

Karst-Derived
Early Pennsylvanian
Conglomerate
in Ness County, Kansas

Doris E. Nodine-Zeller

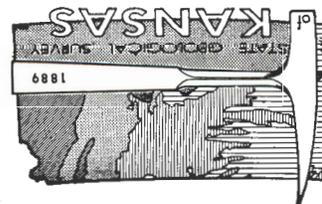
Kansas Geological Survey
Bulletin 222 1981

Karst-Derived Early Pennsylvanian Conglomerate
in Ness County, Kansas:
Subsurface Mississippian-Pennsylvanian Boundary
Delineated in Well Core

By

Doris E. Nodine-Zeller

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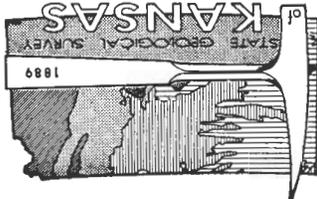
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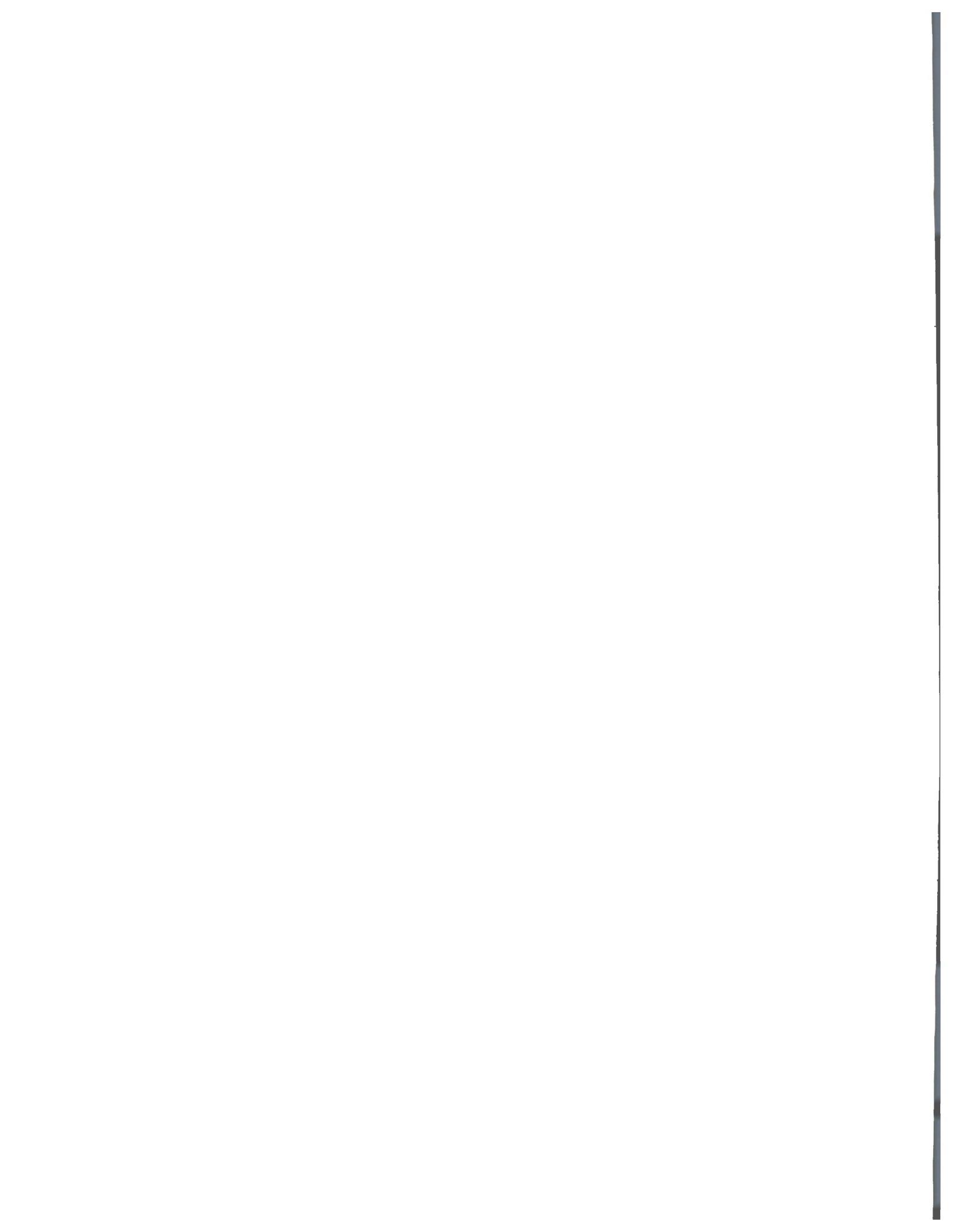
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EXECUTIVE SUMMARY

This Bulletin reports on basic research in the field of micropaleontology, the study of the microscopic fossils found in rocks. By studying these tiny fossils, much information can be obtained about the present characteristics of their host rocks and about the conditions under which those rocks were formed. Various fossils were studied to determine the age and to interpret the succession of events in the formation of host rocks. These events can be compared with those that took place in Texas, Oklahoma, Missouri, Kansas, Colorado, Wyoming, Montana, and the southwestern United States during about the same time and involving the same rock units. For this report, the rocks in a continuous core from the Mid-Continent No. 1 J. G. Collins borehole, Ness County, Kansas, were studied to determine their age and other characteristics. It is important to study this rock horizon as it may contain lead, zinc, and silver ores or be a stratigraphic trap for petroleum and natural gas.

Key Words: Atokan; breccia; Cherokee; conglomerate; Desmoinesian; endothyrids; Fort Scott; fusulmids; Kansas; karst; Marmaton; microfossils; Mississippian; Mississippian-Pennsylvanian boundary; Morrowan; Ness County, Kansas; Osagean; paleosol; Pennsylvanian; southwestern Kansas subsurface.

Contents

ABSTRACT	1
INTRODUCTION	2
GEOLOGIC SETTING	2
PREVIOUS INVESTIGATIONS	2
DESCRIPTION OF CORE	3
INSOLUBLE RESIDUES	19
MICROFAUNA IN CORE	19
INTERPRETATION OF CORE	26
PALEOGEOGRAPHIC AND DEPOSITIONAL HISTORY	26
CORRELATIONS USING ELECTRIC LOGS	28
DISCUSSION	28
ACKNOWLEDGMENTS	29
REFERENCES	29

Figures

1. Location of Mid-Continent No. 1 Collins borehole, Ness County, Kansas	3
2. Geologic section encountered in Mid-Continent No. 1 Collins core	5
3. Fort Scott Limestone core piece	6
4. Thin sections of Fort Scott Limestone	7
5. Thin sections of lithologic types occurring in basal Marmaton Fort Scott Limestone	9
6. Thin sections of lithologic types found in Mid-Continent No. 1 Collins well core	10
7. Beds from Mid-Continent No. 1 Collins well core from paleosol and immediately below	13

ERRATA

Kansas Geological Survey
Bulletin 222, 1981

The following scales were omitted:

Plate 2, page 23
Scale for 16 and 17 shown above 17
————— (1.0 mm)

Plate 3, page 25
Scale for remaining specimens shown in 6
————— (0.5 mm)

Plates

8. Matrix sandstone from breccia-conglomerate	14
9. Beds from Mid-Continent No. 1 Collins well core	15
10. Fractured and weathered brownish-gray Mississippian limestone cobble in breccia-conglomerate	16
11. Matrix conglomerate bed composed of residual materials	16
12. Large fragments of fossilized wood in matrix sandstone	16
13. Matrix bed consisting largely of sand and silt grains	16
14. Matrix bed	17
15. Shaly matrix bed	17
16. Thin sections of lithologic types found in cobbles and boulders in Mid-Continent No. 1 Collins well core	18
17. Dasycladacean alga from Chesteran limestone cobble	19

Tables

1. Microfossils in thin sections of core from the Mid-Continent No. 1 Collins well core	21
2. Microfossils in thin sections of core from the Mid-Continent No. 1 Collins well core	23
3. Microfossils in thin sections of core from the Mid-Continent No. 1 Collins well core	25
4. Correlation cross section A-A' showing relationship of Mississippian and Pennsylvanian units in Mid-Continent No. 1 Collins borehole (in pocket)	25
1. Lithologic description of the Mid-Continent No. 1 J. G. Collins Core	6

Karst-Derived Early Pennsylvanian Conglomerate in Ness County, Kansas: Subsurface Mississippian-Pennsylvanian Boundary Delineated in Well Core

ABSTRACT

Beds encountered in a continuous core from the Mid-Continent No. 1 J. G. Collins borehole, Ness County, Kansas, include a Pennsylvanian breccia-conglomerate of Atokan or lowermost Desmoinesian age, with reworked upper Morrowan or Atokan limestone in a conglomerate or possibly a deeply weathered subsoil zone; a paleosol; a pebble conglomerate of Cherokee age; Cherokee shale representing transgressive-regressive marine episodes; and a normal marine limestone (Fort Scott) of early Permian age. Beds at the bottom of the core rest upon the upper, weathered surface of the Mississippian limestone (Osagean). The breccia-conglomerate contains a mixture of weathered-in-place cobbles and boulders and riverine pebbles and cobbles of Osagean, Meramecian, and Chesteran ages derived in part from an extensive, deeply weathered karst terrain. The chaotic nature of the matrix beds (composed mainly of fresh and white chalky chert; spicular chert; drusy and doubly terminated quartz crystals; white beakite rings; rounded, frosted and clear, angular sand and silt grains; gray and green clay residuum, and minor amounts of

sphalerite and chalcopyrite) represents the mixed weathering products derived from reworking of Mississippian beds in the matrix of the breccia-conglomerate. Also in the enclosing matrix are beds at several levels with carbonized plant remains and coaly fragments of Morrowan or Atokan age and a disjunctive white sandstone conglomerate below the weathered soil zone. A cross section, with base of the Fort Scott Limestone as datum, shows correlation of three of the overlying units above in the vicinity of the No. 1 Collins borehole. Vari-ation in thickness of the "basal Pennsylvanian conglomerate" leads to difficulty in picking the top of the "Mississippi solid." Fusulines, endothyrids, archaedisoids, conodonts, and algae were used for age determination and to interpret succession of stratigraphic, geomorphic, and structural events. These karst-related events can be compared with those that took place in Texas, Oklahoma, Missouri, Kansas, Colorado, Wyoming, Montana, and the southwestern United States during about the same time and that involved the same stratigraphic units. This Mississippian karst horizon is important as a host rock for the emplacement of lead, zinc, and silver ores and as a stratigraphic trap in the accumulation of petroleum and natural gas.

County area to the south. A contour map on top of the Mississippian (Merriam, 1960) indicated regional structural trends in this same direction parallel to, but separated from, the Central Kansas uplift. In addition, plotting of the oil fields in Gove, Lane, Ness, and Hodgeman counties reveals the outline of this trend (Kansas Geological Survey, 1975). This same structural grain seems to have been present from Precambrian through Mississippian and even into earliest Pennsylvanian time.

Subsurface Paleozoic rocks previously reported as present in this area include those of Ordovician age (Arbuckle, Viola), unconformably overlain by the "Kinderhook shale" (Chatanooga, Devonian), and the Gilmore City Limestone (Kinderhookian), unnamed Osagean beds, and the "Warsaw Limestone" (Meramecian) of Mississippian age. Rocks of Pennsylvanian age (Desmoinesian, Missourian, and Virgilian stages) unconformably overlie truncated Mississippian strata. Permian beds rest disconformably upon those of the Pennsylvanian Wabamoose Group. To the east, toward the Central Kansas uplift, younger Pennsylvanian strata successively overlap the older, deeply eroded Paleozoic beds, until, at the crest of the structure, late Pennsylvanian rocks rest directly on Precambrian granites and quartz monzonites. Of the Pennsylvanian beds, only rocks of possible Morrowan or Atokan stages, undifferentiated Cherokee Group, and basal portion of the Marmaton Group (Fort Scott Limestone) have been included in this report.

PREVIOUS INVESTIGATIONS

Thompson and Goebel (1963) included a study of the No. 1 Collins core in their investigation of conodonts in subsurface rocks of Meramecian age in Kansas. In regard to the age of the Mississippian portion of the core, they stated (p. 14):

Correlation using conodonts places the top of the Warsaw at a depth of 4,468 feet. From sample logs the top of the Warsaw is picked on a lithologic basis at 4,527 feet. The depth of 4,527 feet is a level in the core where fine-strate bryozoans are the dominant constituent of the core. The lack of agreement between the lithologic pick and the conodont specimens collected from the No. 1 Collins well led to the several reexaminations of the core and careful redescription. It was observed that no conodonts were recovered from the shaly, clayey, sandy sections of this well. Conodont specimens recovered were apparently from the carbonate chinks of rock. In the section from 4,480 to 4,495 feet, the most significant fossil to the Warsaw Limestone, *Gnathodus texanus*, was found in abundance.

Thompson and Goebel (1968 [1969]) pointed out that the subsurface character of the Warsaw Lime-

The Mid-Continent No. 1 J. G. Collins, C NW NW Sec. 24 T.20S, R.26W, Ness County, Kansas (Fig. 1), was recorded as a dry hole in 1933. This report is based upon beds encountered in a continuous core taken from this borehole from 4297 feet (1310 m) (drillers' Fort Scott Limestone, Marmaton Group, Desmoinesian Stage, Middle Pennsylvanian Series) to 4555 feet (1388 m), reported simply as "Mississippian."

This study is part of a larger on-going project to investigate the Mississippian-Pennsylvanian boundary in southwestern Kansas and to delineate the ages and nature of Mississippian and Pennsylvanian beds encountered in well cores in this area. Specifically, the present study was done to compare age determinations derived by using endothyrid and fusulinid formations with those using conodonts in an earlier study undertaken by Thompson and Goebel (1963; 1968 [1969]) on the No. 1 Collins well core.

No electric or lithology logs were made for this other boreholes in the vicinity of the No. 1 Collins have been studied and correlated with the No. 1 Collins. Approximately half of the core and core pieces had been digested in acetic acid for Thompson and Goebel's studies, and the insoluble residues were described by Welch (1963). The remaining half of breccia-conglomerate pieces and matrix was examined individually for macro- and microfossils. The only fossils visible in the matrix sandstone, clay, and shale of the core were coalified leaves and finely comminuted carbonaceous coaly debris at several intervals between 4496 and 4527 feet (1370 and 1380 m). Thin sections were made of individual pieces of limestone and chert included in the breccia-conglomerate that seemed productive of foraminifers, and the microfossils were identified in the hope that these might at least define the age of the included pebbles, cobbles, and boulders. The "Mississippian" clasts in the core yielded diagnostic endothyrid foraminifera ranging in age from Meramecian to Chesterian; and a diverse microfauna, including fusulinids and algae, was found in Pennsylvanian (Cherokee and Marmaton) rocks at 4430 feet (1350 m) and upwards.

GEOLOGIC SETTING

The No. 1 Collins borehole is situated 48 miles (77 km) west-southwest of the western paleoerosional margin of the Central Kansas uplift (Fig. 1). Schierling (1968) showed a northwest-southeast-trending anticline in Gove County, indicating a positive structure that in all probability extends into the Ness

INTRODUCTION

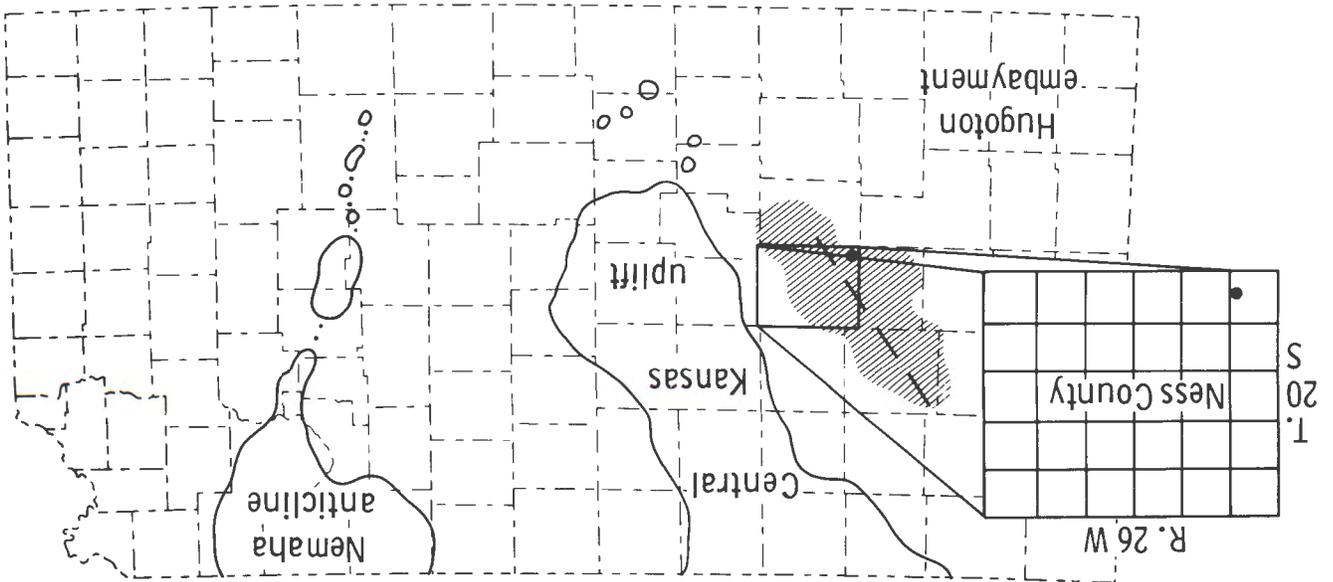


FIGURE 1. Location of Mid-Continent No. 1 Collins borehole, T. 20S, R. 26W, Ness County, Kansas, in relation to Central Kansas uplift. (After fig. 6, Zeller, 1968.)

Genevieve) or are confined solely to the Warsaw dolomite, containing varying amounts of gray, mottled, opaque, microfossiliferous chert, locally occurring geodic and drusy quartz, and disseminated pyrite and glauconite. Molds of fossils are common. Residues from the dolomite beds sometimes contain spongy masses of spicules. They further stated (p. 9):

DESCRIPTION OF CORE

A detailed description of the No. 1 Collins core (Table 1) includes types of fossils found in the rocks, with generic and specific names, where ascertainable. In analyzing the core, both the drillers' log and the core record were utilized in order to augment the visual, microscopic, and thin-section examination of the rocks remaining in the core boxes. Much unconsolidated shale, mudstone, and sandstone was lost in the core recovery operation. Core record descriptions such as "4427 Green shale with chunks of dense Pennsylv. ls. imbedded," "4493.5 Limestone broken by diagonal green clay seams," "4527 Hard gray siltstone, very sandy, it contains small pieces and larger chunks of limestone, green clay, white chert and miscellaneous materials such as black plant fragments," and at 4505, 4548, and 4555 (TD) "Soft gray clay" verified the interpretation that the bulk of the core was a breccia-conglomerate with a variable matrix of detrital materials. When examined microscopically, all the limestone core pieces were found to contain green

In cores from Kiowa and Hodgeman counties . . . , the Warsaw Limestone lies directly below Pennsylvanian rocks and is made up of conglomeratic, sandy, dolomitic shale, containing chert. Post-depositional and post-chert fracturing and faulting is exhibited in the Warsaw Limestone in Hodgeman County. In Ness County, in addition to "mixed-fossil" fragmental limestone . . . , the Warsaw Limestone is breccia with large fragments of finely crystalline limestone interspersed with green, silty, clayey shale. Fowler and Robbe . . . defined this lithology [in Ireland] as "quasi-breccia." Thompson and Goebel . . . applied the term to the Warsaw Limestone in Core F [No. 1 Collins].

Thompson and Goebel (1968 [1969]) concluded, on the basis of conodont evidence, that the interval in the core that I have defined as Atokan or lower Desmoinesian breccia-conglomerate was totally Warsaw in age. However, all of the conodonts listed by them as present in the core (appendix, core F, p. 49) in the "Warsaw" of the No. 1 Collins (4450-4555 ft.) either range from Warsaw through St. Louis (or Ste.

stone; short stringers and pods of glauconite; heterogeneous mixtures of quartz grading from silt to sand size, most frosted and rounded, some clear and angular; intraclasts in the form of short stringers and clunks of glauconitic silt; stringers of sandstone and sandstone clasts; and, in a few places, calcitic or siliceiferous fossil debris. Beds are variable in composition and change abruptly in this unit. One one-inch bed is a coarse white sandstone spotted with iron stain and with dark siltstone-mudstone stringers. Another is a white siliceiferous sandstone one inch thick, fossiliferous, and in places extremely highly weathered and tripolitic (see Fig. 6, D). The white sandstone contains cobbles and pebbles of Mississippian limestone. It is recognized by J. D. Davies as a marker bed on the lithologic logs prepared for the Kansas Geological Society.

The interval from 4447 to 4436 feet (1355-1352 m) represents the deeply weathered soil profile (Fig. 2). It contains, ascending, a red to greenish-gray to deep purple, unctuous claystone, containing shrinkage cracks that have been re cemented with green clay and calcite veins (Fig. 7, B). Above is a gray, noncalcareous, fossiliferous sandstone, with intraclasts of finer-grained, thinly laminated shaly sandstone with sub-bituminous coal debris, sphaerite, pockets of mud, and pieces of godes with "dead" oil. On top of that are "rotten" mudstones and siltstones, just below the paleosol, with iron in differing states of oxidation, from greens of glauconite through bright yellow to bright red to purple.

In this upper interval are pieces of highly weathered limestone, yellow and red with limonite and hematite, and shot through with calcite veins (Fig. 7, A). It is in these limestone cobbles between 4435 and 4440 that the Pennsylvanian conodont *Idiogmathodus* was found. It is not clear whether these limestones were weathered in place or were reworked Pennsylvanian (Morrowan?) cobbles as described in the core record at 4427 feet (see Table 1). Any calcareous microfossils in these limestones would not have survived the intense weathering in the soil zone. In their absence, and because *Idiogmathodus* is non-age-diagnostic unless identified to species, no accurate age could be determined for these particular limestone cobbles, except that they appear to be early Pennsylvanian. Beds from 4436 to 4431 feet (1352-1351 m) represent the upper part of the soil zone. The only lithologic descriptions of this interval (Table 1) were obtained from the drillers' log and the geologists' report as there was no core recovery, only finely divided earthy, sandy debris. They recorded limestone with a

clay and/or silt or sand in tiny crevices and irregularities on the outside, top, or bottom of the piece. In addition, thin-section examination of a "Warsaw" cobble at 4475 feet (1364 m) (Table 1) revealed much interstitial glauconite that, when completely weathered, would have contributed to the volume of green clay shale encountered in the core interval. These observations led to the belief that no solidly bedded Mississippian limestone was encountered by the coring, even to the total depth at 4555 feet (1388 m). In the lithologic documentation of the core (Table 1) some of the footage described by the geologists and driller when the samples came up, and which is no longer intact or was lost in coring, is inserted where it added important information on the nature of the beds.

Beds from the base of the core to 4475 feet (1364 m) (Fig. 2) comprise the distinctive main unit of the breccia-conglomerate, with pebbles, cobbles, and boulders of Mississippian limestones of mixed ages in varying states of weathering from fresh, with well-preserved fossils, to slightly weathered and oxidized, to completely leached, decalcified, and "rotten." Large amounts of chert, occurring as pebbles and cobbles and smaller pieces grading down to millimeter-sized fragments, are found throughout this interval. Some chert fragments are fresh and contain well-preserved endohybrid foraminifers (see Fig. 16, A), sponge spicules, or other fossils. Most pieces are moderately or severely weathered and devitrified to a chalky white composition. The matrix beds enclosing the limestone and chert pieces are predominantly light and dark gray, gray-green and green glauconitic shale, clay, or mudstone; gray, gray-green, and green (glauconitic) siltstone; and greenish-buff, gray, and gray-green (glauconitic) sandstone. The matrix beds are mostly friable, noncemented, nonindurated, and noncalcareous. They contain scattered pyrite; sphaerite; chalcopyrite; geodes; mammillary quartz; doubly terminated quartz crystals; beekite rings; drusy quartz; glauconite grains; coarse, rounded quartz grains; aggregates of pyritized spicules; plant and coaly debris; and rounded, frosted and clear, and subangular to angular sand and silt grains.

Beds from approximately 4447 to 4475 feet (1355-1364 m) are a distinctive white conglomeratic sandstone (Fig. 2), contrasting sharply with the beds below and above in color and composition. The color varies from white to light tan to pinkish-white, to pale gray and dark gray. The sandstone contains much glauconite, scattered pyrite, sphaerite, and mica; beekite rings; thinly laminated, angular intraclasts of pale greenish-gray (glauconitic) sandstone and gray sand-

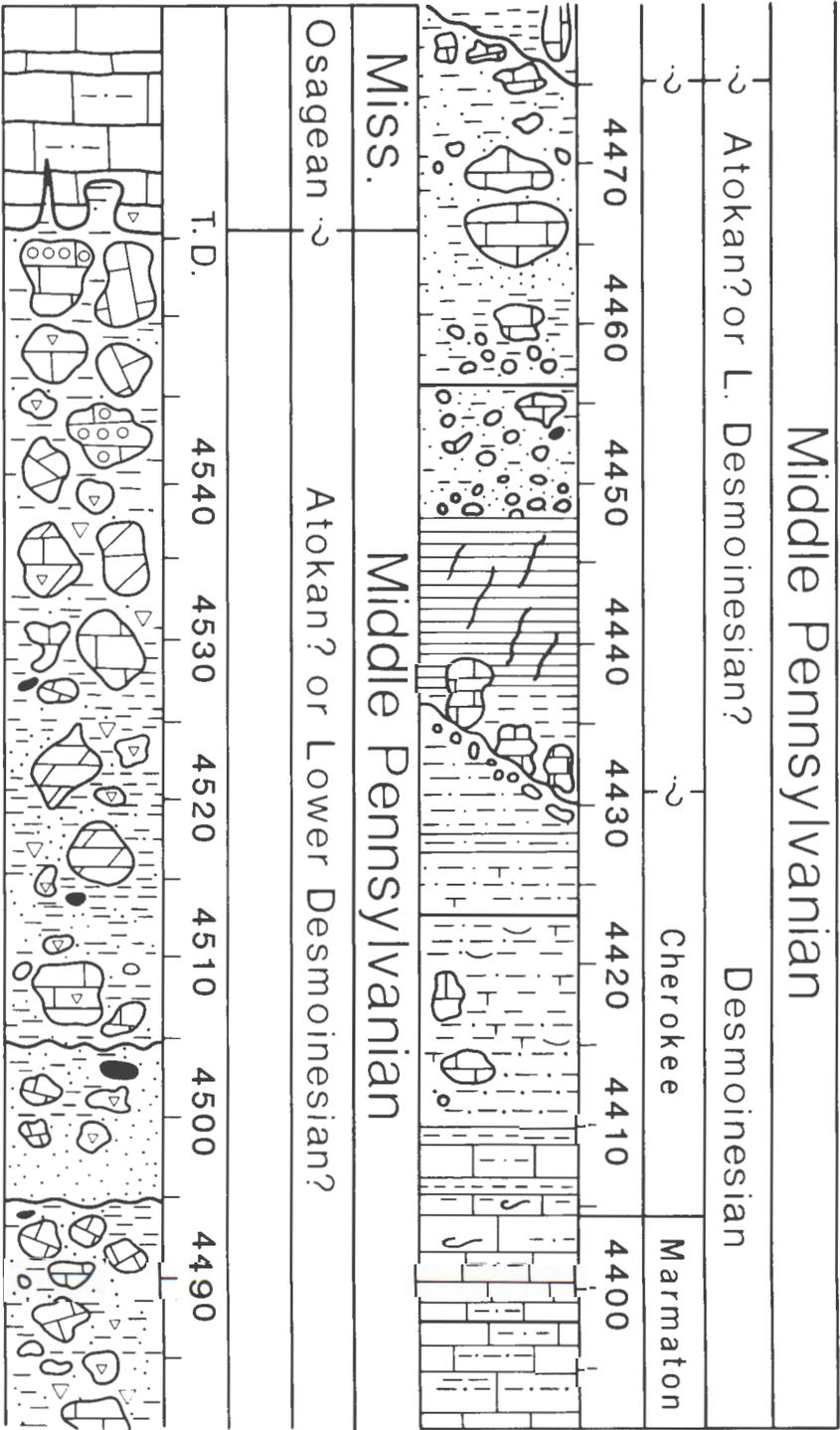


Figure 2. Geologic section encountered in Mid-Continent No. 1 Collins core, C NW NW Sec. 24, T.20S, R.26W, Ness County, Kansas. Conglomerate derived from karst-surface debris lies on top of deeply weathered late Osagean limestones and dolomites. Distinctive white sandstone conglomerate is shown from 4475 to 4447 feet. Paleosol lies at 4430 feet and below. Marine Cherokee and Marmaton beds lie above the basal Cherokee conglomerate.

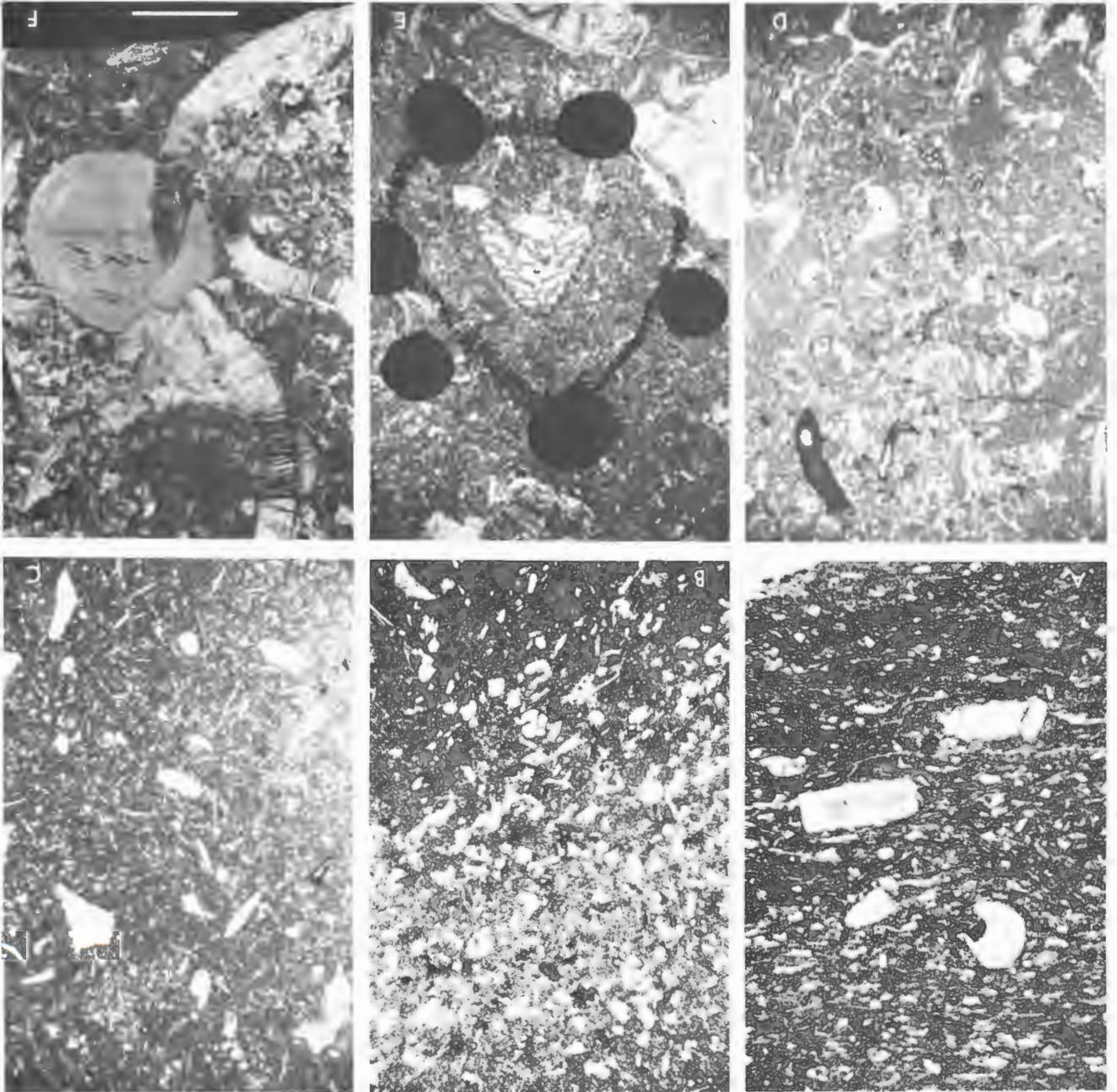


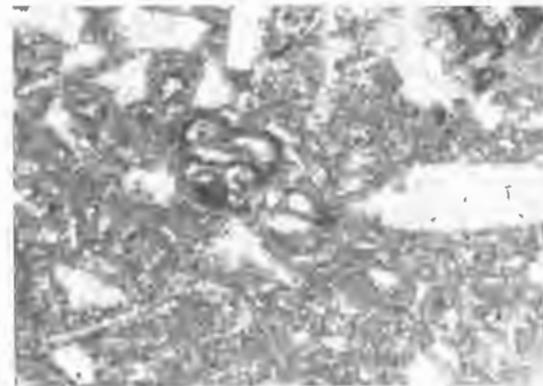
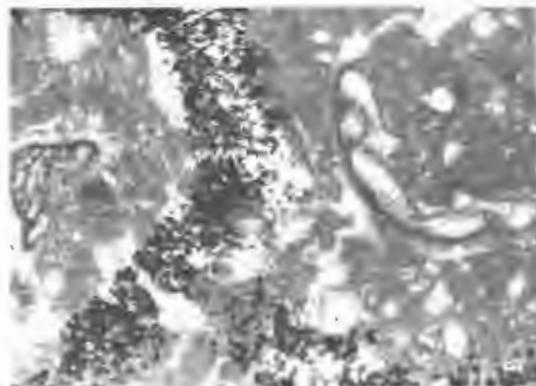
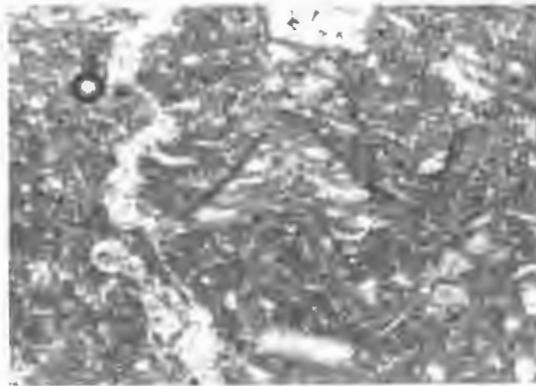
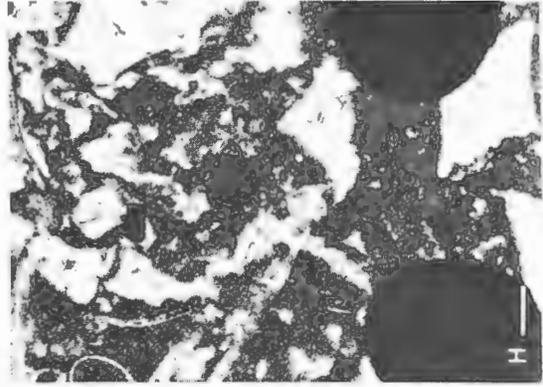
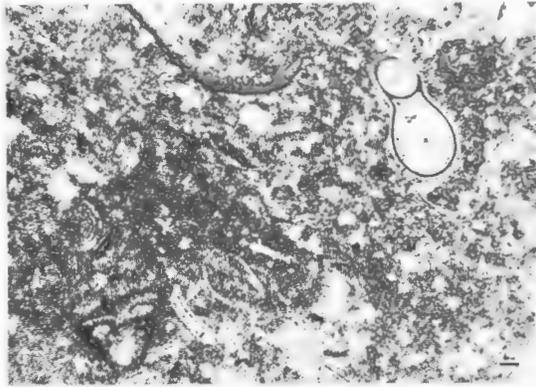
FIGURE 4. Thin sections of Fort Scott Limestone, Marmaton Group, Desmoinesian Stage, Middle Pennsylvanian Series, from Mid-Continent No. 1 Collins well core, Ness County, Kansas (see Table 1; Fig. 2). A, Crinoidal limestone at 4395 feet. B, Calcareous siltstone at 4395 feet with finely broken fossil debris including pieces of holothurian plates, *Globivalvulina*, and other foraminifers. C, Sponge-spicule micrite at 4397 feet. D, Calcareous algal mudstone at 4398 feet with ground mass dolomitized and containing tiny foraminifers. E, Same bed as D, but included to show very large *Howchinia?* in relation to smaller forams occurring with it. F, Biopelmicrite at 4399 feet showing evidence of pressure compaction. Pelecypod shell fragments penetrate cylindrical thallus of *Komia*. Bar scale in F is one millimeter.

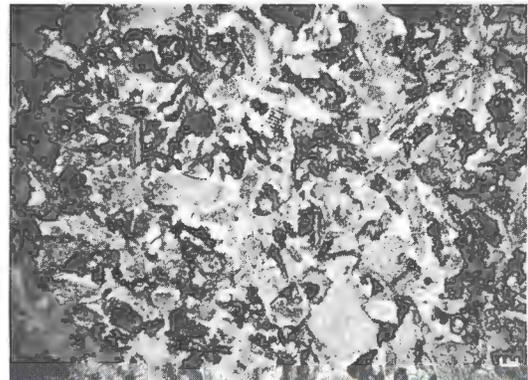
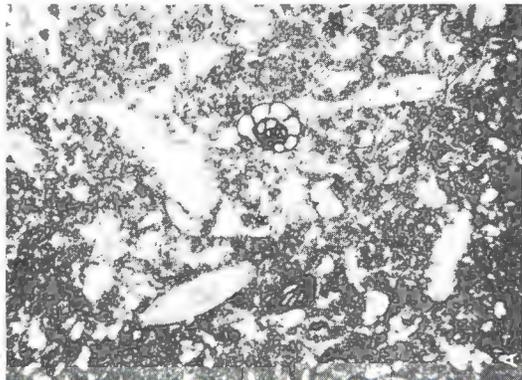
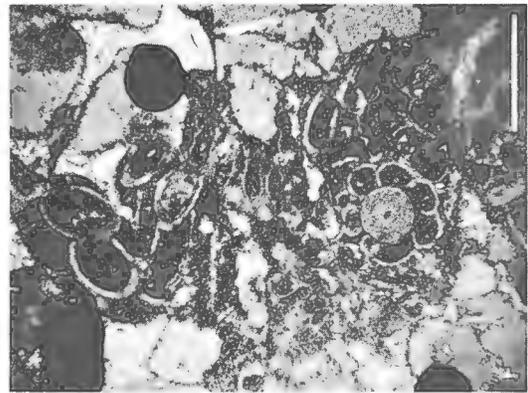
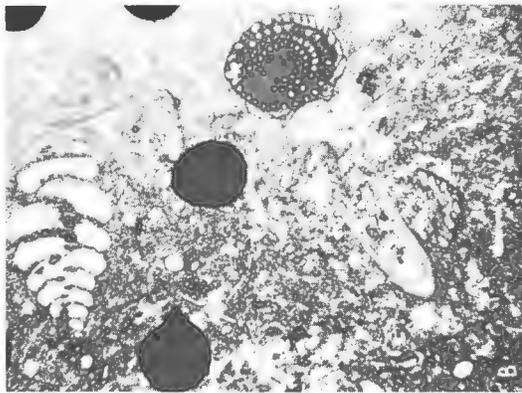
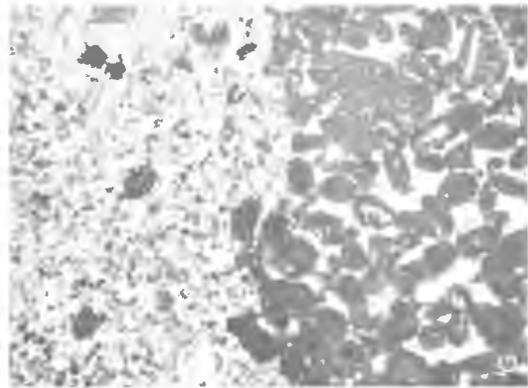
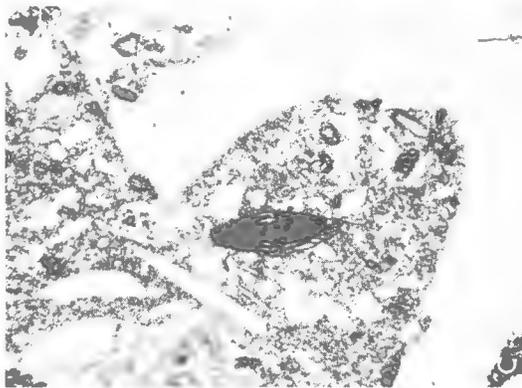
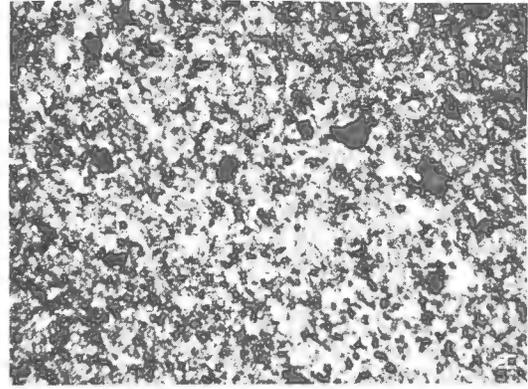
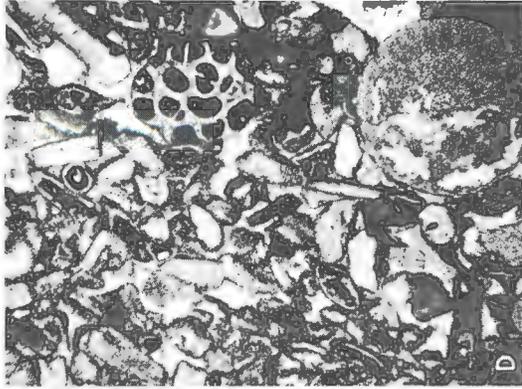
Depth Feet Meters	Description	Depth Feet Meters	Description	
4401 1341	Limestone, light-gray, algal, with <i>Komia</i> , <i>Trubertina</i> , <i>Endothyrax</i> , <i>Howchinia</i> , <i>Tetartax</i> , <i>Ammoliscus</i> , <i>Pseudowedgekinella</i> , <i>Plectogyna</i> , <i>Polytax</i> <i>labeci</i> Cushman and Waters, <i>Tetartax conica</i> Ehrenberg, "Pseudostachroides" of Petyk and Mamer 1972, <i>Millerella</i> , <i>Globivalvulina</i> , <i>Tartarula</i> , <i>Fusulina</i> of type occurring in lower Fort Scott Limestone, ostracodes, pelecypods, bryozoans, amber-colored plant material, pyrite, trace oil stain. *Shale, black. (Pl. 1, Figs. 15, 16; Pl. 2, Figs. 1-6, 8, 10, 12; Fig. 5, D, E)	4411 1345	4409-10 1344	Limestone, tan, more silty, less fossiliferous, fusulinids, brachiopods, ramose bryozoans, Siltstone, tan, calcareous, brachiopod and pelecypod fragments, spines, bryozoans. <i>Ibid.</i> , pyrite, ostracodes, small specks glauconite. And shale, dark gray-green, unfossiliferous, sandy or silty.
4402 1342	Limestone, light-gray, algal, <i>Komia</i> , highly recrystallized, <i>Archaeolithophyllum</i> ?, white porcelaneous encrusting forams, <i>Tubertina</i> , <i>Tetartax</i> , <i>Ammocertella</i> , <i>Globivalvulina</i> , <i>Plectogyna</i> , <i>Endothyrax</i> , <i>Planolites</i> , <i>Millerella</i> , <i>Ammoliscus</i> , bryozoans, pelecypods. (Pl. 2, Figs. 7, 11; Fig. 5, F)	4422 1348	4408 1344	Limestone, tan, more silty, less fossiliferous, fusulinids, brachiopods, ramose bryozoans, Siltstone, tan, calcareous, brachiopod and pelecypod fragments, spines, bryozoans. <i>Ibid.</i> , pyrite, ostracodes, small specks glauconite. And shale, dark gray-green, unfossiliferous, sandy or silty.
4403 1342	Mudstone, light to dark-gray, calcareous, molluscan, algal, <i>Komia</i> , <i>Globivalvulina</i> , <i>Millerella</i> , <i>Tetartax</i> , <i>Howchinia</i> ?, large inflated fusulinids, <i>Roschubertella</i> , <i>Pseudowedgekinella</i> , large "Plectogyna" bryozoans, brachiopod spines, broken fusulinids. (Pl. 2, Figs. 13-15; Fig. 5, G, H)	4427 1349	4407 1343	Limestone, light blue-gray, silty, fossiliferous, mud-streaked, <i>Globivalvulina</i> , <i>Tetartax</i> , fusulinids, brachiopods, ramose bryozoans, conodonts. †Shale, gray, with thin layer of black shale.
4404 1342	Limestone, gray, with molluscan mudstone, beds of coarse fossil fragments, "triamgenular" bryozoans, <i>Millerella</i> . And black shale with brachiopods and bryozoans.	4428 1350	4406 1343	Shale, dark-gray, calcareous, sometimes peaty, sandy, silty, fossiliferous, finely comminuted shell debris.
4405 1343	Siltstone/mudstone, lime, molluscan, black, with small limestone (Mississippian) clasts, traces brachiopods, small crustaceans, pyrite.	4429 1350	4405 1343	Siltstone, tan to gray to brown, lime, oil-stained, reworked high-spired <i>Tetartax</i> , <i>Pseudowedgekinella</i> , algae, <i>Komia</i> , recrystallized bryozoans, porcelaneous white encrusting forams, <i>Wedgekinella</i> (broken, reworked zone), <i>Eoschubertella</i> , <i>Bradigina magna</i> Roth and Skinner, <i>Gibberopsis</i> ? And mudstone, mottled, many stylonites. (Pl. 2, Fig. 17; Pl. 3, Fig. 6, B, C)
4406 1343	Shale, dark-gray, calcareous, sometimes peaty, sandy, silty, fossiliferous, finely comminuted shell debris.	4430 1350	4404 1342	Siltstone, gray-green to tan, more dense in places, with thin stringers and pockets of green silt, burrow mottled. <i>Ibid.</i> , more lime, finely silty beds, burrow mottled, bryozoans, rotund fusulinids, <i>Ammocertella</i> , <i>Glimacamma</i> , <i>Endothyrax</i> , <i>Tubertina</i> , <i>Globivalvulina</i> , <i>Tetartax</i> , bryozoans, algae, rock partially silicified. *Shale, black and some limestone. (Fig. 6, A)
4407 1343	Limestone, light blue-gray, silty, fossiliferous, mud-streaked, <i>Globivalvulina</i> , <i>Tetartax</i> , fusulinids, brachiopods, ramose bryozoans, conodonts. †Shale, gray, with thin layer of black shale.	4431 1349	4403 1342	Mudstone, gray-green, calcareous, silty, different from above, much pelecypod and brachiopod debris, finely broken. †Shale, green, with clumps of dense Pennsylvanian limestone imbedded.
4408 1344	Limestone, tan, more silty, less fossiliferous, fusulinids, brachiopods, ramose bryozoans, Siltstone, tan, calcareous, brachiopod and pelecypod fragments, spines, bryozoans. <i>Ibid.</i> , pyrite, ostracodes, small specks glauconite. And shale, dark gray-green, unfossiliferous, sandy or silty.	4432 1349	4402 1342	Mudstone, as above, increase in pyrite, burrow mottled. <i>Ibid.</i> , and *limestone, broken, and green shale.
4409-10 1344	Limestone, tan, calcareous, brachiopod and pelecypod fragments, spines, bryozoans. <i>Ibid.</i> , pyrite, ostracodes, small specks glauconite. And shale, dark gray-green, unfossiliferous, sandy or silty.	4433 1348	4401 1341	Mudstone, medium-gray, sometimes platy, fossiliferous, intensely burrowed, <i>Fusulinella stoufi</i> ? Thompson (reworked?). (Pl. 2, Fig. 16)
4411 1345	4410-11 1345	4434 1348	4400 1341	Siltstone, gray, calcareous, fragmental, fusulinids (reworked?) and sandstone, green to tan, calcareous, pelecypods, tiny white opaque forams, blastoids, echinoid spines. 4417 †Shale, gray, calcareous and fossiliferous. 4418, 5 †Limestone, dense, pale brown.

Figure 5. Thin sections of lithologic types occurring in the basal Marmaton Fort Scott Limestone (Desmoinesian Stage, Middle Pennsylvanian) in the Mid-Continent No. 1 Collins well core. A, Biomertite at 4399 feet showing dolomitization and replacement of original fabric by sparry calcite. Large foraminifer at top of photograph is *Tetartax*. Mass of very tiny plectogyrts in lower portion is "swarm" of asexual generation of individuals. Adult tests have disintegrated somewhat. B, Spicular algal biomertite at 4400 feet with replacement sparry calcite, showing two different species of *Tetartax*. C, Biomertite with replacement sparry calcite at 4401 feet, showing *Tetartax conica* Ehrenberg. D, Spicular algal biomertite at 4401 feet, with pyrite and silification of some penultimate debris. Algal blades lie above *Tetartax*. E, Spicular algal biomertite at 4402 feet. Algal blades and penultimate plates are replaced with sparry calcite. Crinoidal mass is dolomitized in part. Forclaneous encrusting forams in center of photograph. G, Spicular biomertite at 4403 feet. *Howchinia*? has secondarily enlarged wall. Brachiopod spine lies below it. Remaining fossil debris is penultimate and molluscan. H, Algal biomertite at 4403 feet, from near base of Fort Scott Limestone. Algal material and groundmass are replaced by sparry calcite. *Ammoliscus*-*cus* in center of photograph. (Dark objects are India ink dots.) Scale is 0.1 millimeter.

SOIL-DERIVED BASAL CONGLOMERATE OF THE CHEROKEE

Pebble conglomerate or breccia with sand grains, glauconite, iron (bright red), gray-green silt.



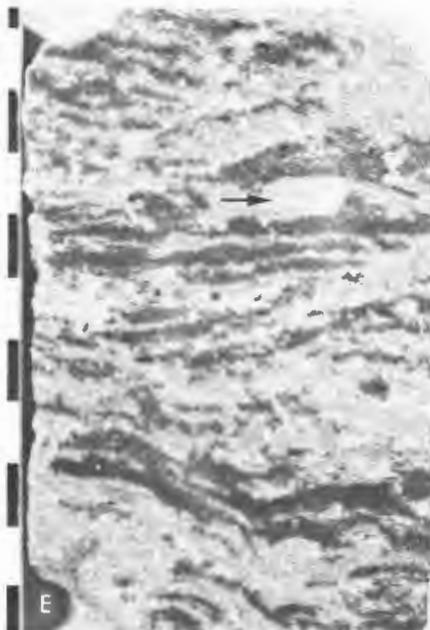
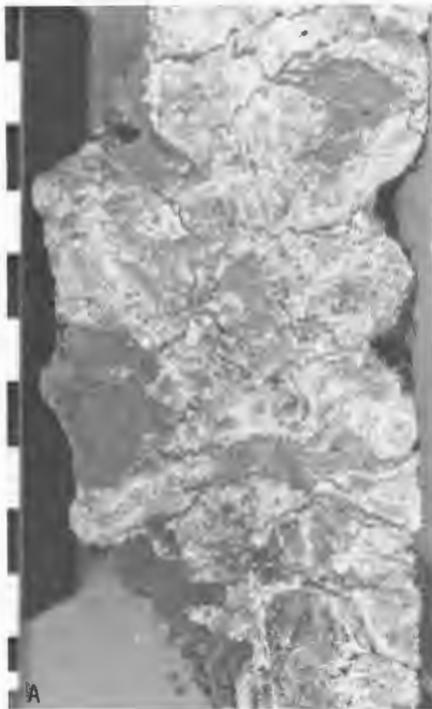


Depth Feet Meters	Description	Depth Feet Meters	Description
4431 1351	*Limestone, slight oil show. (No core recovery.)	4456 1358	<i>Ibid.</i> , more coarse-grained here, glauconitic shale, subrounded to angular clear quartz grains, light-gray siltstone, dark-gray siltstone, gray-green shale, pyrite, iron stains from bright orange to yellow. (Fig. 7, C, D)
4432 1351	†Clay, green, lumpy. *Limestone, shale, sticky. (No core recovery.)	4457 1359	Conglomeratic sandstone, bright green, glauconitic, interbedded green glauconitic silt with quartz beds, enclosing pieces of crinoidal limestone. Seems to be no calcite in matrix. Cobbles of irregular calcite beds intermixed with finely crystalline crinoidal material.
4436 1352	Mudstones and siltstones, highly weathered, below paleosol, quartz grains, iron in different states of oxidation, from greens of reds, veins filled with calcite (see Fig. 7, B). Highly weathered limestone cobbles with <i>Idionothodus</i> from 4435-4440. (Fig. 7, A)	4458-60 1359	Sandstone, conglomeratic, jumbled mixture of structures and materials: light and dark banded, red iron stains, green glauconitic, Meramecian cobble at 4460 feet, algae, *Lime, sandy, and green shale. †4459 Gray, hard, hard, studded rock, calcareous. (Fig. 7, E)
4440-45 1353-55	Claystone, red and greenish-gray to deep purple, mottled, unctuous, slickensided, shinkage cracks, fractured and recemented with green clay, calcite veins. (Fig. 7, B) sided, weathered soil zone.	4461 1360	Siltstone or mudstone, gray-green, with finely disseminated pyrite, highly weathered piece of crinoidal limestone.
4446 1355	Shale, red and gray-green, mottled.	4462 1360	Sandstone conglomerate-breccia, with blocks of highly weathered, recrystallized, dolomitized, yellow-brownish-reddish crinoidal limestone, some glauconitic, also glauconitic shale in tiny crevices in limestone pieces, crinoid stems stained pinkish-orange by oxidized iron. Contains bed of pure green glauconitic sandstone with silt and pyrite. Reworked "alpha coral," holothuran sclerites, echinoid spines, bryozoans, <i>E. kleina</i> ? Woodland in clast of Chesteran age.
4447 1355	Shale, red and gray-green, mottled, "rotten."	4464 1361	Breccia-conglomerate, crinoidal beds in varying states of preservation mixed with sandstone-shale-glaucconite. One piece limestone nearly pure white, very well preserved. Most stained red or ochre-colored by iron.
4448-50 1356	Conglomerate in shale, "rotten," red and green mottled, reworked sandstone pebbles and Mississippi chert and limestone clasts. Angular sandstone, fresh minerals, gray, mostly medium-grained (angular to subrounded, some frosted), all sizes of quartz to silt sized. Very little to no CaCO ₃ glauconite, mica, sphalerite fragments. A thin zone coarse white sandstone, mottled with iron stain and with dark siltstone-mudstone stringers.	4466 1361	Cobble or boulder of white fossiliferous limestone, bryozoans, algae, brachiopods, corals, fossil pellets coated with algae (resemble oolites), crinoids, "alpha coral," <i>Lytopora</i> , <i>Stachaeoides</i> ? <i>Asphaltina cordillensis</i> Manet, <i>E. kleina</i> Woodland, conodonts. (Pl. 3, Fig. 2)
4451-55 1357-58	Sandstone conglomerate, mottled gray and tan, noncalcareous, burrow mottled, intraclasses of banded sandstone, green shale, Mississippi limestone clasts, angular and rounded to subrounded, frosted grains from coarse sand to fine silt. And sandstone, mottled light tan to pinkish-white to pale gray to dark-gray, weathered, burrow-mottled (Fig. 8), some large pieces quartz. 4451 †Clay, variegated, red, yellow, and green. (See Fig. 7, C, D)	4467 1362	Siltstone, green-gray, and yellow-brown fossiliferous limestone boulder/cobble, well-preserved, fenestellid bryozoans, "alpha coral," echinoderm plates.

FIGURE 6. Thin sections of lithologic types found in Mid-Continent No. 1 Collins well core, Ness County, Kansas (see Table 1; Fig. 2). A, Biomictic with sparry calcite at 4429 feet. Cherokee Group. Contains fusulinids, *Endothyranaella*, brachiopod spines. B, Spicular micrite at 4430 feet with sparry calcite and dolomite. Cherokee Group. Contains *Climacommia*, fusulinids, *Tetraxaris*, and *Tubervatina*. (Black objects are India ink dots.) C, Dolomitized spicular micrite at 4430 feet, with fossils replaced by sparry calcite. Cherokee Group. Fusulinid *Wedekindella* at center. D, Coarse-grained biopelmicrite at 4469 feet, composed mainly of pelmatozoan and bryozoan debris. At upper right is reworked, rounded, "alpha coral," so far found only in Meramecian beds. E, Medium-grained biomictic at 4485 feet, made up mostly of pelmatozoan debris. F, Biomictic at 4490 feet, with sparry calcite replacing pelmatozoan plates. Alga *Asphaltina cordillensis* Manet (Meramecian) is at center of photomicrograph. G, This seemingly represents the contact between the Meramecian and Chesteran beds in the area of the No. 1 Collins borehole. Highly weathered, pyritized oolite below, which is reworked into extremely fine grained Chesteran beds above, contains Meramecian endothyrids as centers. This section is from 4495-4509 feet. H, Fine-grained, high-energy deposit of Chesteran age at 4495-4509 feet. Contains some still-intact fragments of dark ooids of Meramecian age. These beds contain the endothyrid *Endothyra kleina* Woodland of Chesteran age. Scale in F is 0.5 mm.

Depth Feet Meters	Description	Depth Feet Meters	Description
4168-69 1362	Sandstone, with silt, grayish-green, pyrite, glauconite, fossiliferous limestone debris in it. Very thin beds of silt and shale, light to medium-gray. Limestone pieces with well-preserved fossils, light-gray, with brownish stain, with gray-green, silt-filled crevices, <i>Asphalmina</i> , bryozoa, brachiopods, horn corals, echinoid spines, small round sponges, profile reworked "alpha coral," pyrite, glauconite, iron stains. Thin bed of white siltified sandstone, fossiliferous, somewhat devitrified. (Fig. 6, D)	4170-72 1363	Sandstone, conglomerate, green, grading coarse to fine, with silt, short strings and lumps of glauconite, contains sandstone clasts, aggregates of pyritized spicules, bcc-kite rings.
4173 1363	<i>Ibid.</i> , with limestone cobbles, tan, fine-grained, fossiliferous, pyrite, few grains bright green glauconite, inclusions of coarse white sandstone.	4174 1361	<i>Ibid.</i> , completely recrystallized, siltified.
4175 1361	Siltstone/sandstone, mixed, green to gray to white, highly glauconitic, containing coaly fragments, pebbles and cobbles Mississippian (Warsaw) limestone, tan, algal-crinoidal-bryozoa, with grains and fillings of fossil cavities with glauconite, some bedding at right angle to core length.	4176 1361	Limestone cobbles and pebbles, tan, silty, medium-grained, slightly crinoidal, tiny bluish-green glauconite, in matrix of gray-green quartz silt with glauconite. Siltified pebble with "alpha coral" bryozoans, brachiopods, lithology like 4174, but better preservation.
4177 1365	Siltstone/sandstone matrix, clastic, with in-tracasts of tan, fossiliferous limestone and green silt, glauconitic shale, dark-gray shale, light-gray shale, "alpha coral," <i>Ibid.</i> , plus dark-gray mudstone, scarce fresh glauconite.	4178 1365	Limestone cobble or boulder, shattered, yellowish-gray, fine-grained, fragmental, massive bedded, with fine cross bedding, finely broken bryozoa debris, bedding at right angle to core length (Fig. 9, A).
4179 1365	Limestone cobble, tan, coarse-grained, stylolites, bryozoa, dolomitized, siltified glauconite, in matrix of gray-green siltstone with coarse quartz grains. Limestone pebbles and cobbles, brownish-gray, fossiliferous, stylolitic, in matrix of siltstone/sandstone, gray-green, with quartz grains, chalky chert.	4180 1366	Same as 4177, limestone in green-gray, white chert imbedded.
4181 1366	Limestone cobble/boulder, light tannish-gray, recrystallized chert. *Lime and shale. Limestone cobble/boulder, light tannish-yellow, dense, fine-grained, crinoids, bryozoans, "alpha coral," stylolites, glauconitic, scarce yellow-orange chert pieces. Lime-stone with fossil "hash," pyrite, contains reworked Meramec fossils (<i>Asphalmina</i> of Mamer), echinoid spines, conodonts, dasyclad algae, <i>Endothyra klinea</i> Woodland (Chesteran) in a matrix of siltstone/sandstone, pale gray-green. (Pl. 3, Figs. 3-5; Fig. 17)	4182 1366	Limestone cobble, tan, siltified, and massive recrystallized chert, yellow and red iron stain, some glauconite, limestone as in 4181, in matrix of siltstone/sandstone, gray-green.
4183 1366	Limestone cobble, brownish-tan, "hashy," fine-grained, tiny plectoyrids and endothyrids, <i>E. klinea</i> Woodland (in Chesteran cobble), <i>Archaeodiscus?</i> , reworked "alpha coral," echinoid plates, clear interstitial calcite, few grains glauconite, contains oil droplets. *Lime and chert. Conglomerate, sandy, with chert fragments.	4184 1367	Limestone cobble, yellowish-tan, fine-grained, stylolites, glauconite, medium-sandstone, Limestone cobble, tannish-pink, medium-sandstone.
4185 1367	Limestone cobble, silty, echinoid spines, fragmented "alpha coral," crinoid plates, bryozoa, ostracodes, some included glauconite. (Meramecian) (Fig. 6, E)	4186 1367	Limestone cobble/boulder, yellowish-tan, very fine grained, stylolites, <i>Loximolima</i> (Chesteran). (Pl. 3, Fig. 6)
4187 1368	Mudstone matrix, green and yellowish-tan, sandstone and limestone clasts, yellow-buff, sandstone, fine-grained, buff, noncalcareous, and limestone cobble, yellow-buff, medium-grained, bryozoa, brachiopods, crinoid plates, stylolites.	4188 1368	Limestone cobble, tan, coarse-grained, stylolites, bryozoa, dolomitized, siltified glauconite, in matrix of gray-green siltstone with coarse quartz grains.
4189 1368	Limestone pebbles and cobbles, brownish-gray, fossiliferous, stylolitic, in matrix of siltstone/sandstone, gray-green, with quartz grains, chalky chert.	4189 1368	Limestone pebbles and cobbles, brownish-gray, fossiliferous, stylolitic, in matrix of siltstone/sandstone, gray-green, with quartz grains, chalky chert.

FIGURE 7. Beds from Mid-Continent No. 1 Collins well core from paleosol and immediately below (see Table 1; Fig. 2). A, Extremely weathered argillaceous limestone (light gray) with secondary mineralization (iron) along stylolites, calcite veins, and ochre-colored limonite clay masses (dark gray), at 4136 feet. The Pennsylvanian conodont *Idiozithodus* sp. was recovered from this zone. B, Maroon, highly weathered, siltken-sided claystone, with cracks and crevices infilled with light-gray clay at 4140 feet. C, Sandstone/siltstone matrix beds at 4156 feet. Lighter colored masses are fine-grained sandstone with iron stain. The groundmass is coarse sand and silt with particles of white chert, glauconite, black flecks of coaly debris, and minute gray claystone intracasts. (Core photograph inverted.) D, Indurated sandstone matrix beds at 4156 feet. This is a uniformly fine-grained deposit of the type commonly used for whetstones. It includes some larger-than-average, rounded quartz grains and an occasional intracast of limestone (upper right). Dark masses are merely iron stain. E, Glauconitic sandstone matrix beds at 4160 feet with intracasts of horizontally reworked sandstone beds (arrow). Differential oxidation of iron produces apparent banding effect. F, Transposed, brecciated Mississippian limestone fragments infilled with sandstone and siltstone matrix beds containing irregular pieces of white chert, from 4175 feet. Darker limestone clasts (arrows) are from a different lithologic suite. Very light colored clasts are devitrified chert. Scale in all photographs is in centimeters.



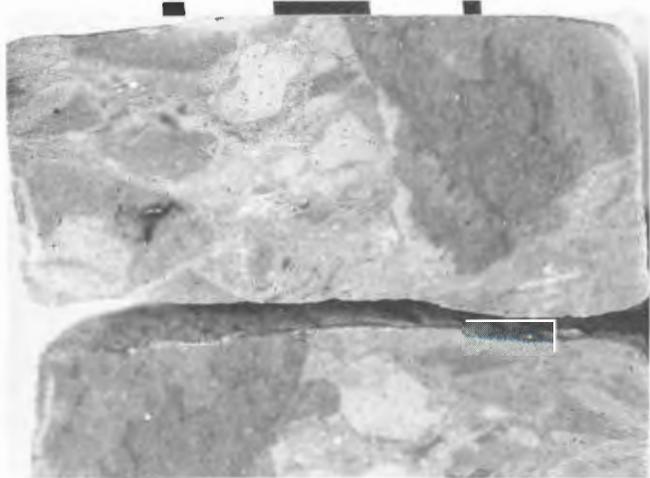
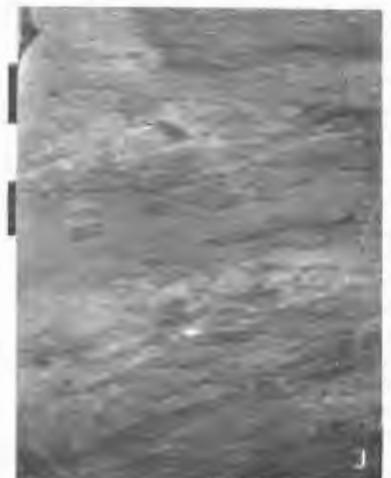
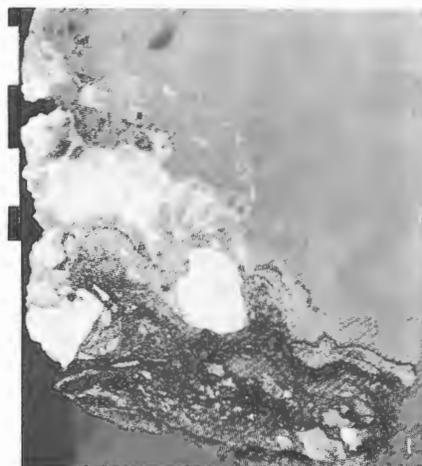
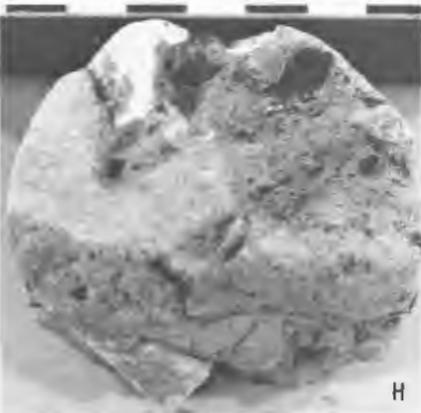
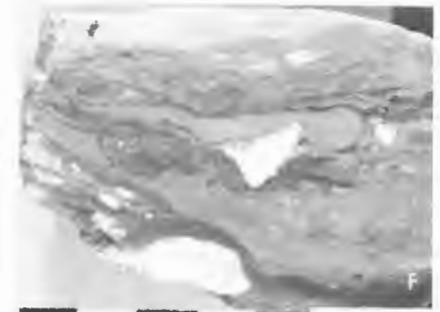
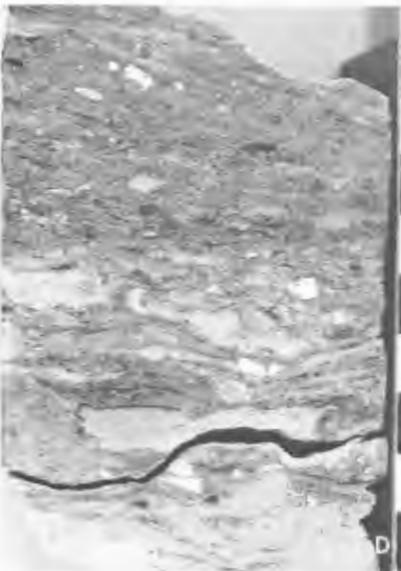
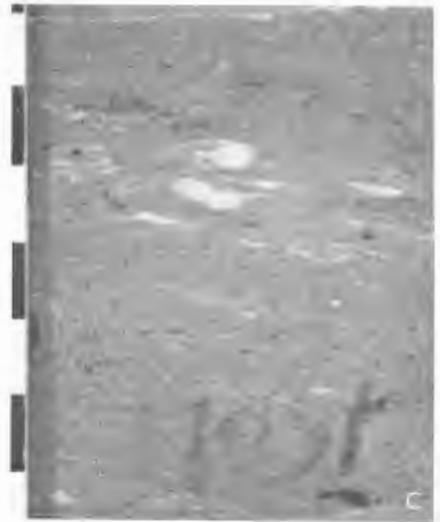


FIGURE 8. Matrix sandstone from breccia-conglomerate at 4455 feet. It is gray to tan, burrow-mottled, and contains tiny pieces of Mississippian limestone and white chert (see Table 1). Scale in centimeters.

Depth Feet Meters	Description
4493 1370	Limestone pebble, yellow-green, crinoidal, highly weathered, with pink iron stain, some algae, three kinds of fenestellid bryozoans, in siltstone/sandstone matrix, yellow-brown, and green, glauconitic, iron-stained where oxidized.
4493.5 1370	Limestone, broken by diagonal green clay seams.
4494 1370	Limestone pebble in breccia-conglomerate, yellow-green to white, algal-bryozoan, "alpha coral," highly weathered, stylolitic.
4495 1370	(Meramec) (Fig. 9, B) Limestone pebbles, reworked pyritized Meramecian fossils, some coarse-grained pieces, some very fine grained "hashy" pieces, some beds dolomitized, some pieces recrystallized to clear spar, stylolites, some reworked, contains tiny pockets and pyrite, sphalerite, contains iron-stained and seams of green silt, "alpha coral," bryozoans, small round sponges, horn corals, pelmatozoan debris. (Chesleran and Meramec) (Pl. 3, Fig. 7; Fig. 6, G, H)
4496-99 1370-71	Sandstone/siltstone matrix, conglomeratic, medium-gray, noncalcareous, with clasts of silt, glauconitic silt, medium-gray silt, carbonized plant remains (wood), with reworked white, chalky Mississippian chert, glauconitic shale and mudstone inclusions, and beds of mudstone, gray-green, with no debris. (Figs. 11, 12)
4499.5 1371	Limestone, pale brown, finely crystalline, with large pyrite lumps near base.
4500 1372	Siltstone, sandy, conglomeratic, medium-gray, as in 4496 above, with mudstone balls (Fig. 13). And limestone (Mississippian) cobbles, gray-brown, coarsely crystalline, fossiliferous, highly weathered, stylolitic.
4504-06 1372-73	Sandstone/siltstone, conglomeratic, gray, with clasts of siltstone, green glauconitic silt, black carbonized plant fragments (wood), and finely comminuted specks of black plant material (subbituminous coal), with white, chalky Mississippian chert and limestone fragments.
4507-09 1374	Siltstone conglomerate, gray to green-gray (Fig. 14), and mudstone, green, with white nodal, in quartz siltstone/sandstone matrix, mottled pink with iron-stain, coarsely crystalline, yellow-tan, limestone cobbles/pebbles, yellow-tan, (Fig. 10)
4491 1369	<i>Ibid.</i> , with vertical and horizontal seams of green shaly silt and white chalky chert.
4492 1369	Limestone cobbles/pebbles, yellow-tan, mottled pink with iron-stain, coarsely crystalline, in quartz siltstone/sandstone matrix, white to pale green.

FIGURE 9. Beds from Mid-Continent No. 1 Collins well core from 4479 to 4555 feet (TD) (see Table 1; Fig. 2). A, Well-preserved fine-grained Meramecian limestone cobble or boulder at 4479 feet, showing cross bedding that is perpendicular to core length (rotated). B, Partially weathered, fossiliferous Meramecian limestone cobble at 4494 feet. Stylolitic at outside edge of cobble. C, Somewhat indurated sandstone matrix beds at 4501 feet, containing stringers of pure white sandstone, finely comminuted coaly debris, devitrified white chert, blebs of glauconite, and intraclasts of gray claystone. D, Glauconitic, conglomeratic sandstone matrix beds at 4509 feet. Contains sand and silt stringers, small angular pieces of white chert and mixed Mississippian limestones, finely comminuted to stringers of coaly debris, masses of and finely disseminated glauconite, and varicolored claystone clasts. E, Sandstone matrix beds at 4510 feet, fairly fine grained, argillaceous, burrow mottled, and containing scarce very tiny fragments of Mississippian limestones. Most dark flecks are quartz grains; rarely they are coaly debris. F, Matrix beds at 4513 feet, a heterogeneous mixture of sand and silt grains, mudstone particles and stringers, devitrified white chert, glauconitic stringers, and blebs of pyrite. G, Dolomitized, decalcified, so-called "Warsaw" cobble showing oomoldic porosity at 4520 feet. Fenestellid bryozoan at left. H, Top view of core piece from 4522 feet, consisting of shaly matrix beds with disseminated glauconite, pyrite, and coaly debris and pieces of angular white chert. I, Dolomitized, desilicified, argillaceous Osganian sponge beds at 4530 feet, with enclosing matrix beds, as described above, still intact, with secondarily developed quartz masses, some with drusy quartz inclusions, and blebs of pyrite. Originally beds looked like those in J (core photo inverted). J, Argillaceous Osganian sponge-bed cobble or boulder at 4549 feet, containing dark sponge debris and white siliceous sponge spicules. Some cobbles and boulders are very well preserved and show the irregular, whipsy, high-angled bedding surfaces; others are severely weathered to a light tan color, demineralized, and dolomitized, with bedding structures obliterated. Scales in centimeters.



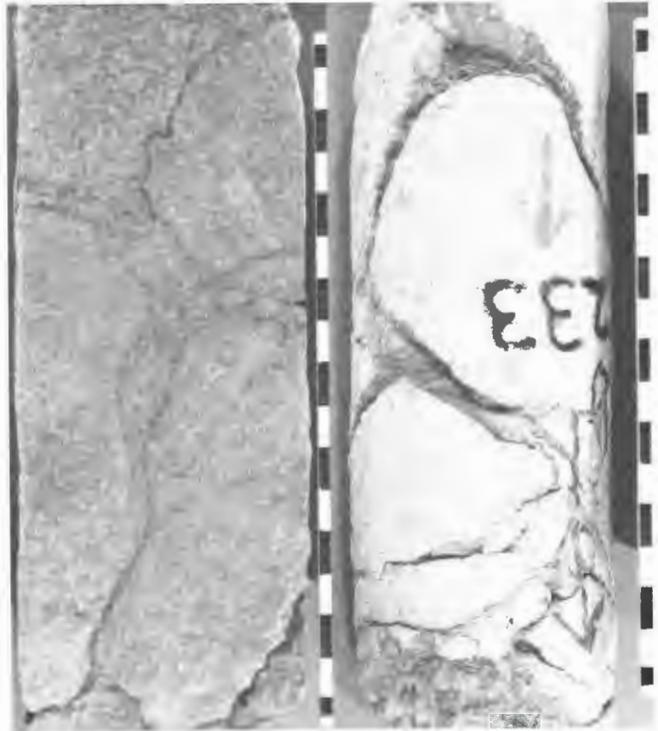


FIGURE 10. Fractured and weathered brownish-gray Mississippian cobble in breccia-conglomerate at 4491 feet. Cracks and crevices are filled with green clay residuum (see Table 1). Scale in centimeters.



FIGURE 11. Matrix conglomerate bed at 4496 feet composed of residual materials derived from weathering of Mississippian rocks (see Table 1). Black material is subbituminous coal fragments; white material is devitrified chert. Scale in centimeters.



FIGURE 13. Matrix bed from 4500 feet consisting largely of sand and silt grains with shale intraclasts and mud balls or pellets, coaly fragments, and pieces of white devitrified Mississippian chert (see Table 1). Scale in centimeters.

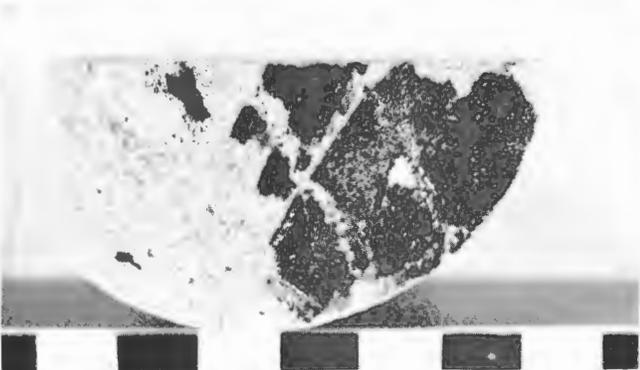


FIGURE 12. Large fragments of fossilized (coaly) wood in matrix sandstone at 4499 feet (see Table 1). Scale in centimeters.

Depth Feet Meters	Description
4527 1380	†Siltstone, hard, gray, very sandy, contains small pieces and larger chunks of limestone, green clay, white chert, miscellaneous materials, such as black plant fragments. Cobbles highly weathered tan to white. Mississippian chert, some surfaces water-worn, and Osagean sponge beds, in shaly, sandy conglomerate matrix, green and gray, irregularly bedded, and with relatively fresh Meramec cobbles of limestone. (Fig. 9, I; Fig. 16, C-G)
4536 1383	Siltstone, tan, fine-grained, dolomitized, with "whispy" sponge texture, sponge spicules. (Osagean)
4537 1383	Siltstone cobble, "billowy" sponge texture, some bedding at high angles. (Osagean)
4538-39 1383-84	Siltstone cobble, tan, dolomitized, highly weathered, chalky, sponge spicules. Dolomitized Salem "pseudo-oolite."
4540.5 1384	Clay, gray, soft.
4541-42 1384	Siltstone sponge beds, gray-white, dolomitized, highly weathered, chalky.
4543 1385	Limestone, blue-gray, dolomitized, tan, mottled with gray-green siltstone, also dolomitized, highly weathered, contains fenestrate bryozoans.
4544 1385	Limestone cobble, dark-gray, weathered, and limestone cobble, tan, dolomitized also, but not as highly weathered as 4543. (Fig. 16, H)

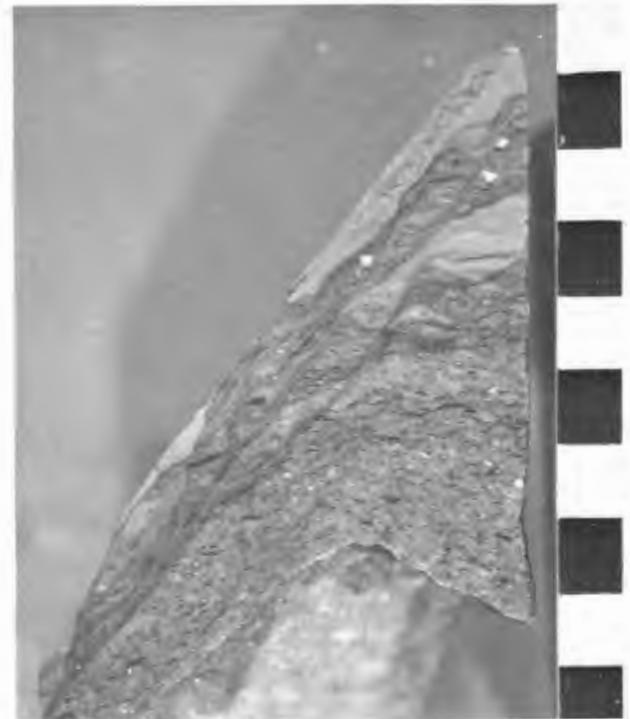


FIGURE 14. Matrix bed at 4508 feet (see Table 1). Sandy, silty residue deposited by fluvial action at high angle of repose around and between pebbles of Mississippian limestones derived from karst weathering terrain. White particles are devitrified white chert; dark particles and streaks are coaly material. Limestone cobble was occupying rounded depression seen at top of photograph. Scale in centimeters.

Depth Feet Meters	Description
4510 1375	Chalky chert fragments, pieces weathered Mississippian limestone, gray-brown, very fine grained, fresh plant fragments. And Mississippian chert with Salem endothyrids. (Pl. 3, Figs. 8-11; Fig. 6, G, H; Fig. 9, D; Fig. 14; Fig. 16, A, B)
4511 1375	Conglomerate, shaly, green-gray, with chert of white devitrified Mississippian chert with Salem endothyrids, sandy pockets, finely comminuted black plant debris, and highly weathered Mississippian limestone cobble. (Pl. 3, Figs. 8-11; Fig. 9, E)
4512 1375	Cobbles, white chalky Mississippian chert, in matrix of gray-green and gray shale, with fine particles carbonaceous wood fragments.
4513-27 1376-80	Sandstone, light-gray-tan, with small clasts of devitrified Mississippian chert pieces, and cobbles, and green shale, bedded, in matrix of light-gray shale and mudstone. Conglomerate, shaly, sandy, gray and green, irregularly bedded, with white, chalky Mississippian chert clasts and pebbles, weathered and water-worn surfaces, dolomite with fossil molds of fenestellid

FIGURE 15. Shaly matrix bed at 4522 feet (see Table 1) with high clay-mineral content, much disseminated pyrite, finely comminuted coaly debris, and angular pieces of rather fresh, white chert. Scale in centimeters.



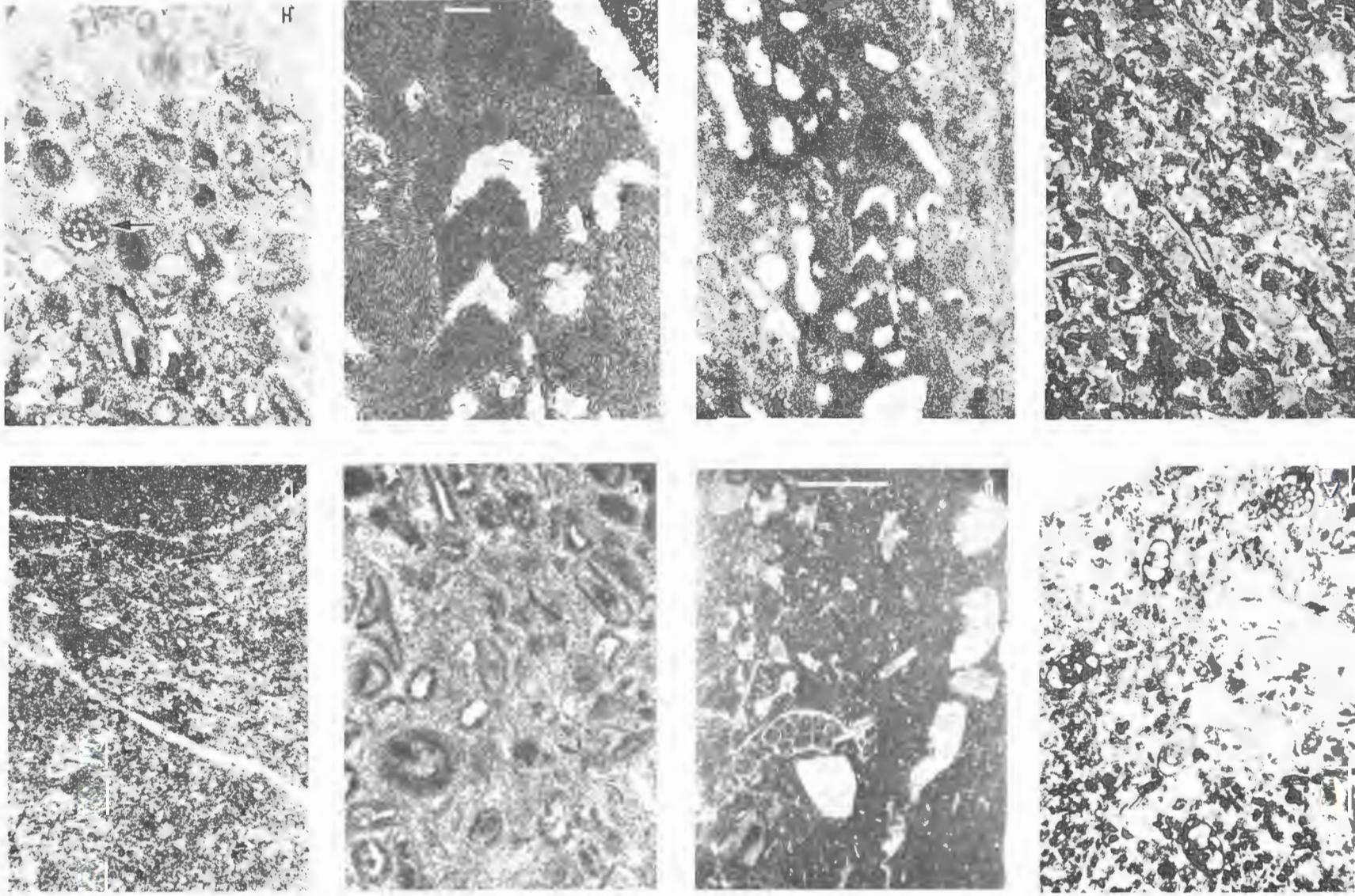


FIGURE 16. Thin sections of lithologic types found in cobbles and boulders in the Mid-Continent No. 1 Collins well core. Scale for A, B, C, D, E, F, H shown in B. A, Chert replacement of Meramecian limestone at 4509-4519 feet containing endothyrid foraminifers. B, Biomierite at 4509-4519 feet containing pelmatozoan and bryozoan debris. Scale is 0.5 mm. C, Dolomitized, partially silicified Meramecian biomierite at 4528-4538 feet. D, Sponge spiculite replaced by chert at 4528-4538 feet. Individual spicules can be seen in lower part of photograph. E, Biopelmierite with finely broken fossil debris, mostly pelmatozoan, at 4528-4538 feet. F, Dolomitized, vuggy biomierite at 4528-4538 feet. All calcitic pelmatozoan fragments are now voids. G, Enlargement of a portion of F showing coarsely granular dolomite in groundmass. Scale is 0.1 mm. H, Dolomitized, vuggy Meramecian biomierite at 4544-4555 feet. "Ghost" of endothyrid foraminifer can still be seen (arrow).

Depth Feet Meters	Description	Depth Feet Meters	Description
4545	Siltstone, mottled gray to tan, dolomitized, highly weathered, fossils obliterated by recrystallization. †Dolomite, has channels, as if weathered by surface waters. (Fig. 16, H)	4549	Siltstone, tan to gray, dolomitized, weathered, sponge spicules. (Fig. 9, J)
4546	Siltstone, streaked gray and tan, dolomitized, pyritized sponge spicules. (Osagean)	4555	Dolomitized Salem "pseudo-white," glauconitic. (Meramecian) (Fig. 16, M).
4547	Siltstone, tan to gray mottled, fine-grained, dolomitized, sponge spicules and debris, some pyrite, some spicules pyritized.		†Clay, gray, soft. *Lime and green shale.
4548	Siltstone, gray-black, "whispy" sponge bed.		LATE OSAGEAN STAGE
			Mississippian beds below core.

INSOLUBLE RESIDUES

Welch (1963) described insoluble residues from the limestone cobbles and pebbles from 4452 to 4555 feet (1357-1388 m) (white sandstone conglomerate beneath soil zone to total breccia-conglomerate core). Throughout this interval, the residues contain rounded, frosted sand and silt grains, often imbedded in green clay. Welch commented on the persistence of gray and green clay shale, white chalky chert, spicular chert and sponge spicules, drusy quartz, white beakite rings, and silicified fossils from limestones. Sphalerite and chalcocopyrite were noted at 4463-4467, 4476-4480, and 4495-4529 feet (1360-1362, 1364-1365, and 1370-1380 m), mostly as disseminated crystals or growing within drusy quartz vugs. The chert, quartz crystals, sand/silt grains, and sphalerite and chalcocopyrite represent reworking of material from Mississippian formations being eroded during the time of formation of the breccia-conglomerate.

MICROFAUNA IN CORE

All the microfossils in the No. 1 Collins core below about 4455 feet (1358 m) are Mississippian in age (see Table 1). They carry a Meramecian and Chesteran fauna including "Endothyra" *macra* and *E. symmetrica* E. J. Zeller (Pl. 3, Figs. 8-11) (Nodine-Zeller, 1972; Ebanks and others, 1977), *E. kleina* Woodland (Pl. 3, Figs. 3,5), *Asphaltina cordillensis* Mamet (Pl. 3, Fig. 2), *Lytropora*, *Stachaeoides*, *Eosigmothina* (Pl. 3, Fig. 6), *Archaeodiscus*, and an unnamed primitive coral (herein referred to as "alpha coral"), so far found only in the Meramecian (Pl. 3, Fig. 7). This coral is found extensively reworked in Chesteran beds in Kansas (Fig. 6, D) that also contain other reworked Meramecian fossils (Nodine-Zeller and Thompson, 1977). The secondary (replacement) cherts in the pebbles and cobbles contain excellently preserved Meramecian endothyrids (Pl. 3, Figs. 8-11; Fig. 16, A). A well-preserved dasyclad alga (Fig. 17) was found in a Chesteran cobble from 4481 feet (1366 m).

Although the family Dasycladaceae ranges from Cambrian to Recent, a similar form, questionably identified as a nanoporid, was illustrated by Brenckle (1977) from the Pitkin Limestone (Chesteran) in Arkansas. Both are extremely small forms. According to Thompson (1965), the limestone cobbles at 4435-4440 feet (1352-1353 m) (Fig. 7, A), just below the paleosol, yielded specimens of the conodont *Idiogoniatodus*, indicating either actual deposits in place or reworking of the earliest Pennsylvanian beds identified in the No. 1 Collins core. *Idiogoniatodus* ranges from uppermost Morrowan (Lower Pennsylvanian) into the Middle Pennsylvanian, but in this instance it must represent marine beds of possible late Morrowan or Atokan age, since Cherokee rocks have been positively identified from the Cherokee basal conglomerate to the base of the Fort Scott Limestone at 4404 feet (1342 m). It is possible, but not likely, that these beds could be earliest Cherokee in age. Interestingly, *Idiogoniatodus delicatus* Cunnell was found in basal Pennsylvanian beds tentatively identified as Atokan (Riverton) by Thompson in the No. 1 Swalley well core, Cherokee County, Kansas (Nodine-Zeller and Thompson, 1977).

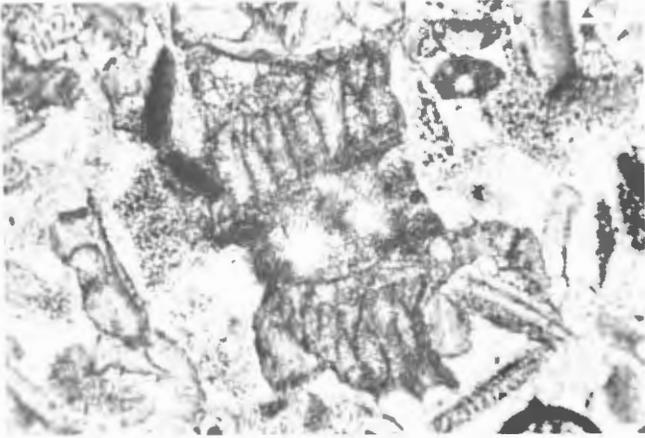


FIGURE 17. Dasycladacean alga (nanoporid?) from Chesteran limestone cobble at 4481 feet. Scale is 0.1 mm.

- PLATE 1. Microfossils in thin sections of core from the Mid-Continent No. 1 Collins well core, Ness County, Kansas. Scale for 1, 3, 5, 6-13, 15 shown in Figure 1 (0.1 mm). Scale for 2 shown in Figure 2 (0.1 mm). Scale for 4, 16 shown in Figure 4 (1.0 mm). Scale for 14, 17 (1.0 mm) shown in conjunction with photographs.
- Figure 1.—*Ozawainella* sp. Horizontal axial section, Fort Scott Limestone, 4395 feet (1340 m).
 2.—*Ammonovertella* sp. Showing initial chamber and convoluted test. Fort Scott Limestone, 4395 feet.
 3.—*Globivalvulina* sp. Fort Scott Limestone, 4395 feet.
 4.—*Fusulina* sp. Compares favorably with type occurring in basal Marnaton elsewhere. Fort Scott Limestone, 4398 feet (1340 m).
 5.—*Howchinia?* sp. This is a large form. Walls have been recrystallized. Fort Scott Limestone, 4398 feet.
 6, 10.—*Ozawainella* sp. Primitive form. Fort Scott Limestone, 6, 4398 feet; 10, 4400 feet (1341 m).
 7, 8, 11.—*Tuberculina* sp. Note wall structure in 8. Fort Scott Limestone, 7, 4399 feet (1341 m); 8, 4398 feet; 11, 4400 feet.
 9.—Plectogyrid foraminifera. Fort Scott Limestone, 4399 feet.
 12.—*Tetralaxis conica* Ehrenberg. Fort Scott Limestone, 4399 feet.
 13.—*Millerella* sp. Fort Scott Limestone, 4398 feet.
 14, 17.—*Konia abundans* Korte. 14. Horizontal section, 17. Longitudinal section. Fort Scott Limestone, 4399 feet.
 15.—*Endothyranella* sp. Fort Scott Limestone, 4401 feet (1341 m).
 16.—*Fusulina* sp. Fort Scott Limestone, 4401 feet.

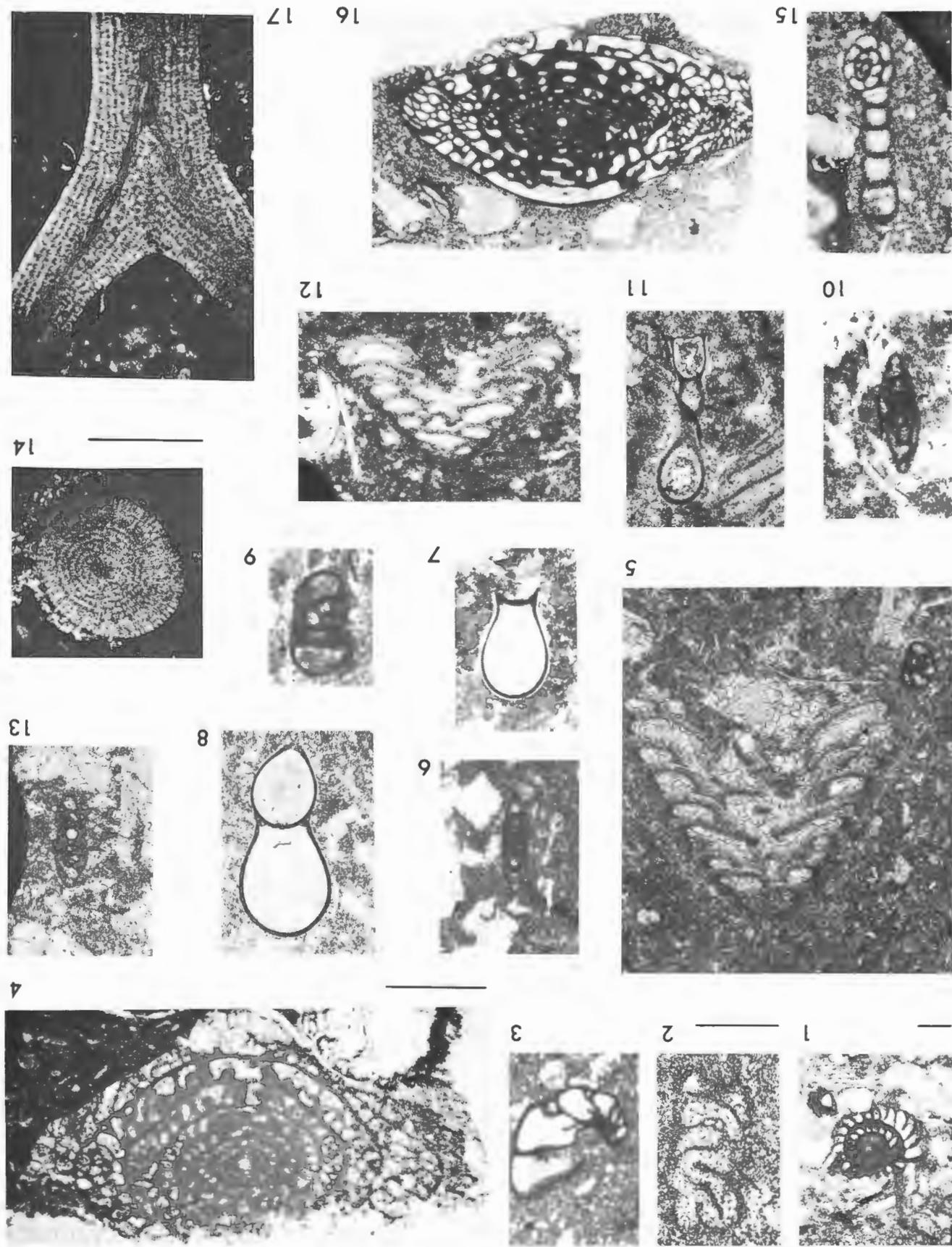
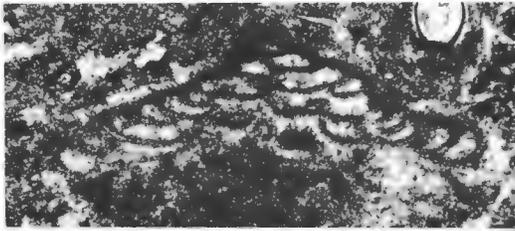
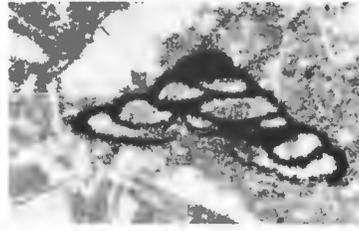


PLATE 2. Microfossils in thin sections of core from the Mid-Continent No. 1 Collins well core, Ness County, Kansas. Scale for 1-15 shown in 2 (0.5 mm), scale for 16 and 17 shown above 17 (1.0 mm).

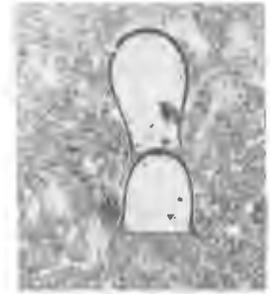
- Figure 1.—*Tetraxia* sp. Fort Scott Limestone. 4401 feet (1341 m). *Tubercina* at upper right.
 2.—Pyritized *Tetraxia* in dolomitized groundmass with clear calcite spar. Fort Scott Limestone. 4401 feet.
 3.—*Tubercina* sp. Fort Scott Limestone. 4401 feet.
 4,13.—*Houchinia*? sp. A very large foraminifera with recrystallized walls. Fort Scott Limestone. 4, 4401 feet; 13, 4403 feet (1342 m).
 5.—*Parlandia* sp. A tubular, attached form showing initial chamber. Fort Scott Limestone. 4401 feet.
 6.—*Indolithyanella* sp., with "neck" broken off. Fort Scott Limestone. 4401 feet.
 7,11.—*Planulincoluta* sp. A porcelaneous encrusting foraminifera. 7, bottom (planar) view; 11, side view. Fort Scott Limestone. 4402 feet (1342 m).
 8.—Juvenile *Wedekindella*? Fort Scott Limestone. 4401 feet.
 9.—*Lochubertella*? sp. a, Skewed sagittal section; b, tangential section, showing chomata; round object is small *Tubercina* sp. Fort Scott Limestone. 4403 feet (1342 m).
 10.—*Tetraxia conica* Ehrenberg. Fort Scott Limestone. 4401 feet.
 12.—*Tetraxia conica*? Ehrenberg. Fort Scott Limestone. 4401 feet.
 14,15.—*Pseudowedekindella* sp., axial sections. Fort Scott Limestone. 4403 feet.
 16.—*Fusulina stoufi*? Thompson, axial section. Cabanis Formation? Cherokee Group. 4413 feet (1345 m). *F. stoufi* has been described from the Mercer and Boggs Limestones (Pottsville) of Ohio.
 17.—*Bradyina*? sp. Cabanis Formation? Cherokee Group. 4430 feet (1305 m).



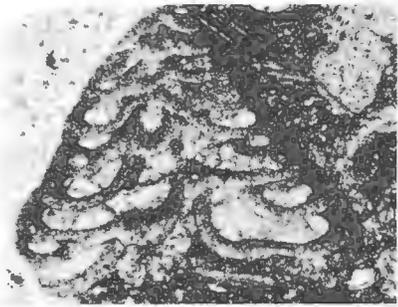
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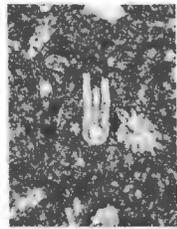
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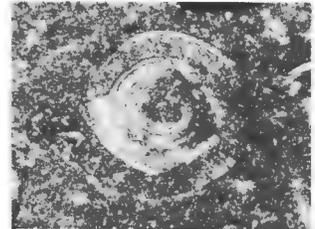
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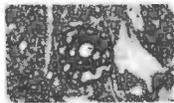
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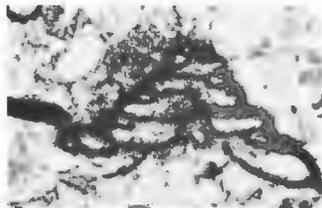


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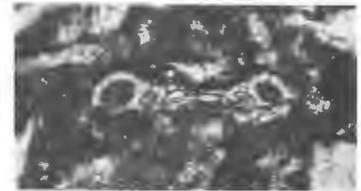


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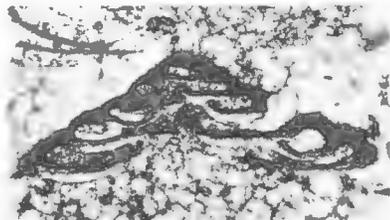
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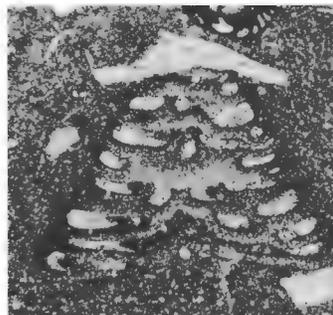
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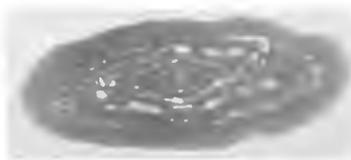
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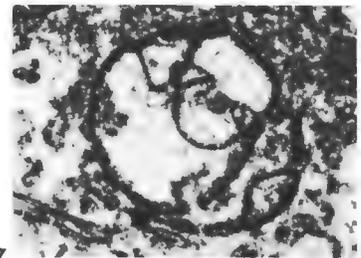
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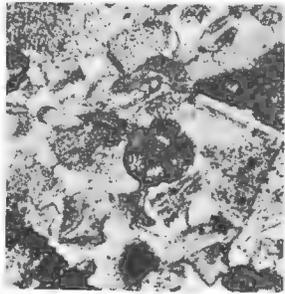
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PLATE 3. Microfossils in thin sections of core from the Mid-Continent No. 1 Collins well core, Ness County, Kansas. Scale for 1, 2, 4, and 7 shown in 2 (1.0 mm); scale for remaining specimens shown in 6 (0.5 mm).

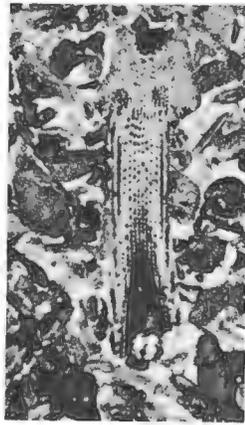
Figure 1.—*Wedekindella henbesti?* (Skinner). Axial section. Cherokee Group. 4430 feet (1350 m).
2.—*Asphallina cordillerensis* Maner. Breccia-conglomerate, probable Meramecian cobble. 4466 feet (1361 m).
3,5.—*Endolhyra klinea* Woodland. Breccia-conglomerate, Chesteran cobble. 4481 feet (1366 m).
4.—Echinoid spine in cobble. 4481 feet.
6.—*Fosimolina* sp. Horizontal axial section. Breccia-conglomerate, Chesteran cobble. 4486 feet (1367 m).
7.—"Alpha coral," showing single septum in central corallites. Breccia-conglomerate, highly weathered Meramecian cobble. 4495 feet (1370 m).
8-11.—"*Endolhyra*" *symmetrica* E. J. Zeller, beautifully preserved in chert. 8,10, horizontal axial sections; 9,11, vertical axial sections. Breccia-conglomerate, Meramecian cobble. Repeated throughout interval 4509-4519 feet (1374-1377 m).



2



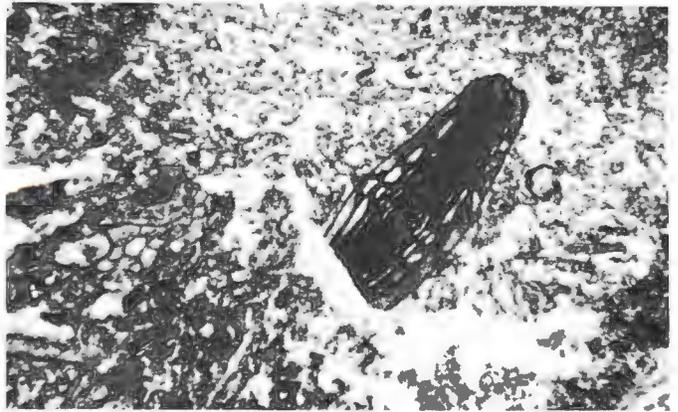
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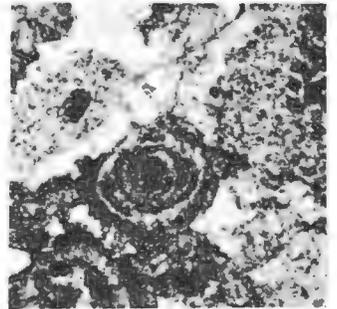
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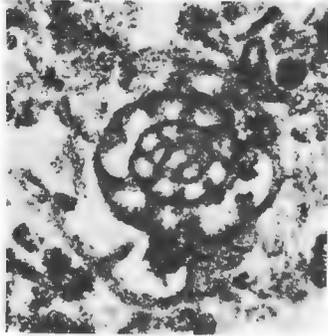
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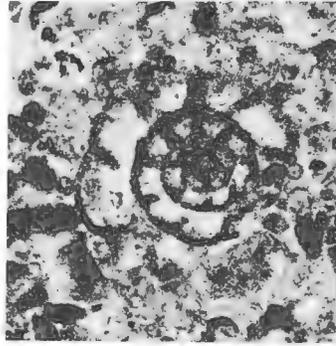
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6



11

derived clasts, or collapse breccia, all reworked and interbedded with fluvial deposits from stream channels crossing the area. A break in sedimentation is indicated by the white sandstone conglomerate. The same materials are found as residues in this conglomerate as in the breccia-conglomerate unit below, meaning that the same Mississippian beds were still being reworked during this time. But the materials are more finely divided in the sandstone conglomerate and are dominated by quartz sand and silt. After this terrestrial unit was deposited, there was a transgression of the seas and a uniform shale bed was deposited, represented by the 10 feet (3 m) of deeply weathered red clay in the soil profile. Above this, limestones of late Morrowan or early Atokan age were deposited, represented in the core by the "rotten," leached limestone (cobbles and pebbles? or in-place weathering?) of early Pennsylvanian age containing *Idiogoniatodus*. After the marine beds were laid down, there was uplift, or regression of the sea, formation of a soil, marine transgression, with subsequent reworking of the soil zone rubble into the Cherokee basal conglomerate, and a second, and perhaps third, regression and transgression of the Cherokee seas, with reworking of previously deposited materials. Finally a stable marine regime was established, as evidenced by normal-marine faunas in the Pennsylvanian rocks encountered at the top of the core.

PALEOGEOGRAPHIC AND DEPOSITIONAL HISTORY

The existence of a karst terrain in early Pennsylvanian time is suggested by the considerable relief of the eroded Mississippian surface, the presence of the early Pennsylvanian breccia-conglomerate of variable thickness that contains highly weathered, demineralized limestone and chert; silification of limestone beds; smooth, water-worn surfaces on the limestones, sometimes exhibiting solution fluting; inclusions of reworked cave-derived travertine in Pennsylvanian limestones; and the preservation of a well-developed soil profile at the top of the conglomerate.

Paleokarst topography has been reported in the subsurface developed on the Cambro-Ordovician Arbuckle dolomite in a number of places in Kansas (Walters, 1946); on Ordovician-Silurian rocks in the Williston Basin in Montana (Rogh, 1967); and on Mississippian limestones in Kansas, Oklahoma, Missouri, Colorado, Montana, Wyoming, and the southwestern United States (on Meramecian-St. Louis Limestone and on Chesteran beds) (Mapel and others, 1979; Clupper, 1978; Orgren, 1979; McKnight and

Beds from 4430 feet (1350 m) (Cherokee basal conglomerate) and above are Pennsylvanian in age (Fig. 2; see Table 1 for faunal lists). Fusulinids were used to interpret age of the rocks in this sequence. No *Eostaffella* or *Millerella* of the type *M. marblensis* Thompson, occurring in the Morrowan, were found in the limestone cobbles in the core. Nor was *Profusulinella*, mainly restricted to Atokan-age rocks (Thompson and Zeller, 1956), recovered from this core. *Millerella marblensis* has been reported (Nordine-Zeller, 1977) from the Boyd Shale (Morrowan) of Arkansas and in core samples from the Kearny Formation (Morrowan), Kearny County, Kansas (Thompson, 1944). It is likely that the specimen of *Idiogoniatodus*, of a highly indistinguishable mineralogical composition, came from extremely altered, highly weathered and mineralized limestone from the phreatic zone, and that no delicate calcareous microfossils would have survived such extreme weathering processes. In addition, well-preserved remnants of Morrowan or Atokan age were probably not encountered at the particular location of the core, since this zone of reworked Pennsylvanian cobbles (or limestone beds?) is relatively thin.

Desmoinesian (Cherokee) rocks in the core (4430-4405 ft.; 1350-1343 m) contain *Fusulinella*, *Eoschubertella*, *Ozawunella*, *Fusulina* (Pl. 2, Fig. 16), *Wedekindella* (Pl. 3, Fig. 1), and *Pseudowedekindella* (Pl. 2, Figs. 14-15). The genus *Fusulinella* first appears in rocks of Atokan age and ranges into the Desmoinesian (Thompson, 1945). The presence of *Wedekindella* indicates that these beds are lower Desmoinesian (Cherokee) in age (Thompson, 1945). The distinctive, highly inflated, and highly evolved species of *Fusulina* (Pl. 1, Fig. 4) at 4398 feet (1341 m) is similar to those occurring in the basal Marmaton Fort Scott Limestone. Thompson (1945, p. 41-42) stated:

Fusulina is believed to have been derived from *Fusulinella* and they overlap stratigraphically. In fact, forms occur near the top of the Desmoinesian that seem to be intermediate in nature between *Fusulinella* and *Fusulina*. *Millerella*, *Pseudostaffella*, *Eoschubertella*, "Nankinella," and *Wedekindella* are associated with *Fusulina* in the Desmoinesian.

INTERPRETATION OF CORE

From analysis of the data available, it is clear that rocks represented by the sequence in the No. 1 Collins core are highly weathered pebbles, cobbles, and boulders derived from an older karst limestone terrain that was created by the block faulting of an anticline referred to earlier. These pebbles, cobbles, and boulders are intermixed with fresh, more angular talus-

(Desmoinesian Stage). In places, there is a thin basal conglomerate in the lowermost shale of the Krebs that contains irregularly rounded chert and limestone pebbles one to three inches (3-8 cm) in diameter. These pebbles are derived from weathered, karsted Mississippian formations. Orogenic movements affecting this part of the Midcontinent began in early Atokan time and occurred intermittently into early Desmoinesian time.

The karst surfaces exposed in recent times in Colorado and the western United States can be studied in outcrop. These ancient karst terrains had developed considerable relief and were extensive throughout the whole North American continent. Lee (1940, p. 76) stated that in Sumner County (south-central Kansas) the weathered zone (basal Pennsylvanian conglomerate) is 163 feet (50 m) thick and extends downward to within 70 feet (21 m) of the top of the Chattanooga, cutting out the entire Mississippian section. In another well the weathered zone is 150 feet (46 m) thick and extends to within 35 feet (11 m) of the top of the Chattanooga. He pointed out that considerable caving was indicated by the presence of dark shale with the chert and that all the wells in which there is an important development of white (highly weathered) chert are on topographically high areas near the axis of the Nemaha fault structure and, thus, would have been subjected to karsting. The relationship between white, deeply weathered chert and uplifted, extremely deeply weathered Mississippian rocks has been noted by Fowler and others (1934), Lee (1940), McKnight and Fischer (1970), and Maslyn (1977b).

DeVoto and Maslyn (1977) described a karst topography in Colorado with breccia-filled sinkholes up to 250 feet (76 m) deep and cave systems developed on top of the Leadville Limestone (Mississippian) in late Mississippian or earliest Pennsylvanian time. There, as in the No. 1 Collins core, thoroughly weathered white chalky chert and earthy limestone fragments are believed to represent material from exposed sinkholes or cave-filling deposits reworked into the conglomerate. The 125 feet (38 m) of conglomerate in the Collins well rests upon eroded late Osgan beds. In southwestern Colorado the breccia-conglomerate rests upon the eroded surfaces of Osgan and Kinderhookian to Devonian beds. Throughout the Four-Corners area similar shales and siltstones with limestone-chert boulders rest upon a paleokarst surface of Mississippian or Devonian limestone. In northern New Mexico, Smith (1980) described karst formation on Mississippian limestone with intermittent and irregular deposition of clastic debris on the karst surface, creating a series of com-

Fischer, 1970; Merrill and Winar, 1958; Roberts, 1979; Mallory, 1979; Armstrong, 1979; McKee, 1979; Smith, 1980).
 Clupper (1978, p. 44) described the irregular Chesteran (post-Pitkin) karst surface in northeastern Oklahoma:

... the unconformable surface is marked by solution cavities. Plate 7, Figure 2 shows a large Pitkin block, about 10 feet in maximum observed dimension, which is overlain by a Morrowan sandy limestone. Bedding planes within the Pitkin block are practically vertical relative to the overlying beds which drape over the block. [See Fig. 9, A this paper.] Large solution vugs are . . . present on the surface of the boulder. Filling these vugs is a sandy carbonate. Numerous rock burrows are present on the surface of the block, which are filled with silt and sand. All these features strongly suggest that localized areas within southern Adair County stood relatively emerged during post-Pitkin, pre-Morrowan time.

Orgren (1979) found a similar post-Pitkin karst erosional surface in northeastern Oklahoma (Wagoner, Cherokee, and Muskogee counties) developed on sub-aerially exposed Chesteran rocks. Cobble and boulder-sized fragments of the Pitkin are here incorporated within a matrix of calcareous sandstone. The sandy matrix material contains *Declinognathodus noduliferus* and *Adeotognathus latus*, conodonts thought to be of Morrowan age. The Morrowan sands and fragments of Pitkin limestone are found lying in solution cavities 16 feet (5 m) below the Chesteran-Morrowan unconformity surface. A fragment of breccia-conglomerate shown in a photograph (Orgren, 1979, pl. 12, fig. B) is strikingly similar to those found in the No. 1 Collins well core (see Fig. 7, F).

McKnight and Fischer (1970) described uplift and development of a karst terrain at the end of Chesteran time in the Tri-State lead-zinc district. They stated (p. 56):

In places the whole [Chesteran] series was removed by erosion before the deposition of the Pennsylvanian, in other places, near-maximum thicknesses of Chester strata are preserved in sharply defined structural basins and stupps that formed in the interval between the deposition of the Chester and the beginning of Pennsylvanian deposition.

The Chesteran in that area is a maximum of 200 feet (61 m). The thickness of Morrowan rocks is about 80 feet (24 m). They consist of sandstones with interbedded shales and some limestones. The sandstones contain fossil plants and some of them are colored black with inclusions of finely carbonaceous or bituminous material. Morrowan rocks are truncated by the overlying Atokan rocks. The Atokan, in turn, is truncated and overlapped by rocks of the Krebs Group

porosity vary rapidly from well to well, as they do in the breccia-conglomerate in the area of the No. 1 Collins.

CORRELATIONS USING ELECTRIC LOGS

The continuous core in the No. 1 Collins from the bottom of the basal Pennsylvanian conglomerate to well above Marmaton beds made it possible to pick with great accuracy the contacts between the conglomerate, the paleosol, the basal Cherokee conglomerate, the Cherokee beds, and the Marmaton (base of the Fort Scott Limestone) using a combination of lithology and microfossils. Because there is no electric log for the No. 1 Collins, a cross section (Pl. 4) was constructed tying in the No. 1 Collins stratigraphic units by extrapolation with those identified on well logs in the vicinity. The easternmost log, Cities Service Oil Company No. 1 E. Moore B (Schaben Field), was included because formations had been identified previously by Hilpman (1965) and because the Fort Scott Limestone "kick" was readily identifiable and easy to correlate with that on the accompanying logs on the traverse. In analyzing data from Davies' lithology logs, top cards, drillers' logs, and electric logs, and comparing them with directly visible information from the well core, several facts became apparent:

- (1) There was considerable disagreement in the "pick" of the top of the Fort Scott; therefore, its thickness, varied considerably;
- (2) The position of the Cherokee and, therefore, the top of the "Pennsylvanian conglomerate" and its thickness were not uniformly agreed upon;
- (3) No unanimity existed in picking the top of the "Mississippi solid"; and
- (4) There was considerable disagreement on the age of the Mississippiian beds below the "top pick."

DISCUSSION

From the paleontological and physical evidence, it is clear that deposits in the No. 1 Collins well represent a breccia-conglomerate of earliest Pennsylvanian age (or possibly early Cherokee?) in which is incorporated an inchoate mixture of pebbles, cobbles, and boulders of Mississippian limestone and chert in a silty, sandy, shaly matrix that is largely derived from end-stage weathering debris from Mississippian formations ranging from Osagean to Chesterian in age (455-4475 ft.; 1388-1364 m). On top of the breccia-

plex depositional episodes. Latest Mississippian orogenic movements, as it was in the Midcontinent area. Earliest Pennsylvanian worldwide sea-level change (Vail and others, 1977; Kerr, 1980) with regression of seas, coupled with coincident geotectonic episodes, resulted in the karsting of widespread solution-prone Mississippian formations.

The karst landforms found in the United States in the Mammoth Cave (Kentucky), Carlsbad Cavern (New Mexico), and Missouri cavern systems represent ancient exhumed karst terrains, augmented by further karstification since the time they were uplifted and exposed to surface weathering on the present land surface. Solely modern karst surfaces, such as those on the Gulf Coast and in the Caribbean, are believed to have developed since some time in the Tertiary (Quinlan, 1972).

One aspect of the breccia-conglomerate that has been puzzling is the unoxidized state of the iron in the matrix beds. None of the sandstones, siltstones, shales, or mudstones in the Collins conglomerate are red. They are either white, tan, light to dark gray, or green (unoxidized), and in no case did the drillers' log record anything but green and gray shale in this interval. Most sinkholes or paleokarst caves are filled with red (oxidized) siltstones or clays (Bretz, 1942; Maslyn, 1977; Walters, 1946; Jennings, 1971), and these would remain red even though they were reworked at a later time. Either the accumulation of the breccia-conglomerate was so rapid that in-place weathering and oxidation did not have a chance to take place, or the breccia-conglomerate was deposited rapidly in a nearshore quasi-saline site under conditions that allowed little or no oxidation to take place. It is possible that the dissemination of carbonaceous and coaly debris that is so prevalent in the section may have prevented the oxidation of iron in the breccia-conglomerate. Alternatively, it may be that red, oxidized surface deposits were stripped off by erosive action of the encroaching seas. Two possible depositional modes for the breccia-conglomerate are inferred from evidence accumulated in this study: (1) deposition in fault-controlled channels upon a karst surface and (2) deposition adjacent to prominent sea cliffs, perhaps in a quasi-marine environment. In both, additional deposits of riverine origin would be intercalated. Guinan (1966) described a well in western Texas penetrating a paleogeomorphic trap in Wolfcampian conglomerate of chert and limestone pebbles in a shaly matrix, which produces from an irregularly developed channel deposit crossing a faulted anticline. Total thickness of the conglomerate and net

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- conglomerate, denoting a pronounced change in sedimentation, is a sandstone conglomerate containing very small pieces of recognizable reworked Mississippian limestone (4475-4448 ft.; 1364-1356 m). On top of this is highly weathered, unctuous red clay shale containing highly weathered limestone cobbles and boulders (4448-4435 ft.; 1356-1351 m). It is in this unit that a lower Pennsylvanian cobble was found that contained *Idiogoniatodus*. At 4435-4430 feet (1352-1350 m) is the paleosol, with "rotten" limestone, colored bright orange and yellow by hematite and limonite, and transected by calcite veins. On top of the soil zone are red and green claystone and sandstone with small pieces of Mississippian and lower Pennsylvanian? limestone and chert in a thin basal conglomerate reworked by transgressing Cherokee seas (4430-4429 ft.; 1350 m). And finally the deposition of the transgressive-regressive marine limestones and shales of middle to late Cherokee age, containing a fauna typical of the Verdigris Limestone, was begun. Lying above the Cherokee beds is the Fort Scott Limestone of the Marmaton Group (4404-4385 ft.; 1342-1337 m), the highest Pennsylvanian identified in this study. These highest beds examined contain tiny, angular clasts of Mississippian limestones, cherts, and Cherokee? shales, and travertine presumably derived from reworking of former karst-surface cave deposits.

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