Peoria and Loveland formations		
	5	5
Sand, fine to very coarse		12
Pliocene		
Ogallala Formation		
Silt, very limy, light tan	8	20
Silt, very sandy, limy, tan	10	30
Sand, medium to coarse, clean, well		
	10	40
Sand, medium to very coarse, clean,		
	10	50
Sand, medium to very coarse; con-		
tains streaks of silty clay in lower	10	60
Silt, very limy; contains cemented		
streaks and thin lavers of fine to		
	5	65
Silt, clayey, tough, brown	5	70
Sand, fine to medium, loose	20	90
Sand, fine to coarse, loose	10	100
Sand, medium to very coarse, clean,		
	10	110
Gravel, fine; contains very coarse		
	8	118
CRETACEOUS		
<b>Upper Cretaceous</b>		
Niobrara Chalk—Smoky Hill Chalk Member		
Shale, light grav upper part, dark		
gray lower part  12		130

11-22-1**0dad.** Sample log of test hole in SE NE SE c. 10, T 11 S, R 22 W, 250 feet north of bridge, 75 feet south of driveway to farm home, at edge of road: augered July 1960. Altitude of land surface 2,161,5 feet; depth to water 26.1 feet.



11-22-11ccb . -Sample log of test hole in NW SW SW sec. 11 , T11 S, R22 W, 0.15 mile north of SW cor. sec. 11, 150 feet south of bridge, at edge of road; augered July 1960. Altitude of land surface 2,139.8 feet.





11-22-32aaa . -Sample log of test hole inNE NE NE sec. 32, T 11 S, R22 W, 40 feet west and <sup>8</sup> feet south of center of crossroads; drilled September 1960. Altitude of land surface 2,378.1 feet; depth to water 77.2 feet.  $TL11L$ Depih ,





 $C$ R

feet 11-22-36bba. - Sample log of test hole in NE NW NW sec. 36, T 11 S, R22 W, 0.2 mile east of NW cor. sec. 36, 150 feet west of bridge, at edge of road; drilled October 1957. Altitude of land surface 2,283.2 feet.





11-23-7dad. Sample log of test hole in SE NE SE  $-23-7$  data.  $-$  -sample tog of test hole in SE NE SE.<br>Sec. 7, T 11 S, R 23 W, 0.3 mile north of SE cor, sec.<br>7, 100 feet south of driveway to farm home, at edge of Tead : augered July 1960. Altitude of land surface  $24.264.9$  feet; depth to water  $21.9$  feet.



11-23-8bbb . -Sample log of test hole in NW NW NW sec . 8, T 11 S, R23 W, 30 feet east and 15 feet sauth of center of crossroads; augered July 1960. Altitude of land surface 2,290.6 feet.





11-23-8bcb . -Sample log of test hole in NW SW NW sec. 8, T 11S , R 23 W , 0.3 mile south of NW cor. sec. 8, at edge of road: augered July 1960. Altitud of land surface 2,265.2 feet; depth to water 22.67 feet. Thickness, Depih, feet feel NEOGENE Pleistocene Alluvium Silt , sandy , brown 5. 5.Silt , sandy , tan brown <sup>7</sup> 12 Silt , heavy , brown <sup>8</sup> 20 Silt, slightly sandy, light brown ...... 8 28 Silt, very sandy, light brown  $\ldots$ ........ 9  $\qquad$  37 **CRETACEOUS** 

Upper Cretaceous Niobrara Chalk --Smoky Hill Chalk Member Shale , blue gray 6.5 43.5

11-23-8cbb . -Sample log of test hole in NW NW SW sec. 8, T 11 S, R 23 W, 0.4 mile north of SW cor. sec. 8, 150 feet south of bridge , at edge of road ; augered July 1960. Altitude of land surface 2,249.4 feet ; depth to water 7.0 feet.



11-23-8ccc. Sample log of test hole in SW SW SW sec. 8, T 11 S, R 23 W, 0.1 mile north of SW cor. sec. 8, at edge of road; augered July 1960. Altitude of land surface  $2,284.4$  feet; depth to water  $35.15$  feet.



11-23-35bbb . -Sample log of test hole in NW NW NW sec. 35,TIL S, R23 W, 25 feet eastand 25 feet south of NW cor, sec. 35; drilled September 1960. Altitude of land surface 2,417.6 feet .

 $\mathbf{v}_k$ 





tude of land surface 2,282.9 feet; depth to water 5.5

11-24-11bbc.——Sample log of test hole in SW NW NW sec. 11 , T 11 S, R 24 W, 50 feet south of farm driveway, at edge of road; drilled October 1957. Alti tude of land surface 2,290.2 feet; depth to water 12.8 fcet .



feet.

11-24-32aaa. - Sample log of test hole in NE NE NE sec. 32, T 11 S, R 24 W, 40 feet SW of center of crossroads ; drilled September 1960. Altitude of land sur face 2,519.4 fect .



11-25-2dda . -Sample log test hole NE SE SE sec. 2, T 11 S, R 25 W , 0.25 mile north of SE cor . sec .2, at edge of road ; augered July 1958. Altitude of Land surface 2,401.2 feet.

Thickness, Depth, feet feet





 $11-25-2d d$ . Sample log of test hole in SE SE SE Silt, sandy, tan  $\frac{1}{2}$  s. T 11 S, R 25 W, 0.1 mile north of SE cor. sec. 2, at edge of road: augered July 1958. Altitude of land surface 2.391.6 feet.  $T$ kich Depth,



NEOGENE



11-25-5bbc . -Sample log of test hole in SW NW NW sec. 5, T 11 S, R 25 W, 0.2 mile south of NW cor. sec. 5, 500 feet north of bridge, at edge of road; augered July 1960. Altitude of land surface 2,423.0 feet .



11-25-5bcb.<br>NW sec. 5, T 11 S, R 25 W, 0.25 mile south of NW cor. sec. 5, 50 feet north of bridge, at edge of road; augered July 1960. Altitude of land surface 2,414.5 feet.  $T$ hickness



11-25-6aaa . -Sample log of test hole in NE NE NE sec. 6, T 11 S, R 25 W, 250 feet south of NE cor . sec. 0, 1 11 S, R 25 W, 250 feet south of NE cor. sec.<br>6, at edge of road; augered July 1960. Altitude of land surface  $2,429.6$  feet; depth to water  $17.5$  feet.



sec. 6, at edge of road across from driveway to field;<br>augered July 1960. Altitude of land surface 2,434.5 at of feet; depth to water 24.2 feet.







11-25.11aaa . -Sample log of test hole in NE NE NE e. 11, T 11 S, R 25 W, near NE cor. sec. 11; augered July 1958. Altitude of land surface 2,390.2 feet.

Thickness, Depth, feel feet

NEOGENE





11-25-1 1aad . -Sample log of test hole in SE NE NE sec. 11, T 11 S, R 25 W, 0.15 mile south of NE cor. e. 11, at edge of road; augered July 1958. Altitud of land surface 2,366,5 feet.

	Thickness, Depth, feet feet	
<b>NEOGENE</b>		
Pleistocene		
Grand Island, Sappa, and Crete formations,		
undifferentiated (terrace deposits)		
		5
		12
	3	15
Sand, fine to medium  10		25
Sand, fine to medium, and dark silt-	$\sim$	28
CRETACEOUS		
Upper Cretaceous		
Niobrara Chalk-Smoky Hill Chalk Member		

11-25-11ada.———Sample log of test hole in NE SE NE sec. 11, T 11 S, R 25 W, 0.3 mile south of NE cor. sec. 11, at edge of road; augered July 1958. Altitud of land surface 2,352.4 feet; depth to water 17.85 feet. Thickness, Depih, feet feet







11-25-11bbb.——Sample log of test hole in NW NW NW sec. 11, T 11 S, R  $25$  W, 100 feet east of bridge, at edge of road; drilled September 1960. Altitude of land surface 2,363.9 feet; depth to water 6.3 feet.

Thickness, Depih , feet feei Pleistocene Alluvium Silt, sandy, brown  $\ldots$  $\ldots$  $\ldots$  $\ldots$  $\ldots$  $\ldots$ 7Sand, hne to coarse  $\ldots$   $\ldots$   $\ldots$   $\ldots$   $\ldots$   $\ldots$ 12 Sand, medium to very coarse, and fine gravel <sup>8</sup> 20 Sand, medium, well sorted, clean ... 9 29

- CRETACEOUS Upper Cretaceous Niobrara Chalk-Smoky Hill Chalk Membe Shale , gray 11 40
- 11-25-12bcc.——Sample log of test hole in SW SV NW sec. 12, T 11 S, R 25 W, 70 feet south of bridge, at edge of road; augered July 1958. Altitude of land surface 2,338.7 feet.



11-25-17bbb . -Sample log of test hole in NW NW NW scc. 17, T 11 S, R 25 W, 40 feet east and 10 fee south of center of crossroads; drilled August 1960. Altitude of land surface 2,553.8 feet.





11-25-21aad . -Sample log of test hole in SE NE NE sec .21 , T 11S , R 25 W, near west side of bridge ; hand augered June 1958. Altitude of land surface  $2.466.8$  feet; depth to water  $14.0$  feet.



11-25-27aaa.——Sample log of test hole in NE NE NE sec. 27, T 11 S, R 25 W, 30 feet SW of NE cor. sec.<br>27; drilled August 1960. Altitude of land surface<br>2,537.4 feet: depth to water 71.0 feet.

Thickness, Depih,



11-25-30ddd. Sample log of test hole in SE SE SE<br>sec. 30, T 11 S, R 25 W, 50 feet north and 10 feet west of center of crossroads; drilled August 1960. Altitude of land surface 2,528.6 feet .



Upper Cretaceous

- Niobrara Chalk-Smoky Hill Chalk Member Shale, chalky, brown and white ...... 7 50
	-
- 11-25-34ddd. Sample log of test hole in SE SE SE<br>sec. 34, T 11 S, R 25 W, 50 feet north and 10 feet west of SE cor. sec. 34; drilled September 1957. Alti-<br>tude of land surface 2,555.6 feet; depth to water 76.63 feet.



-26-13ddd.<br>SE sec. 13, T 11 S, R 26 W, Gove County (Hodson) and Wahl, 1960, p. 93); drilled June 1952 by Schaef er Geophysics. Altitude of land surface 2,540.3 feet; depth to water 37.8 feet.

Thickness, Depih, feer feet





5, at edge of school yard ; drilled October 1957. Alti tude of land surface 2,341.2 feet .



Thickness, Depih,



12-21-5ddd1.——Sample log of test hole in SE SE SE sec. 5, T 12 S, R 21 W, 150 feet north of SE cor . sec.





12-21-14aaa . Sample log of test hole in NE NE NE sec. 14, T 12 S, R 21 W , 60 feet west and 10 feet south of NE cor . sec. 14 ; drilled September 1960 . Altitude of land surface 2,338.5 feet; depth to wate 51.36 feet.







12-22-**6ddd.** Sample log of test hole in SE SE SE<br>sec. 6. T 12 S, R 22 W, 100 feet north and 20 feet west of SE cor, sec. 6, at edge of road; drilled October<br>1957. Altitude of land surface 2,399.1 feet; depth to water 91.09 feet.





12-22-12bbb. Sample log of test hole in NW NW<br>NW scc. 12, T 12 S, R 22 W, 350 feet south and 5 teet east of center of crossroads; drilled October 1957. Aiatude of land surface 2,340.4 feet.

> Thickness, Depth, feet feet

155

157



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Niobrara Chalk—Smoky Hill Chalk Member<br>Shale, brown and dark gray ............ 1 122

12-22-23dda.--Sample log of test hole in NE SE SE sec. 23, T 12 S, R 22 W, between highway and railroad, near fence line 90 feet north of highway; drilled October 1957. Altitude of land surface 2,366.8 feet.



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Niobrara Chalk - Smoky Hill Chalk Member Ochre <sup>2</sup> 123

 $^{12-25-22 \text{cbb}}$ . Drillers log of test hole in NW NW<br>SW sec 22, T 12 S, R 23 W, NW of pasture south of Erickson house; drilled July 1954 by Elmer Corder.





12-23-27ccc . -Sample log of test hole in SW SW SW sec. 27, T 12 S, R 23 W, 50 feet east and 10 feet north of center of crossroads ; drilled October 1957. Altitude of land surface 2,331.5 feet.

> Thickness, Depih, feet feet



12-23-28aad.——Drillers log of test hole in SE NF NE sec. 28, T 12 S, R 23 W , near 3rd telephone pole south of Dietz house; drilled July 1954 by Elme Corder.



1



Thickness, Depth,







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feet f





12-23-36bbc. Sample log of test hole in SW NW NW xc. 36 , T 12 S, R 23 W , 0.25 mile south of NW cor. sec. 36, at edge of road near driveway to field; augered July 1960. Altitude of land surface 2,337.7 feet; depth to water 18.5 feet. Thickness, Depih,



SEOGHVE Pleistocene



12-23-36cbc. ---------Sample log of test hole in SW NW SW sec. 36, T 12 S, R 23 W, 0.32 mile north of SW cor. sec. 36, 50 feet south of bridge, at edge of road; augered July 1960. Altitude of land surface 2,301.5 feet; depth to water 3.95 feet. Thickness, Depth,



- Niobrara Chalk—Smoky Hill Chalk Member Shale , blue gray 2.5 18.5
- $12.24.1$ dcc. $-$ -Drillers log of test hole in SW SW SE c. 1, T 12 S, R 24 W, 2.5 miles west of WaKeeney; drilled July 1954 by Elmer Corder .

Thickness, Depih, feet feet VEGEVE Pleistocene and Pliocene (undifferentiated) 5 Top , dark <sup>5</sup> 15 Clay , brown 10 Clay , sandy 11 26 Clay , gray 17 43 47 Limestone, soft  $\ldots$   $\frac{1}{2}$ Clay 16 63 56<br>56<br>71<br>75 66 Clay, sandy  $\ldots$   $\ldots$ Limestone <sup>1</sup> 67 70 Sand <sup>3</sup> W Limestone and sand <sup>1</sup> - A Limestone, cemented hard  $\ldots$   $\ldots$   $\ldots$ 80 % Sand , cemented hard <sup>5</sup> CRETACLOCS Cpper Cretaceous Niobrara Chalk - Smoky Hill Chalk Member Ochre 3

12-24-8ddc .-----------Sample log of test hole in SW SE SE sec. 8, T 12 S, R 24 W, 0.23 mile west of SE cor. sec. 8, at edge of road: drilled September 1957. Altitude of land surface 2,507.2 feet; depth to water 77.47 feet. Thickness, Depih,



12-24-9bcc. -Drillers log of test hole in SW SW NW sec .by sec. 9, T 12 S, R 24 W, drilled by Kansas Highway<br>Department. Depth to water 19.8 feet.<br>Thickness, Depth.



**12-24-13aaa.———**Sample log of test hole in NE NI NE scc. 13, T12 S, R24 W , 75 feet south and <sup>8</sup> feet west of NE cor . sec. 13 ; drilled September 1957. Al titude of land surface 2,462.1 feet ; depth to water 74.0 feet.



th,

	Thickness, Depth leet	teet
	-5	15
Pliocene		
Ogallala Formation		
Silt, very limy, light gray, and fine		
		18
Sand, fine to coarse, and tan silt	3522	23
Silt, very limy, sandy, white		25
Sand, medium to coarse		27
Sand and silt, very limy, cemented		
hard	13	40
Sand and silt, cemented very hard	3	43
Sand, medium to coarse, and gravel	5	48
	$\overline{\mathbf{c}}$	50
Sand, fine to coarse, silty	$\bar{2}$	52
Silt and sand, very limy, cemented	6	58
	2	60
	10	70
Sand, fine; contains thin lavers of		
		88
	18	
Sand, fine, and green silt	4	92
<b>CRETACEOUS</b>		
Upper Cretaceous		
Niobrara Chalk-Smoky Hill Chalk Member		
Shale, weathered, gray and white	כ	97
Shale, brown and dark gray	3	100

12-24-30bccl.---------Sample log of test hole in SW SW NW sec. 30, T12 S, R 24 W , 0.48 mile south of NW cor. sec. 30, at edge of road; drilled August 1960. Altitude of land surface 2,445.4 feet; depth to water 7.6 feet.



1**2-24-32ddd.————**Sample log of test hole in SE SE SF sec. 32, T 12 S, R 24 W, 50 feet north and 10 feet west of center of crossroads; drilled August 1960. Altitude of land surface 2,512.9 feet.





12-25-7ddd. Sample log of test hole in SE SE SE<br>sec. 7, T 12 S, R 25 W, 50 feet north and 10 fee

west of center of crossroads ; drilled August 1960 . Altitude of land surface 2,581.8 feet.



12-25-9add. -Drillers log of test hole in SE SENE c. 9, T 12 S, R 25 W; drilled by Kansas Highway Department. Altitude of land surface 2,544.1 feet. Thickness, Depih ,







12-25-9bcc. -Drillers log of test hole in SW SW VW sec. 9, T <sup>12</sup> S, <sup>R</sup> <sup>25</sup> <sup>W</sup> , drilled by Kansas High way Department. Altitude of land surface 2,496.5 feet. Thickness, Depih, feet<sup>-</sup>



12-25-10adc .---------------------- Drillers log of test hole in SW SE NE sec. 10, T 12 S, R 25 W; drilled by Kansas Highway Department. Altitude of land surface 2,539.2 feet. Thickness, Depth,<br>feet feet feet feet



12-25-15aaa. Sample log of test hole in NE NE<br>NE sec. 15. T 12 S, R 25 W, 20 feet south and 10 feet west of NE cor . sec . 15; drilled September 1957. Altitude of land surface 2,534.6 feet; depth to water 54.7 feet .

	Thickness, Depth. teer	teet
NFOGENE		
Pleistocene		
Peoria and Loveland formations		
	6	6
	8	14
Silt, reddish; contains limy streaks		15
Pliocene		
Ogaliala Formation		
	x	23
Silt, very limy, white, and fine sand	6	29
		40
Silt, sandy, very limy, light tan and		
pink	8	48
Silt, sandy, very limy, white; con-		
tains cemented streaks	8	56
Sand, medium to coarse	1	57
Sand, limy, cemented hard	2	59
Sand, fine, and very limy silt  11		70
Sand, medium to coarse, and very		
	8	78
<b><i>CRETACEOUS</i></b>		
<b>Upper Cretaceous</b>		

Niobrara Chalk -Smoky Hill Chalk Member



12-25-18bbb. ----------Sample log of test hole in NW NW NW sec. 18, T 12 S, R 25 W ; drilled September 1957. Altitude of land surface 2,583.5 feet; depth to water 57.68 feet.





12-25-24ddd. Sample log of test hole in SE SE SF sec. 24, T 12 S, R 25 W, 40 feet north and 10 feet west of center of crossroads ; drilled August 1960. Altitude of land surface 2,449.1 fect; depth to water 14.6 feet.





Thickness, Depih, *fcct* **CRETACEOUS** Upper Cretaceous Niobrara Chalk - Smoky Hill Chalk Member Shale, brown and gray 8 80 12-25-25ada. ----------Sample log of test hole in NE SE NE sec.  $25$ , T 12 S, R  $25$  W,  $0.\overline{3}$  mile south of NE cor. sec. 25 and 20 feet west of road center; drilled August 1960. Altitude of land surface 2,445.65 feet; depth to water 10.7 feet. Thickness, Depth, feet NEOGENE



12-25-31ddd. -Sample log of test hole in SE SE SE sec. 31 , T 12 S, R 25 W, 50 feet north and 10 feet west of center of crossroads; drilled August 1960. Altitude of land surface 2,568.2 feet.







13-21-8abb.———Drillers log of test hole in NW NW NE scc . 8, T 13 S, R21 W, drilled June 1961 by Bureau of Reclamation . Altitude of land surface 2,249,3 feet.  $\overline{a}$ .





Niobrara Chalk -





13-21-8dad . -Drillers log of test hole in SE NE SE sec. 8, T 13 S, R 21 W, drilled June 1961 by Bureau of Reclamation . Altitude of land surface 2,223.2 feet ; depth to water 43.0 feet.





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13-21-8ddd.——Sample log of test hole in SE SE SE Sec. 8, T 13 S, R21 W, 50 feet west and 10 feet north of center of crossroads ; drilled September 1960. Alti tude of land surface 2,205.9 feet; depth to water 28.2 .ut<br>.et



13-21-16bab . -Drillers log of test hole in NW NE VW sec . 16, T 13 S, R21 W, drilled July 1961 by Bureau of Reclamation . Altitude of land surface 2.213.4 feet: depth to water 49.7 feet.



Silt, sandy lower part, light brown ... 10<br>Silt, very sandy, light tan .................. 10 20 Sand, medium to very coarse, and fine gravel; upper part contains  $\lim y$  silt layers; lower 2 feet cemented <sup>8</sup> 28 **CRETACEOUS** Upper Cretaceous Niobrara Chalk—Smoky Hill Chalk Membe 40

Shale , gray 12

13-21-16cbc.———Sample log of test hole in SW NW SW sec. 16, T 13 S, R 21 W, 0.3 mile north of SW cor. sec. 16, 35 feet south of bridge, at edge of road; augered July 1960. Allitude of land surface 2,173.0



13-21-16dad.———Drillers log of test hole in SE NE SE sec. 16, T 13 S, R 21 W, drilled August 1961 by Bureau of Reclamation. Altitude of fand surface<br>2,190.3 feet; depth to water 29.8 feet.<br>Thickness, Depth,<br>feet for of Reclamation. Altitude of



Thickness, Depih,

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Thickness, Depth, feet Silt, sandy, brown ................................ 17 25 Sand, fine to medium, and chalk gravel 2 27 Sand, fine, silty 5 32 **CRETACEOUS** Upper Cretaceous Niobrara Chalk-Fort Hays Limestone Member Chalk, weathered, light yellow ...... 34 66 Carlile Shale Shale, very dark gray; contains 0.5 foot fine sand at top 3 69 13-21-16dbb.————Drillers log of test hole in NW NW SE sec. 16, T 13 S, R 21 W, drilled July 1961 by

Bureau of Reclamation . Altitude of land surface



13-21-16dda . -Drillers log of test hole inNE SE SE sec. 16, T 13 S, R 21 W, drilled August 1961 by Bureau of Reclamation . Altitude of land surface 2,187.3 feet; depth to water 26.0 fect.



13-21-17daa . -Sample log of test hole in NE NE SE sec. 17, T 13 S, R 21 W, 0.45 mile north of SE cor. sec. 17, 500 feet north of bridge, at edge of road; augered July 1960. Altitude of land surface 2,186.0 feet; depth to water 15.35 fcet .

64

Shale 2008. 2008. 2010.



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13-21-17dda. Sample log of test hole in NE SE SE<br>sec. 17, T 13 S, R 21 W, 0.18 mile north of SE cor.<br>sec. 17, 250 feet south of driveway to farm home, at cdge of road; augered July 1960. Altitude of land surface 2,193.3 feet; depth to water 20.1 feet.



- Niobrara Chalk—Smoky Hill Chalk Member<br>Shale, chalky, light blue gray ........ 3.5 43.5
- 13-21-20aad. Sample log of test hole in SENE NE sec. 20, T 13 S, R21 W, 0.2 mile south of NE cor . sec. 20, at edge of road; augered July 1960. Altitude of land surface 2,209.6 feet.



13-21-21bbb. ———Sample log of test hole in NW NW<br>NW sec. 21, T 13 S, R 21 W, 25 feet south and 8 fee east of center of crossroads; augered July 1960. Altitude of land surface 2,197.6 feet; depth to water 22.6 fect .





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Lpper Cretaceous

Vjobrara Chalk --Smoky Hill Chalk Member Chalk, gray to

13-21-31bbb. Sample log of test hole in NW NW<br>NW sec. 31, T 13 S, R 21 W, 25 feet east and 10 feet south of center of crossroads ; drilled September 1960. Altitude of land surface 2,360.0 feet; depth to water  $2$  5.5 feet.



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13-23-2bbb.--Sample log of test hole in NW NW NW sec. 2, T 13 S, R 23 W, 40 feet east of NW cor . sec. 2, at edge of road; drilled September 1960. Altitude of land surface 2,400.0 feet; depth to water 42.0 feet.





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13-25-19aaa. Sample log of test hole in NE NE NE<br>sec. 19, T 13 S, R 25 W, 30 feet west and 10 feet south of center of crossroads ; drilled September 1960. Altitude of land surface 2,450.4 feet.





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Original from UNIVERSITY OF MINNESOTA 14.21-28cbb. -Drillers log of observation well in NW NW SW sec.  $28$ , T 14 S, R  $21$  W, drilled March 1961 by Bureau of Reclamation . Altitude of land sur face 2,118.2 feet; depth to water 35.8 feet.



14-21-28ccb. Sample log of test hole in NW SW SW sec. 28, T 14 S, R 21 W (Leonard and Berry,  $1961$ , p.  $122$ ); jetted August 1950. Altitude of land surface 2,096.7 feet; depth to water 12.5 feet.

Thickness, Depih,

NEOGENE



14-21-30cdd. Sample log of test hole in SE SE SW sec.  $30$ , T 14 S, R 21 W (Leonard and Berry, 1961, p. 122): jetted July 1950. Altitude of p. 122); jetted July 1950. Altitude of land surface<br>2,065.4 feet; depth to water 13.08 feet.<br>Thickness, Depth,



14-21-30dad . -Sample log of test hole in SE NE SE sec. 30, T 14 S, R 21 W (Leonard and Berry , 1961 , D. 122 ) ; jctted September 1950. Altitude of land sur face 2,122.7 feet; dry hole.



14-21-30ddd. Sample log of test hole in SE SE SE<br>sec. 30, T 14 S, R 21 W (Leonard and Berry, 1961,<br>p. 122); jetted July 1950. Altitude of land surface 2,112.9 feet: dry hole.





NEOGENE Pleistocene

14-21-32add . -Sample log of test hole in SE SE NE sec. 32, T 14 S, R 21 W (Leonard and Berry, 1961, p. 123) ; jetted August 1950. Altitude of land surface 2,049.1 feet; depth to water 6.5 feet.



14-21-33bbc. Sample log of test hole in SW NW<br>NW sec. 33, T 14 S, R 21 W (Leonard and Berry, 1961, p. 123 ) ; jetted August 1950. Altitude of land surface 2,057.7 feet.



14-21-33bcc . -Drillers log of in **SW NW sec.** 33, T 14 S, R 21 W, 560 feet north of 1/2-mile line and 30 feet east of section line, drilled March 1961 by Bureau of Reclamation. Altitude of land surface 2,048.6 feet ; depth to water 7.9 feet.



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14-21-34ada.——Sample log of test hole in NE SE NF sec. 34, T 14 S, R 21 W, 0.35 mile south of NE cor. sec. 34, at edge of road (Leonard and Berry, 1961, p 124); drilled October 1949. Altitude of land surface 2,087.3 feet.

Shale, noncalcareous, blue black .... 3 37



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14-21-34dad. Sample log of test hole in SE NE SE sec.  $34$ , T 14 S, R 21 W, 0.3 mile north of SE cor.<br>sec.  $34$ , at edge of road (Leonard and Berry, 1961, p.124 ) ; drilled October 1949. Altitude of land sur face 2,041.4 feet; depth to water 16.9 feet.

> Thickness, Depih , feet feet



14-21-34ddd. Sample log of test hole in SE SE SE sec. 34, T 14 S, R 21 W, 300 feet north and 20 feet west of SE cor. sec. 34 (Leonard and Berry, 1961, p. 125 ) ; drilled October 1949. Altitude of land surface 2,028.4 feet.

feet Thickness, Depih , feet



14-21-35bcc . -Drillers log of observation well in Sw SW NW sec. 35, T 14 S, R 21 W, 185 feet north  $\rm _{of}$ 



Thickness, Depih,

feet

Thickness, Depih,

1. -mile line and 30 feet east of section line; drilled March <sup>1961</sup> by Bureau of Reclamation . Altitude of land surface 2,078.3 feet.



14-21-35cbcl . -Sample log of test hole in SW NW SW sec. <sup>35</sup> , T <sup>14</sup> S, <sup>R</sup> <sup>21</sup> W (Leonard and Berry , 1961, p. 125 ) ; jetted August 1950. Altitude of land surface 2,047.4 feet; dry hole. Thickness, Depth,<br>feet feet

NEOGENE Pleistocene Grand Island, Sappa, and Crete formations, undifferentiated (terrace deposits) Silt; contains very coarse sand ........ 1 1<br>Sand, very coarse, poorly sorted ...... 5 6 Sand, very coarse, poorly sorted ...... 5 CEETACEOUS L'pper Cretaceous Carlile Shale Shale, clayey, calcareous, dark gray; weathers tan 5 11

14-21-35cbc2 . -Drillers log of observation well in SW NW SW sec. 35, T 14 S, R 21 W, 1,130 feet south of 12-mile line and 30 feet east of section line ; drilled March 1961 by Bureau of Reclamation . Alti tude of land surface 2,039.9 feet; depth to water 15.5 feet.



14-21-35ccc. -Drillers log of observation well in SW SW SW sec. 35, T 14 S, R 21 W, 600 feet north and 30 fect cast of SW cor , sec .35 ; drilled March 1961 by of Reclamation. Altitude of Bureau of Reclamation. Altitude of land surface<br>2,035.2 feet; depth to water 11.4 feet.<br>Thickness, Depth,







14-21-36bbb . -Drillers log of observation well in NW NW NW sec. 36, T 14 S, R 21 W, at edge of road intersection; drilled March 1961 by Bureau Reclamation. Altitude of land surface 2,072.6 feet; depth to water 10.9 feet.



14-21-36bcc . -Drillers log of observation well in SW SW NW sec. 36, T 14 S, R 21 W, 20 feet north of  $\frac{1}{2}$ -mile line and 23 feet east of section line; drilled March 1961 by Bureau of Reclamation . Altitude of land surface 2,075.7 feet; depth to water 46.0 feet.



Thickness, Depth,  $f_{\text{ccl}}$ 

14-21-36ddd . -Drillers log of observation well in SE SE SE sec. 36, T 14 S, R 21 W ; drilled March 1961 by Bureau of Reclamation. Altitude of land surface 2,019.3 feet; depth to water 8.5 feet. Thickness, Depth,<br>feet feet **NEOGENE** Pleistocene Alluvium Silt , sandy, dark brown 2.7 2.7 5.3 ........... Sand, fine to medium, silty Sand, fine to coarse, light brown  $\ldots$  2 10 Sand, fine to coarse, grayish brown 19 29 **CRETACEOUS** Upper Cretaceous Carlile Shale Shale, blue gray 0.5 29.5  $\frac{8.0}{10}$ va 14-22-26aca. -Drillers log of test hole in NE SW NE sec. 26, T 14 S, R 22 W; drilled by Bureau of Reclamation. Altitude of land surface 2,162.0 feet. Thickness, Depih, feet feet NEOGENE Pleistocene Grand Island, Sappa, and Crete formations, undifferentiated (terrace deposits ) Silt , sandy 24 24 Sand, fine to medium 30 54 Silt , clayey 13 67 Sand; contains thin layers of clay  $\ldots$  29 96<br>Silt  $\ldots$  100 Silt 4 100 Sand, medium; contains pebbles .... 10 110 CRETACEOUS Upper Cretaceous Carlile Shale Shale <sup>3</sup> 113 14-22-33aac.——Sample log of test hole in SW NE NE sec. 33, T 14 S, R22 W, 0.2 mile SW of NE cor . sec. 33; augered July 1960. Depth to water 21.0 feet. Thickness, Depth, feet feel NEOGENE Pleistocene Grand Island, Sappa, and Crete formations, undifferentiated (terrace deposits ) Silt , sandy , tan 3,5 3.5 Silt , sandy , heavy , tan <sup>5</sup> 8.5 Silt and fine to medium sand .......... 6.5 15 Silt , heavy , brown 3.5 18.5 Sand , fine tomedium <sup>5</sup> 23.5 Sand , fine to coarse <sup>5</sup> 28.5 Sand, medium to coarse, very clean - 6.5 35 **CRETACEOUS** Upper Cretaceous Carlile Shale Shale, tough, blue gray 3.5 38.5 uwania улол 14-22-36aab.--Sample log of test hole in NW NE NE sec. 36, T 14 S, R22 W, 0.2 mile west of NE cor. sec. 36, at edge of road (Leonard and Berry, 1961, p.126 ) ; drilled October 1949. Altitude of land sur face 2,114.7 feet. Thickness, Depih, feel feel NEOGENE Pleistocene Grand Island, Sappa, and Crete formations, undifferentiated (terrace deposits) Silt, dark brown; contains fine to



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NEOGENE

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14-22-36dad.——Sample log of test hole in SE NE SF c. 36, T 14 S, R 22 W (Leonard and Berry, 1961, p. 127); jetted September 1950. Altitude of land surface  $2,068.3$  feet; depth to water  $6.5$  feet.

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coarse gravel and some coarsesand <sup>1</sup>



14-22-36dda . -Sample log of test hole in NE SE SE sec.  $36$ , T 14 S, R 22 W (Leonard and Berry, 1961, p. 127): jetted September 1950. Altitude of land sur face  $2,063.4$  feet; depth to water  $3.85$  feet.



14-23-28ccc. Sample log of test hole in SW SW SW SW SW SW SW SW SW SW, 35 feet east and 10 feet north of SW cor. sec. 28; augered July 1960. Altitude of land surface 2,218.2 feet.  $\sim$  . . .



14-23-32ada. Sample log of test hole in NE SE NE  $\frac{1}{2}$  c. 32, T 14 S, R 23 W, 0.33 mile south of NE cor. sec. 32, at edge of road; augered July 1960. Altitud of land surface 2,200.2 feet.



sec. 32, T 14 S, R 23 W, 0.6 mile south of NE cor.<br>sec. 32, at edge of road; augered July 1960. Altitude of land surface 2,164.9 feet; depth to water 17.05 feet. Thickness, Depih, feet feet NEOGENE Pleistocene Alluvium Silt , sandy , light tan 8.5 8.5 Silt , sandy , tan 1.5 10 brown and gray <sup>23</sup> نم مه Sand , fine 18 Silt , heavy , dark brown 20 Silt, heavy, limy, رد **CRETACEOUS** Upper Cretaceous Carlile Shale Shale , dark gray <sup>7</sup> 30 14-23-32dad. Sample log of test hole in SE NE SE<br>sec. 32, T 14 S, R 23 W, 0.65 mile south of NE cor. sec. 32, 250 feet north of bridge and 35 feet west of road center; drilled September 1960. Altitude of land surface 2,161.8 feet. Thickness, Dep! h, feet feet NEOGENE Pleistocene Alluvium Sand, very fine to medium, red  $tan \frac{3.5}{5.5}$  3.5 Silt, heavy, greenish brown  $\ldots$   $\ldots$  6.5 10 Sand, very fine to fine,  $\tan$   $\ldots$   $8.5$  18.5 Sand, fine to coarse; contains small amount fine gravel <sup>5</sup> 23.5 Sand, fine to very coarse; contains small amount chalk pebbles ......... 6.5 30<br>and, medium to very coarse .......... 5 35 Sand ,medium to very coarse <sup>5</sup> 35 3,5  $6.5$ <br> $8.5$ **CRETACEOUS** Upper Cretaceous Carlile Shale 8.5 – 8.5 Shale, chalky, dark gray .................. 8.5 – 43.5 14-23-33cbc . -Sample log of test hole in SW NW SW sec. 33, T 14 S, R23 W, 0.3 mile north of SW cor. sec. 33, 50 feet north of bridge, and 30 feet east

14-23-32daa. Sample log of test hole in NE NE SE<br>sec. 32, T 14 S, R 23 W, 0.6 mile south of NE cor.

of center of highway; augered July 1960. Altitude of land surface 2,158.9 feet; depth to water 12.6 feet. Thickness, Depth, feet feet NEOGENE Pleistocene Alluvium Silt , sandy , tan 3.5 Silt , sandy , brown 8.5 Silt, very sandy, tan brown  $\dots\dots\dots\dots\dots$  7.5 16 Silt, clayey, sandy, tough,  $\tan$   $\frac{3}{2}$   $\frac{19}{2}$ Silt , sandy , loose 27 Silt , clayey , tough , tan brown 38 اهمادج **CRETACEOUS** Upper Cretaceous Carlile Shale Shale , blue gray 10.5 48.5  $\frac{3.5}{8.5}$  $\frac{3.5}{5}$ 

14-24-27dda.--Sample log of test hole inNE SE SE sec. 27, T 14 S, R 24 W, 0.23 mile north of SE cor. sec. 27, at edge of road: augered July 1960. Altitud of land surface 2,252.5 feet.

NEOGENE Pleistocene

Grand Island, Sappa, and Crete formations, undifferentiated (terrace deposits )




14-24-27ddd. Sample log of test hole in SE SE SE sec. 27, T 14 S, R 24 W, 50 feet north and 15 feet west of center of crossroads; augered July 1960. Alti tude of land surface 2,250.2 feet. Thickness, Depth,



14-24-29bcb.——Sample log of test hole in NW SW NW sec. 29, T 14 S, R 24 W, 650 feet north of river , at edge of road ; drilled October 1957. Altitude of land surface 2,205.7 feet; depth to water 5.94 feet.



14-24-29bcc.——Sample log of test hole in SW SW NW sec. 29, T 14 S, R24 W, 50 feet north of river , at edge of road; drilled October 1957. Altitude of land surface  $2,205.3$  feet; depth to water  $5.93$  feet. Thickness, Depih,



Thickness, Depth,<br>fect fect feet **f** 



14-24-30bbc. ---------Sample log of test hole in SW NW NW sec . 30, T 14 S, R 24 W; drilled August 1947. Altitude of land surface 2,213.4 feet; depth to water 9.0 feet.



14-24-30bcb.——Sample log of test hole in NW SV NW sec .30, T 14 S, R24 W; drilled August 1947. Altitude of land surface 2,212.6 feet; depth to water 6.2 feet.

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14-24-30bcc. Sample log of test hole in SW SW NW sec. <sup>30</sup> , T <sup>14</sup> S, <sup>R</sup> <sup>24</sup> W ; drilled August 1947. Altitude of land surface 2,211.6 feet; depth to water 5.1 feet. Thickness, Depth,



14-24-30ccc.----------Sample log of test hole in SW SW SW sec. 30, T 14 S, R 24 W; drilled August 1947. Altitude of land surface 2,275.9 feet. Thickness, Depth,



14-24-31 cbb. Sample log of test hole in NW NW SW sec. 31, T 14 S, R 24 W; drilled August 1947. Alutude of land surface 2,288.6 feet.



14-24-34aaa.——Sample log of test hole in NE NE NE sec. 34, T 14 S, R 24 W, 0.1 mile south of NE cor. sec. 34, at edge of road; augered July 1960. Altitud of land surface 2,234.3 feet .

NEOGENE

- Pleistocene Grand Island, Sappa, and Crete formations,
	- undifferentiated (terrace deposits)

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14-24-34ddd. Sample log of test hole in SE SE SE sec. 34, T 14 S, R 24 W, 50 feet north of SE cor. sec. 34, at edge of road; augered July 1960. Altitude of land surface 2,195.3 feet. Thickness, Depih,



14-24-35bcb. Sample log of test hole in NW SW NW sec.  $35$ , T 14 S, R 24 W, 0.3 mile south of NW cor.  $sec. 35$ , at eage of road; augered july 1900. Although tude of land surface  $2,200.4$  feet; depth to water 35, at edge of road; augered July 1960. Alti-14.37 feet.



**14-24-35bcc.**——Sample log of test hole in SW SW NW sec. 35, T 14 S, R 24 W, 0.48 mile south of NW cor. sec. 35, at edge of road: drilled September 1960. Altitude of land surface 2,182.2 feet; depth to water 7.8 feet .



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14-24-35cbc. -Sample log of test hole in SW NW SW sec. 35, T 14 S, R 24 W, 0.3 mile north of SW cor. sec. 35, at edge of road across from farm home; augered July 1960. Altitude of land surface 2,182.3 feet. Thickness,



14-25-9ddd. Sample log of test hole in SE SE SE sec. 9, T 14 S, R 25 W, 150 feet north of SE cor. sec.<br>9, at edge of road; augered July 1960. Altitude of land surface 2,273,5 fect,



**14-25-15bbb.———**-Sample log of test hole in NW NW NW sec. 15, T 14 S, R 25 W, 500 feet south of NW cor. sec. 15, at edge of road: augered July 1960. Altitude of land surface 2,271.6 feet.



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east of center of crossroads (Leonard and Berry , 1961 ,

p. 148 ) ; drilled October 1949. Altitude of land sur face 2,031.4 feet. Thickness, Depih,



15-21-3add . -Sample log of test hole in SE SE NE sec. 3. T <sup>15</sup> S, <sup>R</sup> <sup>21</sup> W , 0.45 mile south of NE cor . sec. 3, at edge of road (Leonard and Berry, 1961, p. 148) ; drilled October 1949. Altitude of land surface 2,100.7 feet. Thickness, Depth,



15-21-3ddd. Sample log of test hole in SE SE SE Sand, fine to medium, clean .................................. sec. 3, T 15 S, R 21 W, 100 feet north and <sup>7</sup> feet west of center of crossroads (Leonard and Berry, 1961, p.148 ) :drilled October 1949. Altitude of land sur face 2,148.8 feet .



15-22-labb . -Sample log of test hole in NW NW NE sec. 1, T 15 S, R 22 W (Leonard and Berry, 1961, p. 149) ; jetted July 1950. Altitude of land surface 2,068.6 feet; depth to water 8.3 feet. Thickness, Depth,



15-23-5aaa.——Sample log of test hole in NE NE NI sec. 5, T 15 S, R 23 W , 50 feet west and 10 feet south of NE cor. sec. 5; augered July 1960. Altitude of land surface 2,158.5 feet; depth to water 10.57 feet.



Thickness, Depih,



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R 24 W R 25 W  $\frac{2322}{25}$ 2435  $30<sub>0</sub>$ Railroad Collyer Union Pacific  $101$ U.S. 40  $\frac{2526}{50}$  $Nn$ 39°00'  $Nal$  $2320$  $N<sub>t</sub>$ Creek  $2243$  $Nt$ N  $Na$ Kns  $38°45'$ Kns No  $\frac{2472}{25}$ Base and drainage compiled from maps prepared by the Soil Conservation Service

# **MAP OF TREGO COUNTY, KANSAS**<br>showing areal geology, water-table contours, and location<br>of wells and test holes for which records are given

# By Warren G. Hodson

1965



# **EXPLANATION**



Scale in miles

 $T$ 

 $\mathsf{S}$ 

 $T$ 

 $S$ 

Areal geology mapped by Warren G. Hodson in 1959









 $39°00'$ 

 $2396$ 

Nal

 $Nal$ 

 $N!$ 

 $38^{\circ}45'$ 

**No** 

Base and drainage compiled from maps<br>prepared by the Soil Conservation Service

 $Nal$ 

Kns

 $2320$ 

Kns

Npl

Ao.

R 25 W

Collyer Union

Pacific U.S. 40

R 24 W

2322

Railroad

 $2279$ 

 $37^\circ$ 

 $Knf$ 

2214

 $\frac{2338}{29}$ 

 $\frac{2472}{25}$ 

2207

Creek)

Nal

 $\frac{2243}{44}$ 

272

2391



# **EXPLANATION**



